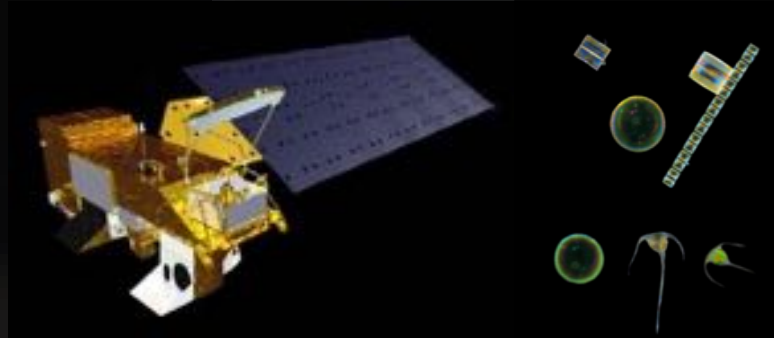


Ocean Biogeochemical Cycles

Bob Brewin



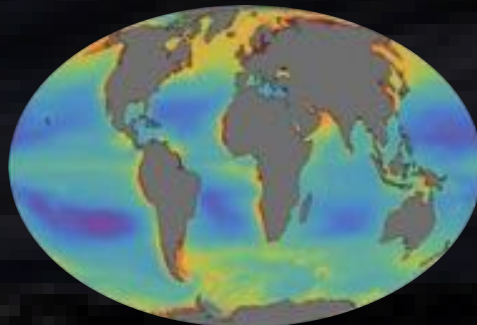
2003-2010



2010-2019



2019-present



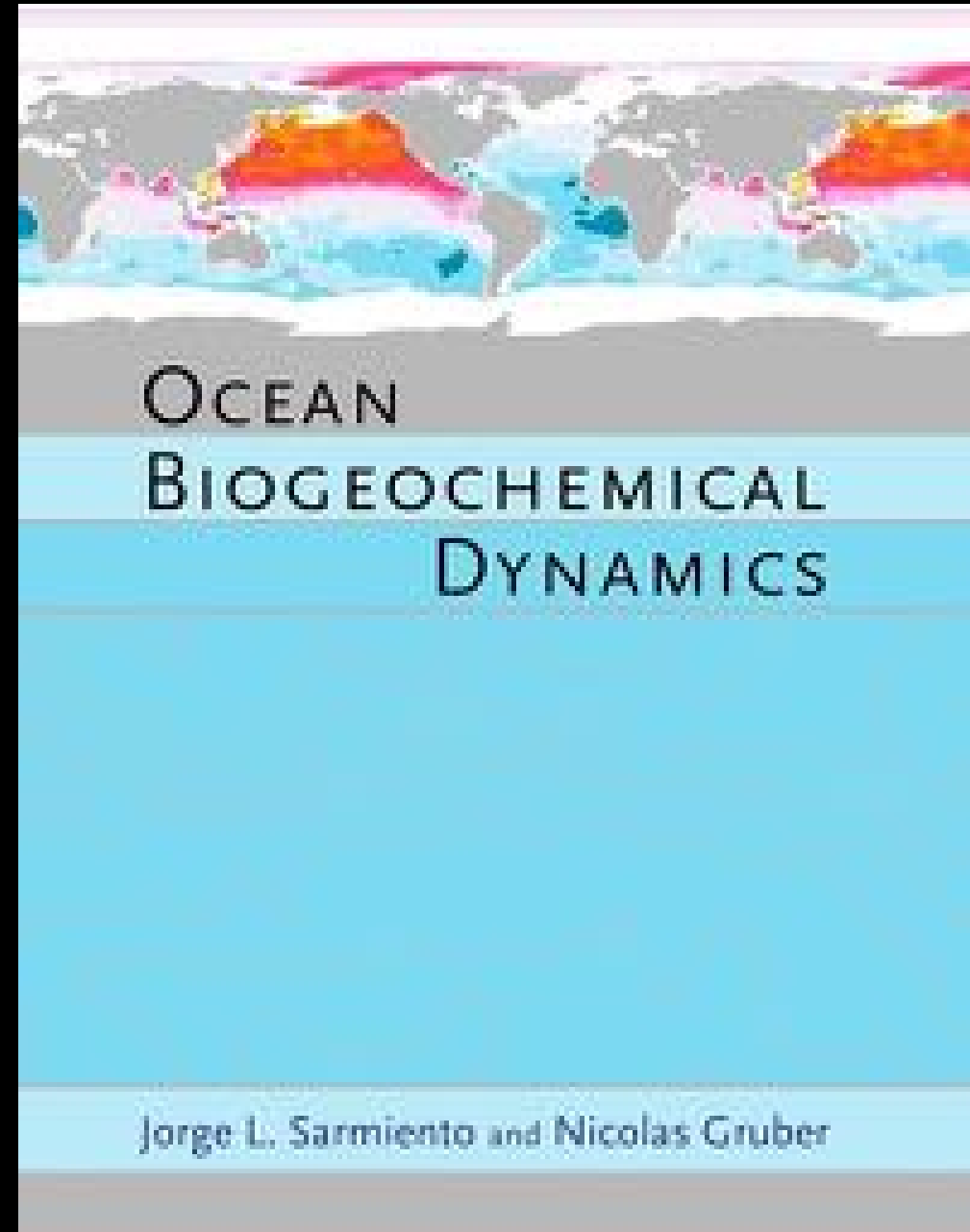
Lecture outline

1. Introduction to marine biogeochemical cycles
2. The Water cycle
3. The Redfield Ratio
4. The Carbon cycle
5. The Ocean Biological Carbon Pump
6. The Nitrogen cycle
7. The Iron cycle

Biogeochemical cycles are the pathways in which chemical substances (including elements and compounds) move through and within the environment.

Chemical substances can be combined and broken apart. But chemical elements can not be created or destroyed. Therefore, elements are recycled or stored (sequestered) somewhere on the planet.

The chemical substances can move between the marine environment and other components of the Earth System, such as atmosphere, ocean floor, terrestrial realms and polar regions.



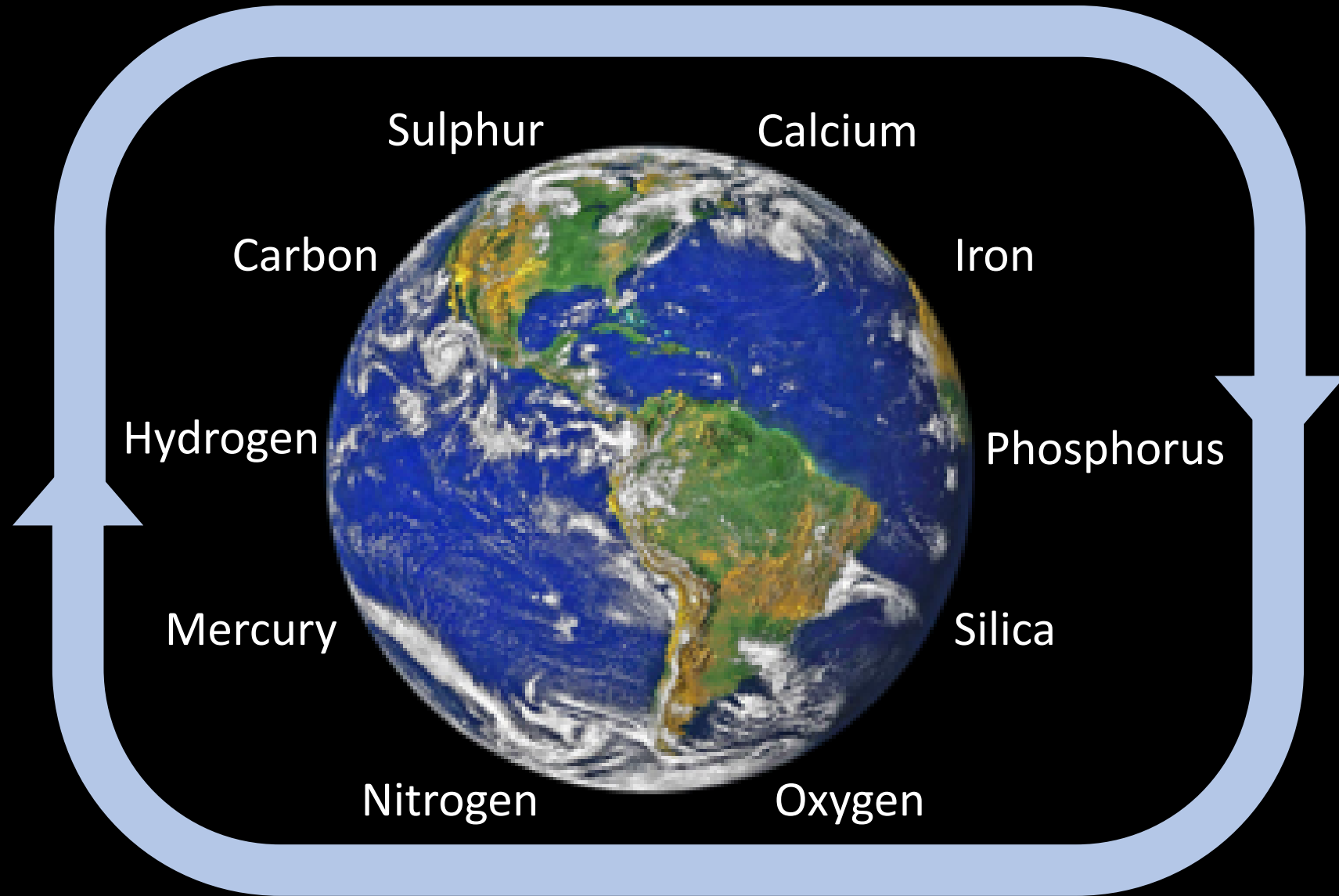
Biogeochemical cycles are driven by a range of forces, including metabolic processes in organisms (**bio**); geological processes, e.g. the movement of the Earth (**geo**); and chemical reactions among the substances (**chemical**).

Hence the name **biogeochemical** cycles (**BGC** cycle).

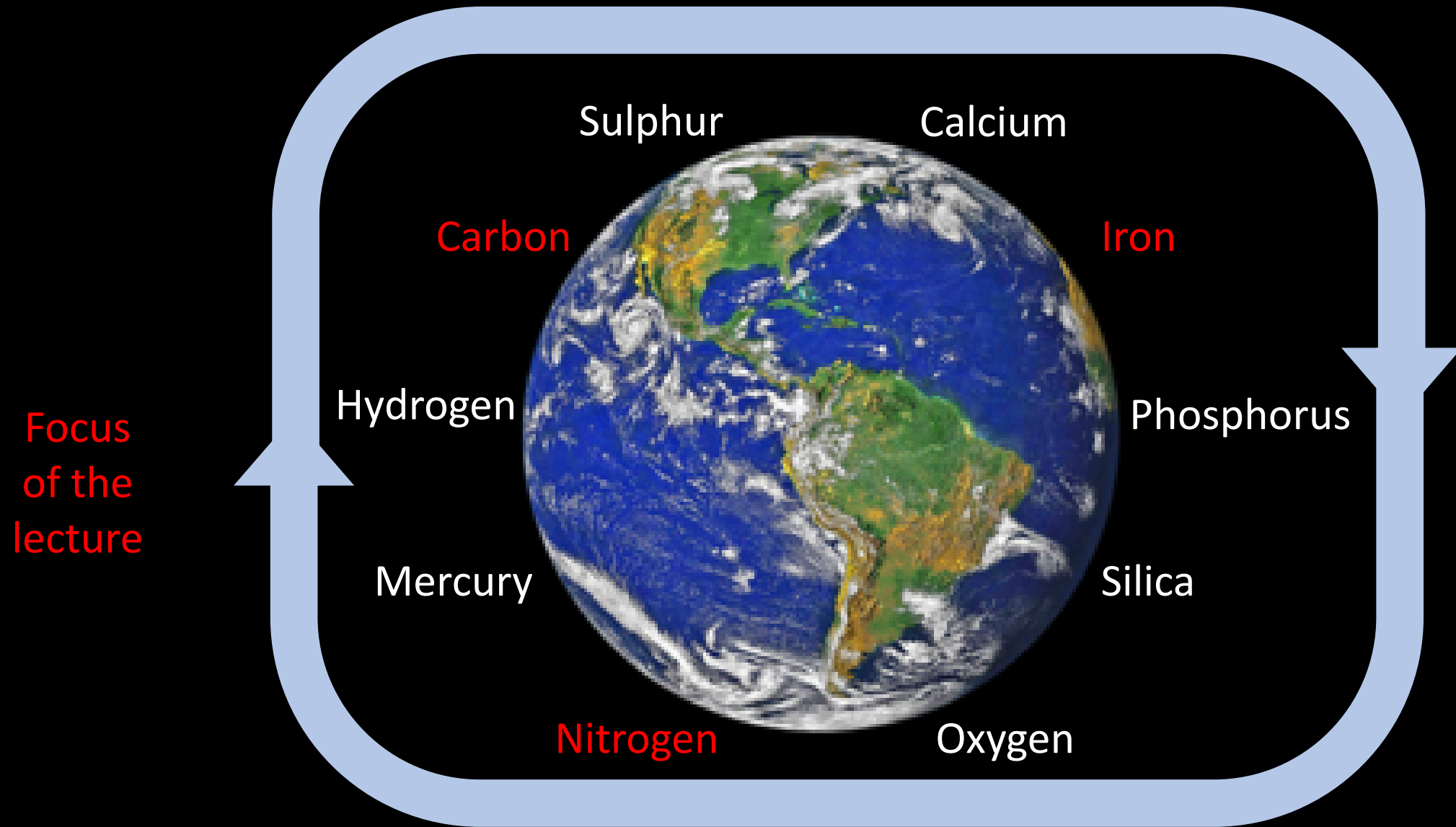
Marine organisms (microorganisms in particular) are critical to the functioning of many **BGC** cycles.

Image: Sardet (2015) Plankton wonders of drifting world. Univ. Chicago Press

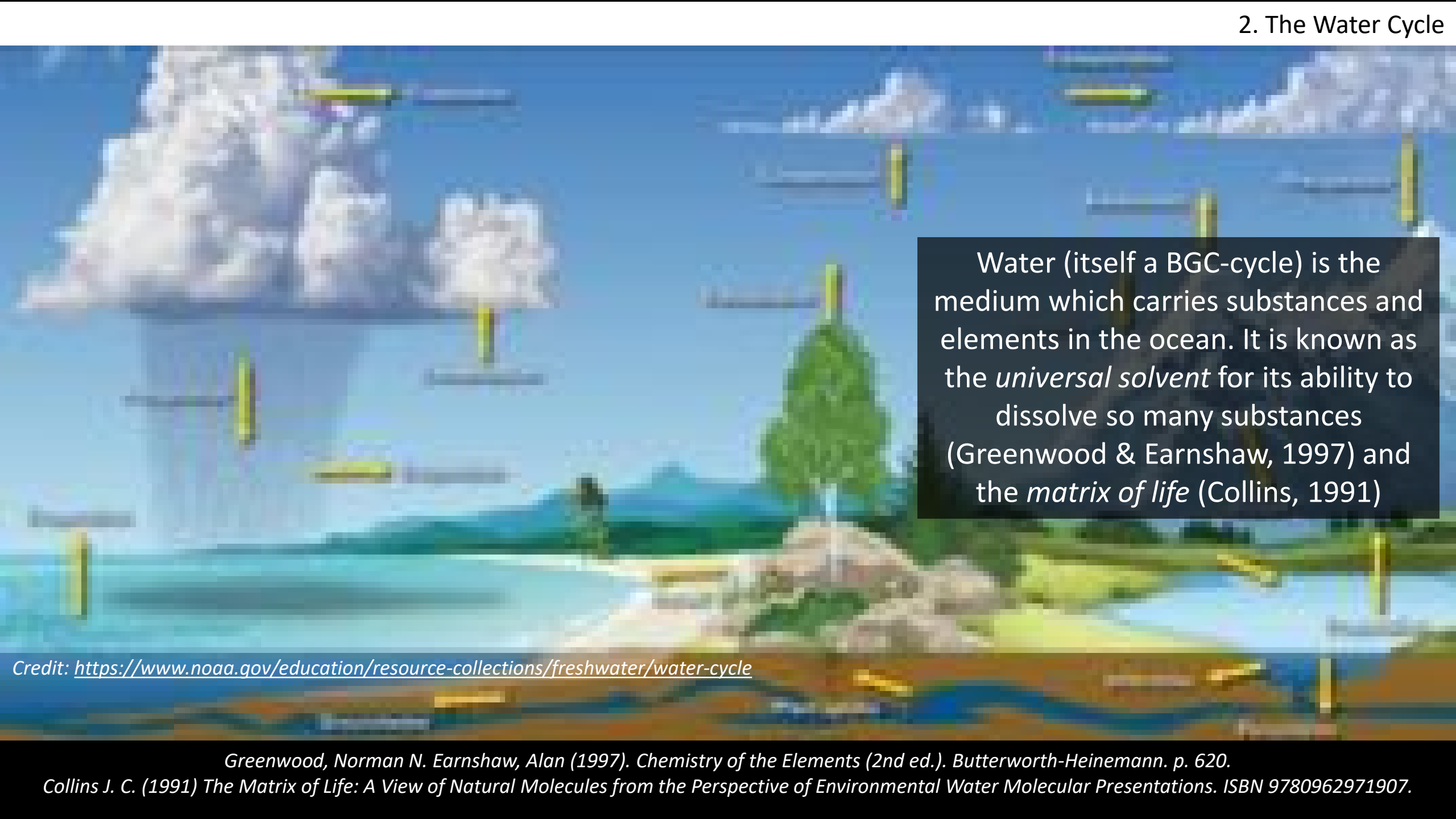




Many of these elemental cycles are intimately linked



Many of these elemental cycles are intimately linked

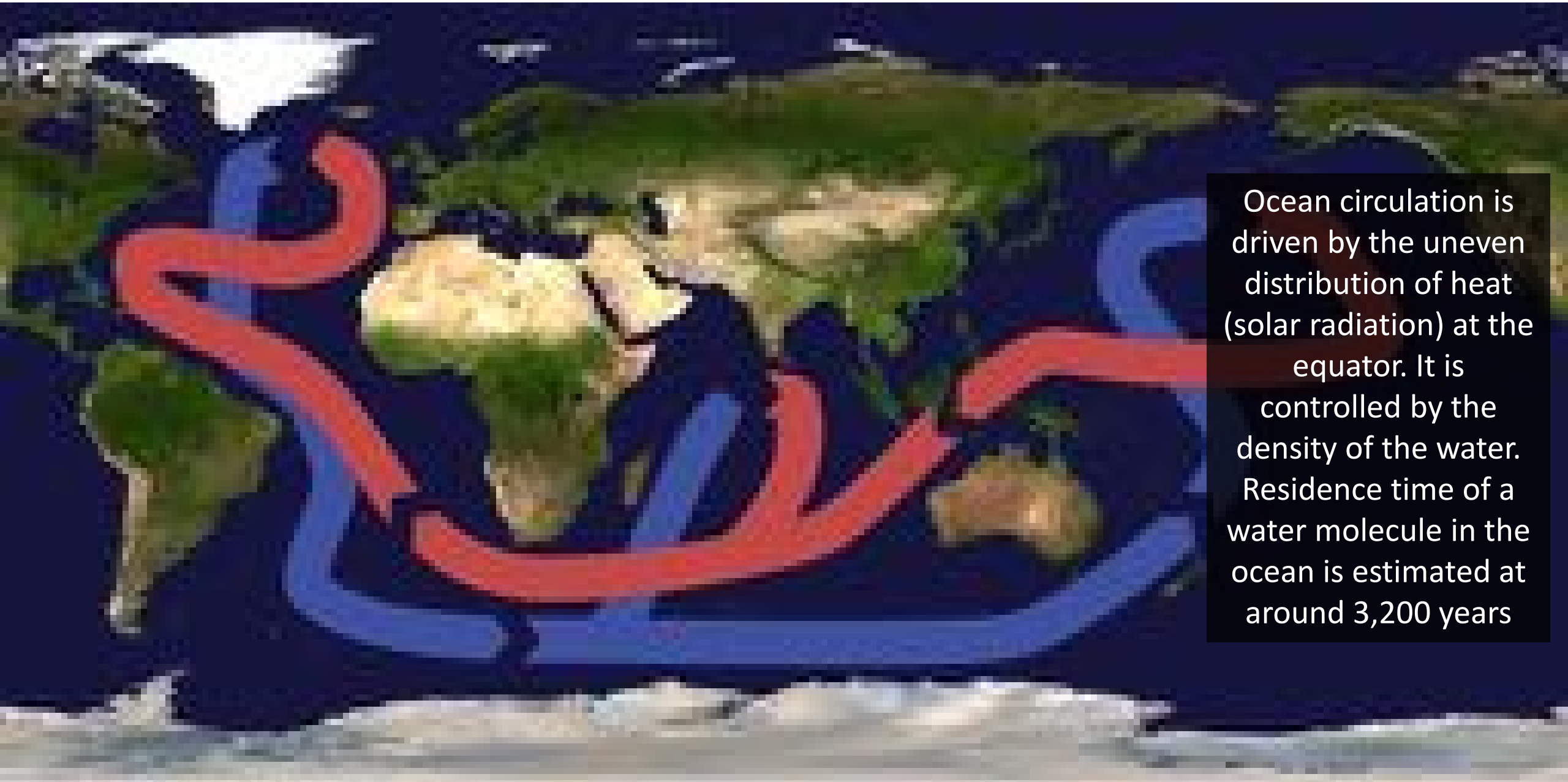
A scenic landscape featuring a bright blue sky with large, fluffy white clouds. In the foreground, there is a body of water reflecting the sky. A single, tall, green coniferous tree stands prominently on a small island or peninsula in the middle ground. The overall scene is peaceful and natural.

Water (itself a BGC-cycle) is the medium which carries substances and elements in the ocean. It is known as the *universal solvent* for its ability to dissolve so many substances (Greenwood & Earnshaw, 1997) and the *matrix of life* (Collins, 1991)

Credit: <https://www.noaa.gov/education/resource-collections/freshwater/water-cycle>

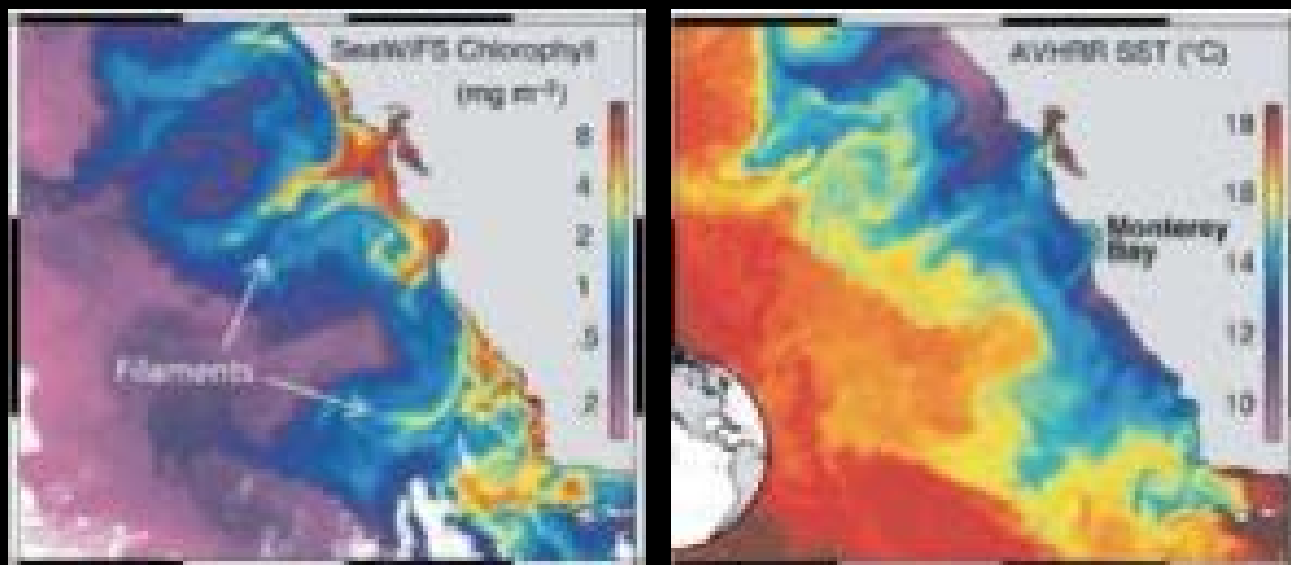
Greenwood, Norman N. Earnshaw, Alan (1997). *Chemistry of the Elements* (2nd ed.). Butterworth-Heinemann. p. 620.

Collins J. C. (1991) *The Matrix of Life: A View of Natural Molecules from the Perspective of Environmental Water Molecular Presentations*. ISBN 9780962971907.

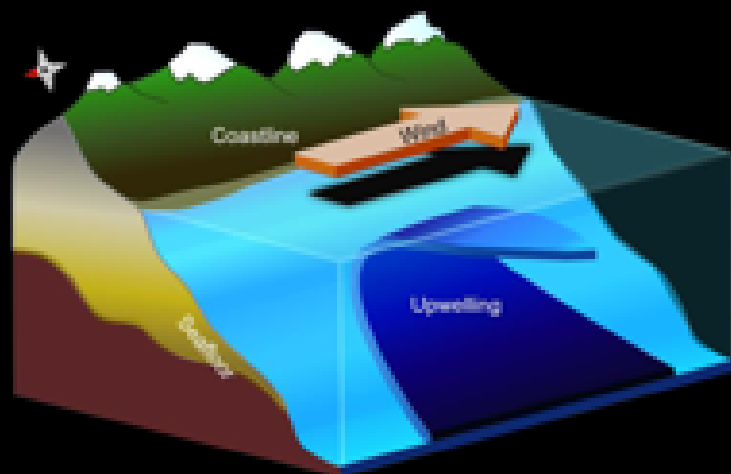


Ocean circulation is driven by the uneven distribution of heat (solar radiation) at the equator. It is controlled by the density of the water. Residence time of a water molecule in the ocean is estimated at around 3,200 years

Coastal upwelling

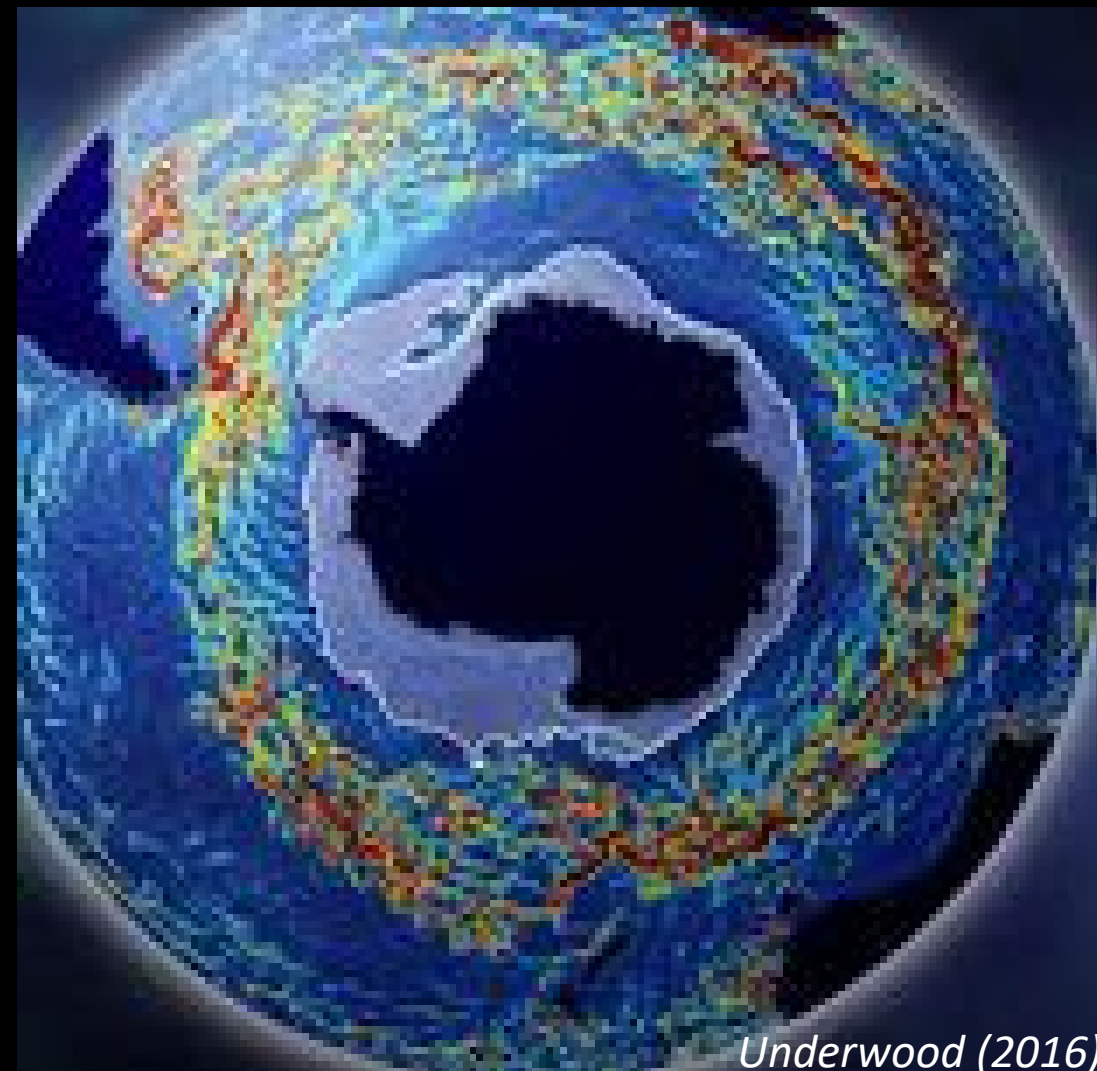


Kämpf & Chapman (2016)



Credit: CNX OpenStax
https://en.wikipedia.org/wiki/Marine_biogeochemical_cycles#/media/File:Figure_44_02_04.png

Wind-driven open ocean upwelling



Underwood (2016)

In 1934, Alfred C. Redfield reported a suite of dissolved nitrate, phosphate and oxygen measurements from various depths in the Atlantic, Pacific and Indian oceans. These data showed a remarkable consistency.

With nitrate and phosphate occurring in a ratio 16:1, and expanded to include a ratio of carbon to phosphate of 106:1, this Redfield ratio, consistent for dissolved and particulate matter, has come to define our understanding of ocean biogeochemical cycling.

Deviations from this ratio are now providing insights into the nutrient dynamics of oceans modern and ancient.



Alfred Redfield

C:N:P
106:16:1

A diagnostic for understanding how these elements move through the ocean!

Carbon cycling from nutrients

Carbon is harder to measure than nitrate or phosphate. Because of Redfield, scientists can estimate from nutrient observations: biological carbon uptake, respiration & organic matter export.

Identify nutrient limitation

Phytoplankton typically require nutrients in a 16:1 nitrogen-to-phosphorus ratio.

$N:P < 16 \rightarrow$ nitrogen be limiting

$N:P > 16 \rightarrow$ phosphorus may be limiting.

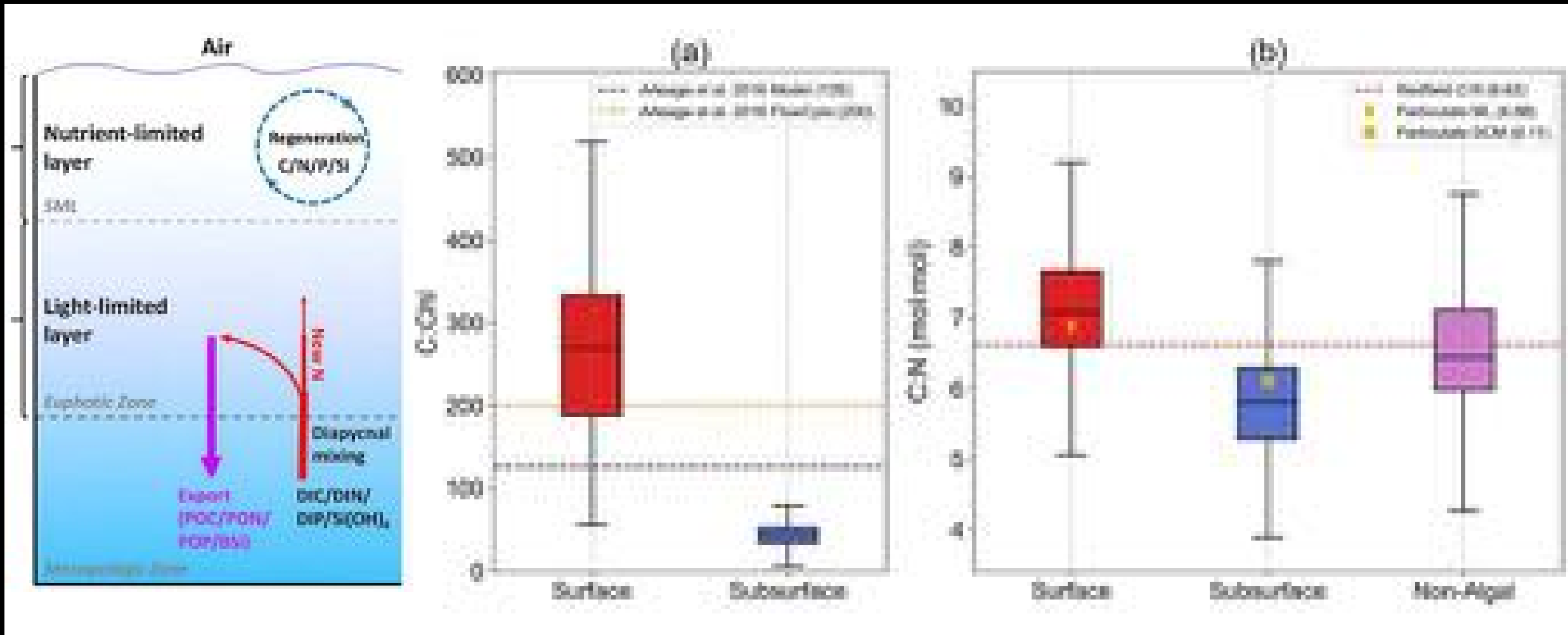
C:N:P
106:16:1

Ecosystem state

Fast-growing phytoplankton lower C:P ratios.

Nutrient-stressed cells may accumulate excess carbon, increasing C:N and C:P.

Cyanobacteria that fix atmospheric nitrogen can alter local N:P ratios.
Iron limitation in parts of the ocean changes nutrient utilization.



Dugdale (1967)

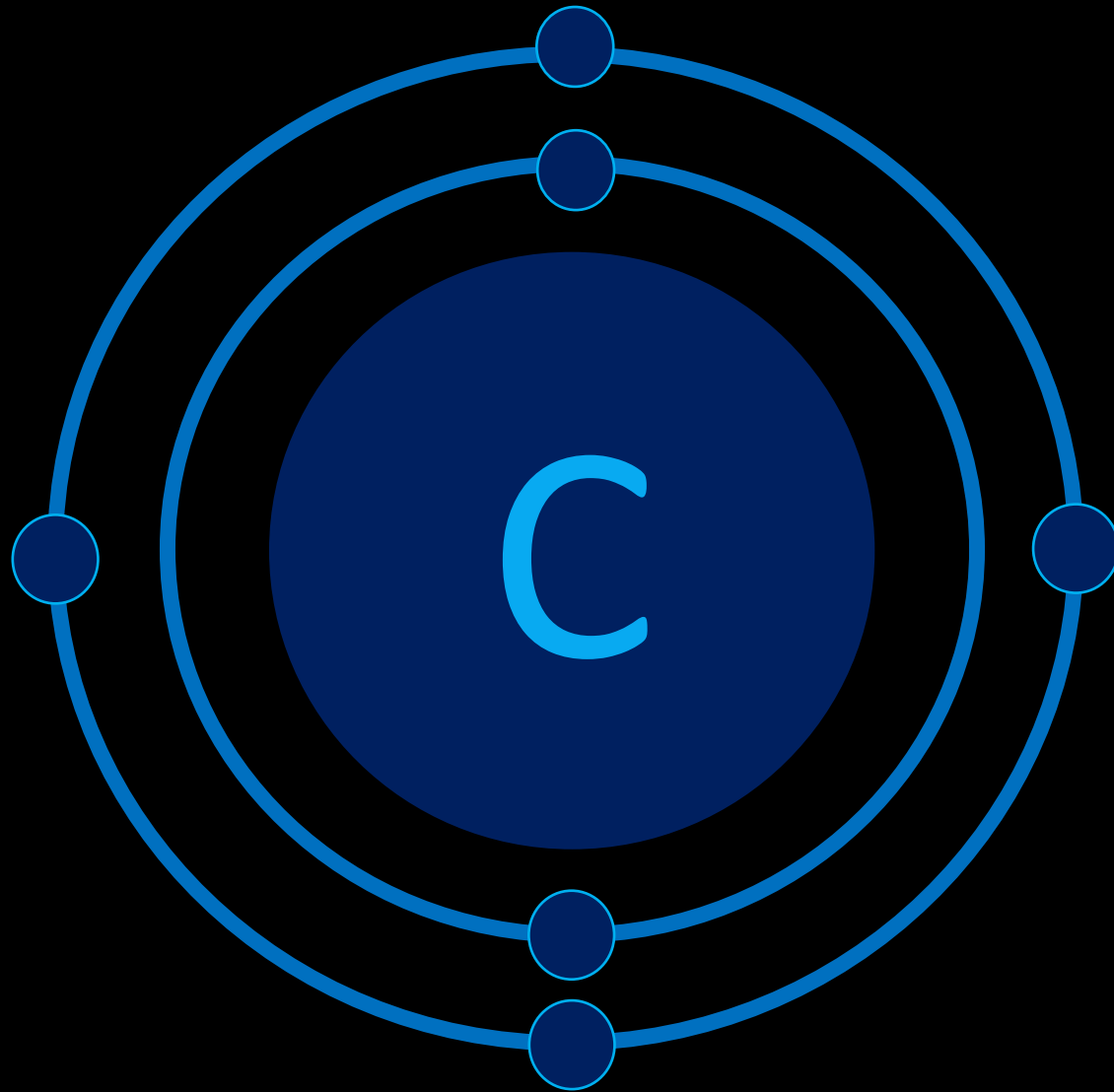
<https://doi.org/10.4319/lo.1967.12.4.0685>

Dai et al. (2023)

<https://doi.org/10.1029/2022RG000800>

Viljoen et al. (Under Review)

preprint <https://doi.org/10.5194/essd-2025-659>

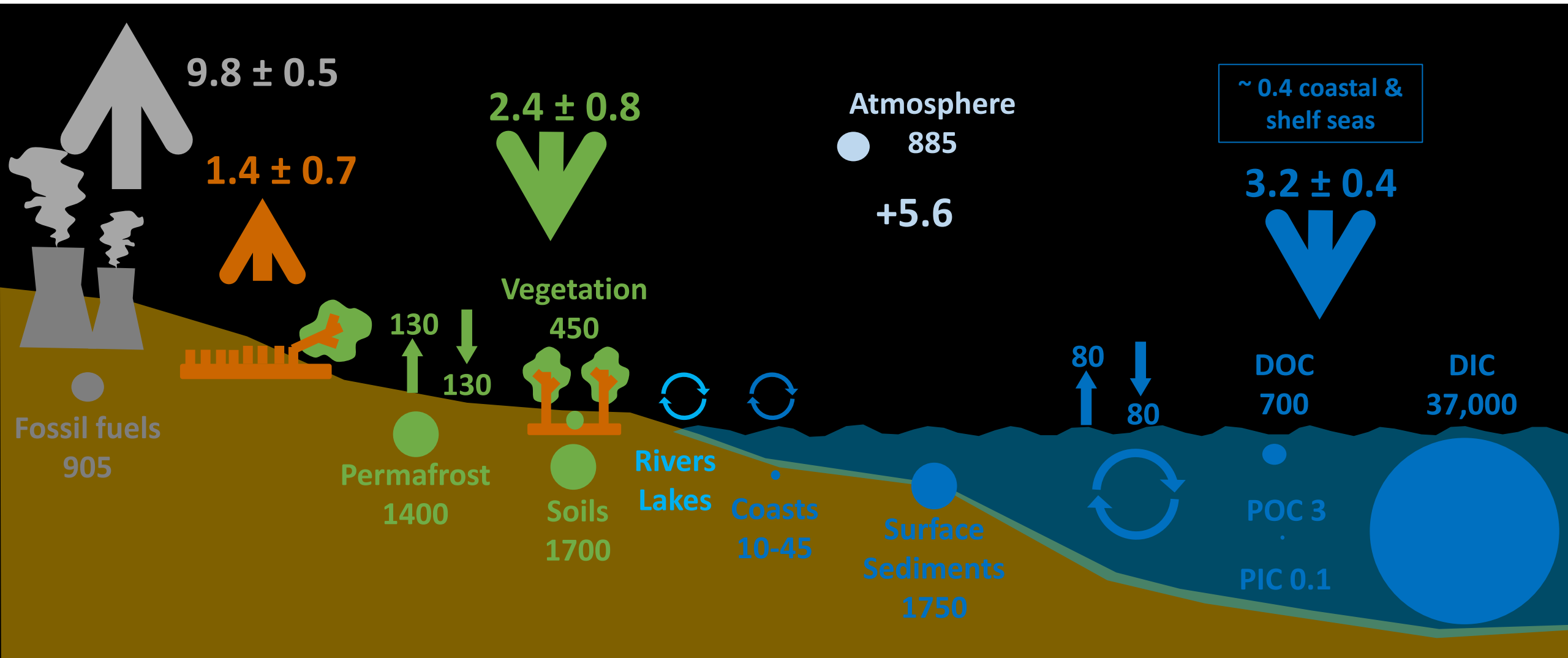


The most versatile
element on Earth

Basis of all life

Fuels human
civilization

Earth's carbon
cycle controls climate

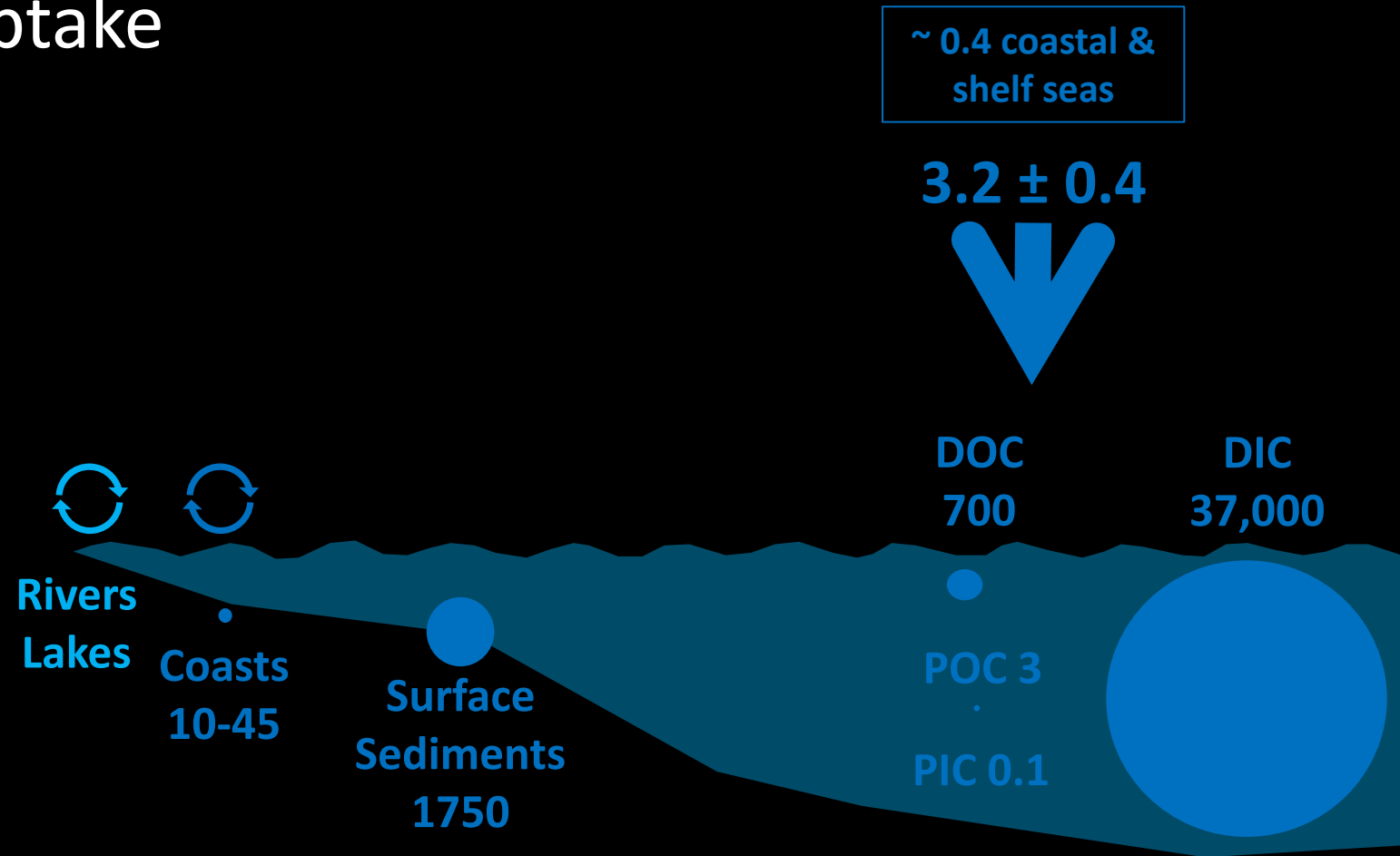


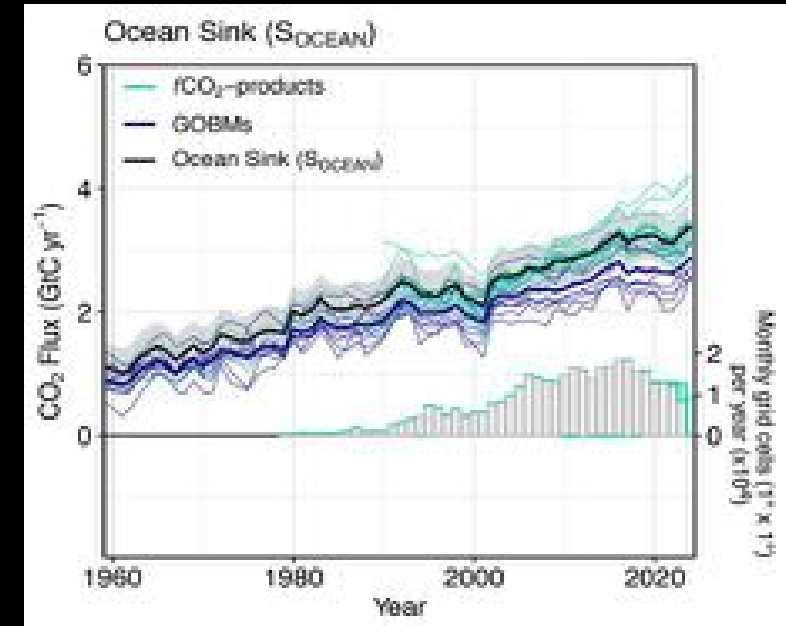
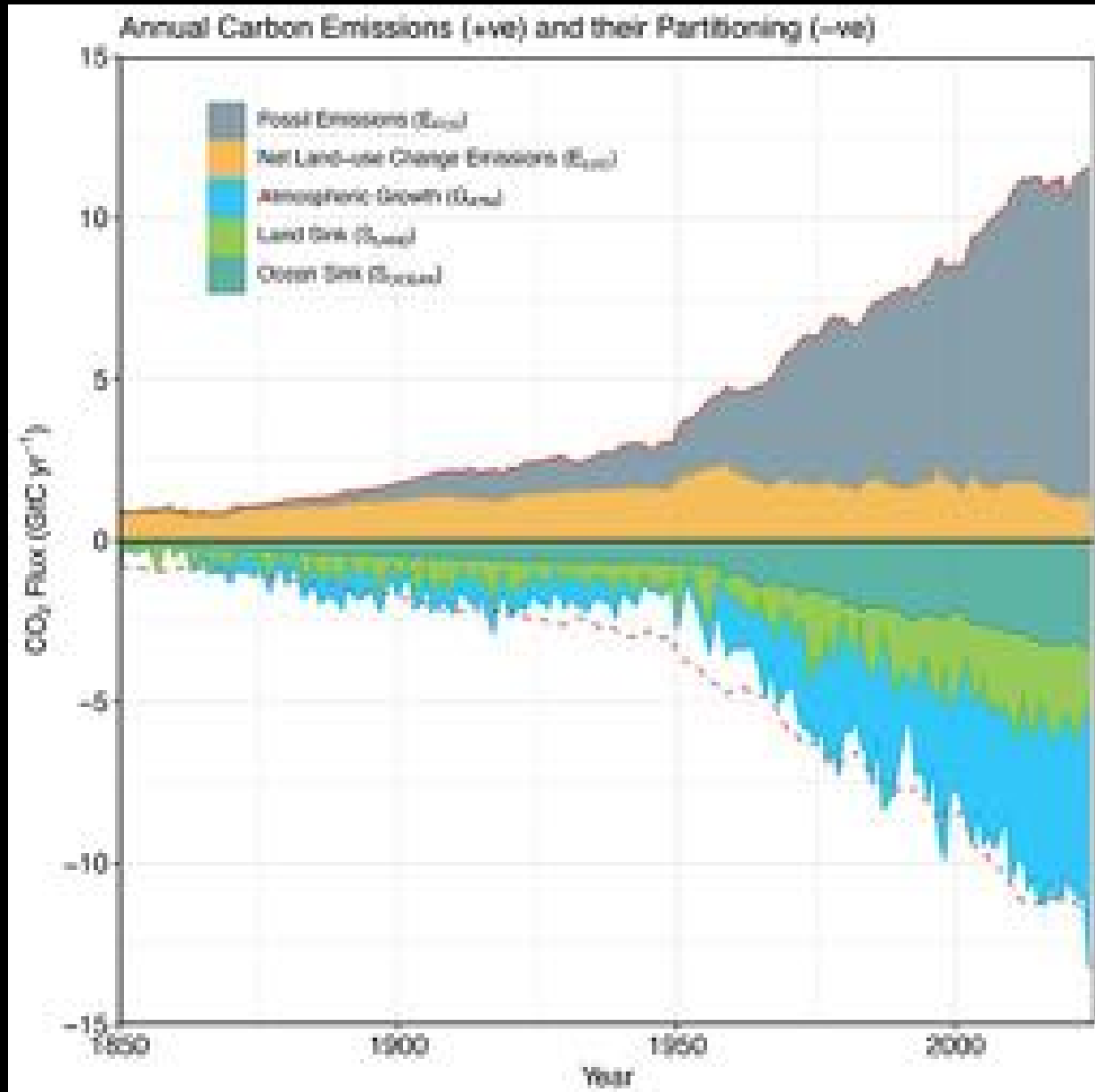
30% anthropogenic CO₂ uptake

Largest reservoir of active carbon of Earth

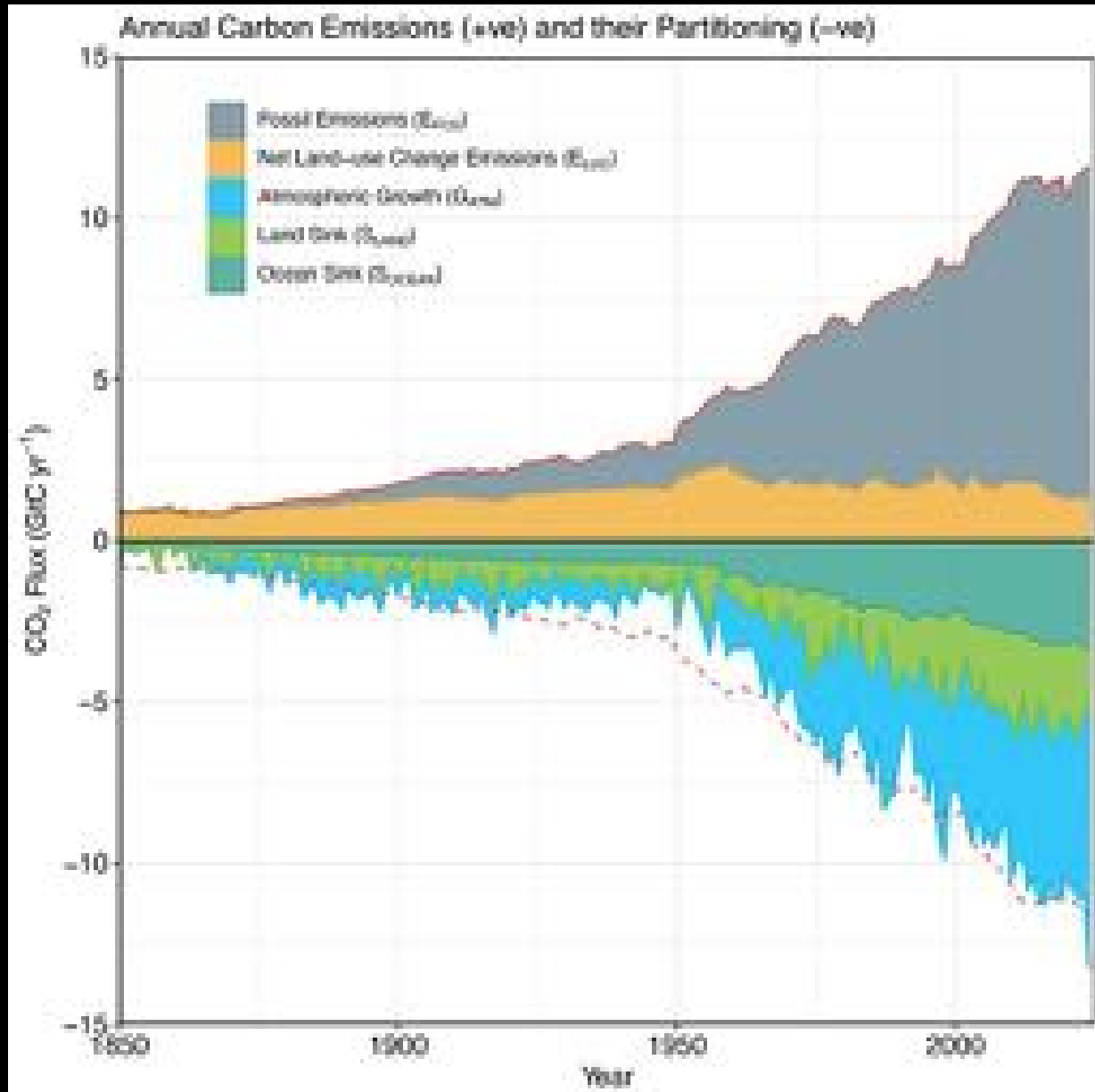
50% of planetary photosynthesis

One of the two main observational constraints on budget calculations

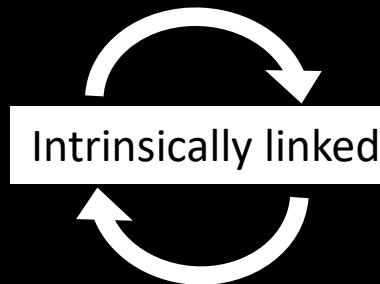




Over the period (2014-2024) the ocean CO_2 sink appears to have intensified. This emphasises the importance of the oceans in the global carbon cycle. Yet, large uncertainties remain in how the oceans will respond to increased CO_2 emissions in the future.



The solubility pump
delivers cold, dense,
DIC-rich waters to
depth mostly at high
latitudes



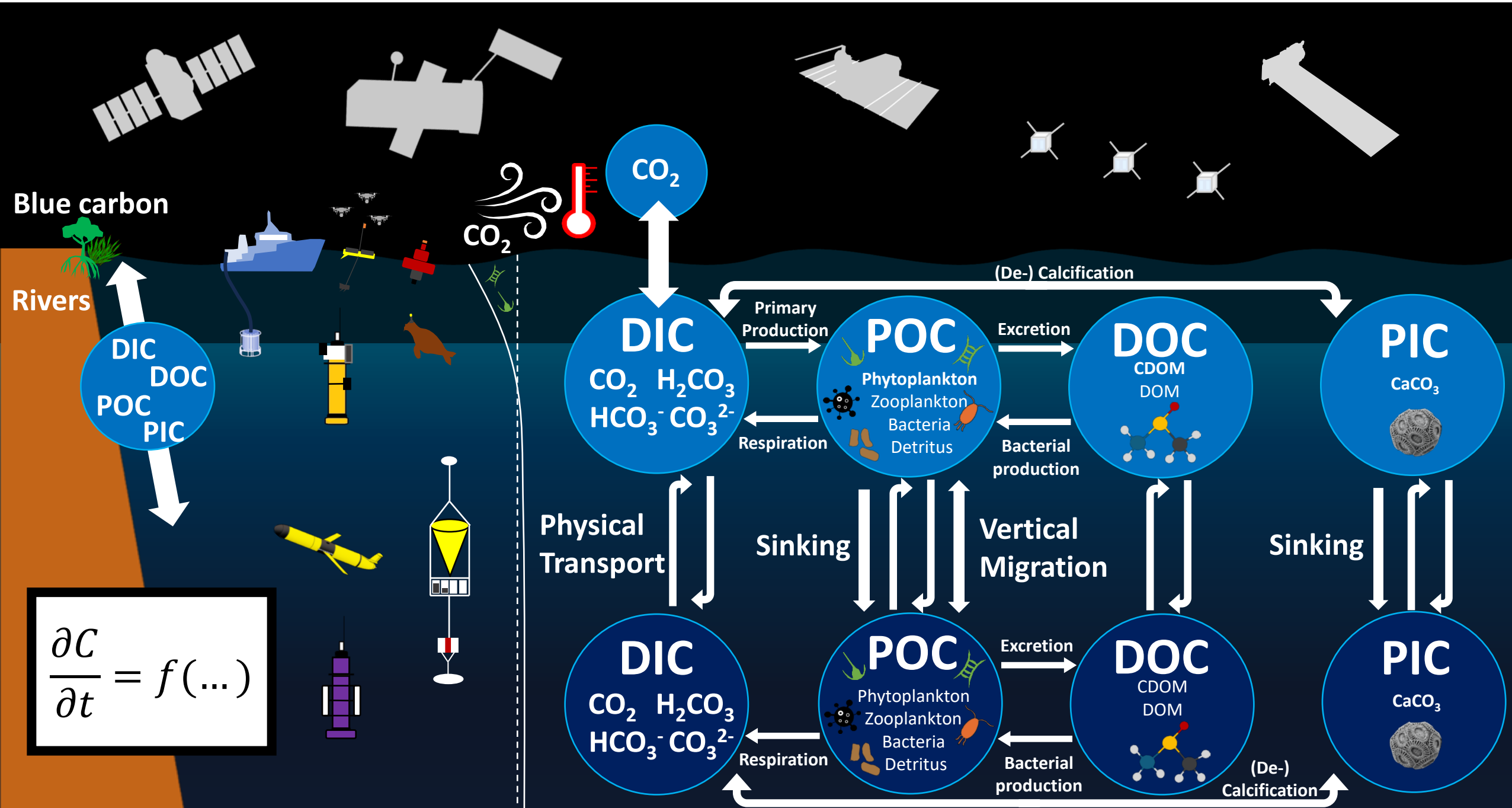
The biological pump
globally exports
organic carbon from
surface waters

SHORTER TERM

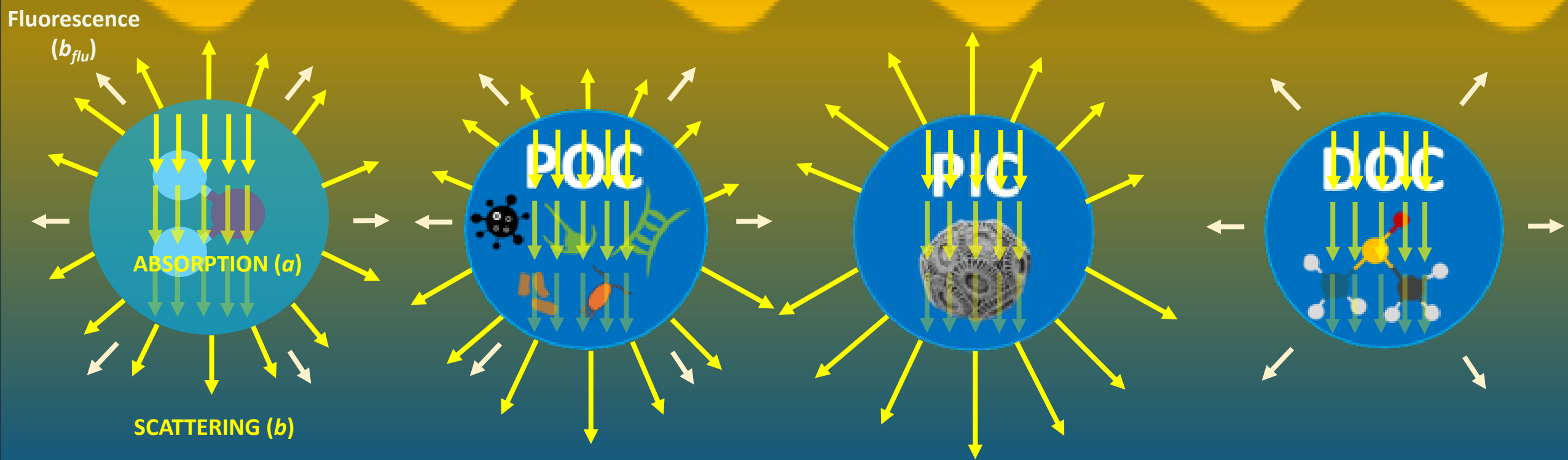
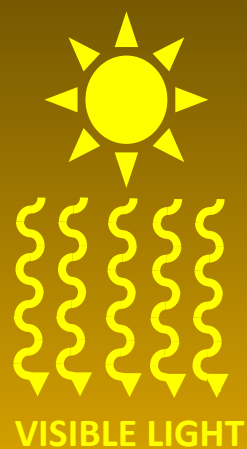
Thought to be
responsible for
uptake of
majority of
anthropogenic
 CO_2

LONGER TERM

Thought to be
responsible
fluctuations in
atmospheric CO_2
over long
timescales



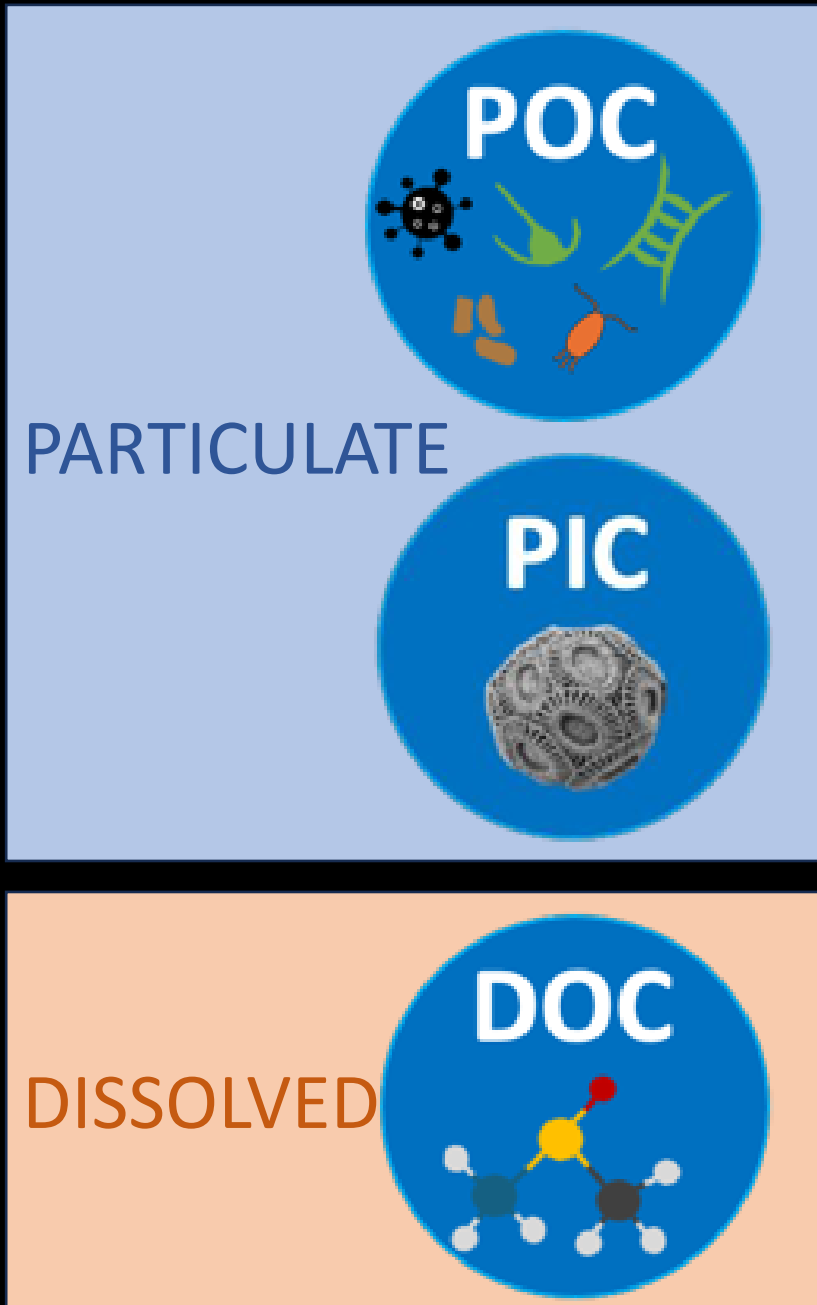
$$R_{rs}(\lambda) = \frac{L_w(\lambda)}{E_d(\lambda)} \approx r \frac{b_b(\lambda)}{a(\lambda)}$$



4. The Ocean Carbon Cycle



Credit <https://www.seabird.com/> <https://www.sequoiasci.com/> http://otronix.com/en/index.php/product/etc/?board_name=ect08&order_by=fn_pid&order_type=desc&vid=10 <https://rbr-global.com/rbrtridente-a-closer-look/>

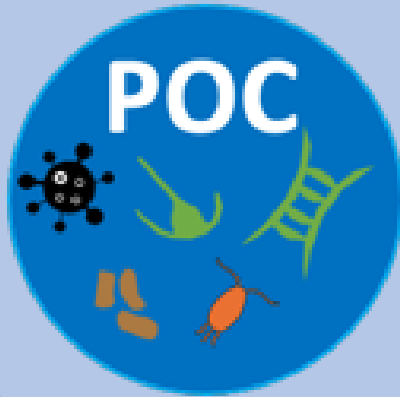


Some Broad Challenges

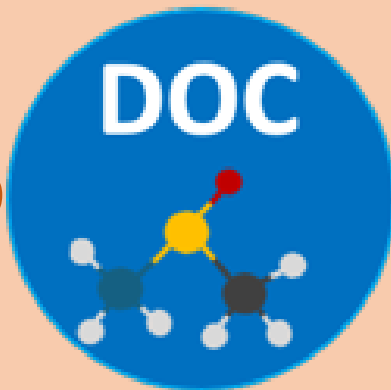
- Challenges to measuring pools of carbon analytically (e.g. see IOCCG Protocols <https://ioccg.org/what-we-do/ioccg-publications/ioccg-protocols/>, measuring phytoplankton carbon (C_{phyto}) very difficult)
- Partitioning of optical signals into contributions of POC and PIC.
- Partitioning of optical signals related to POC into its components (phytoplankton (C_{phyto}) and non-algal organic particles, particle sizes and types)
- Relationship between DOC and CDOM highly variable and often not well correlated (CDOM and DOC are largely decoupled in the open ocean)

Some Opportunities

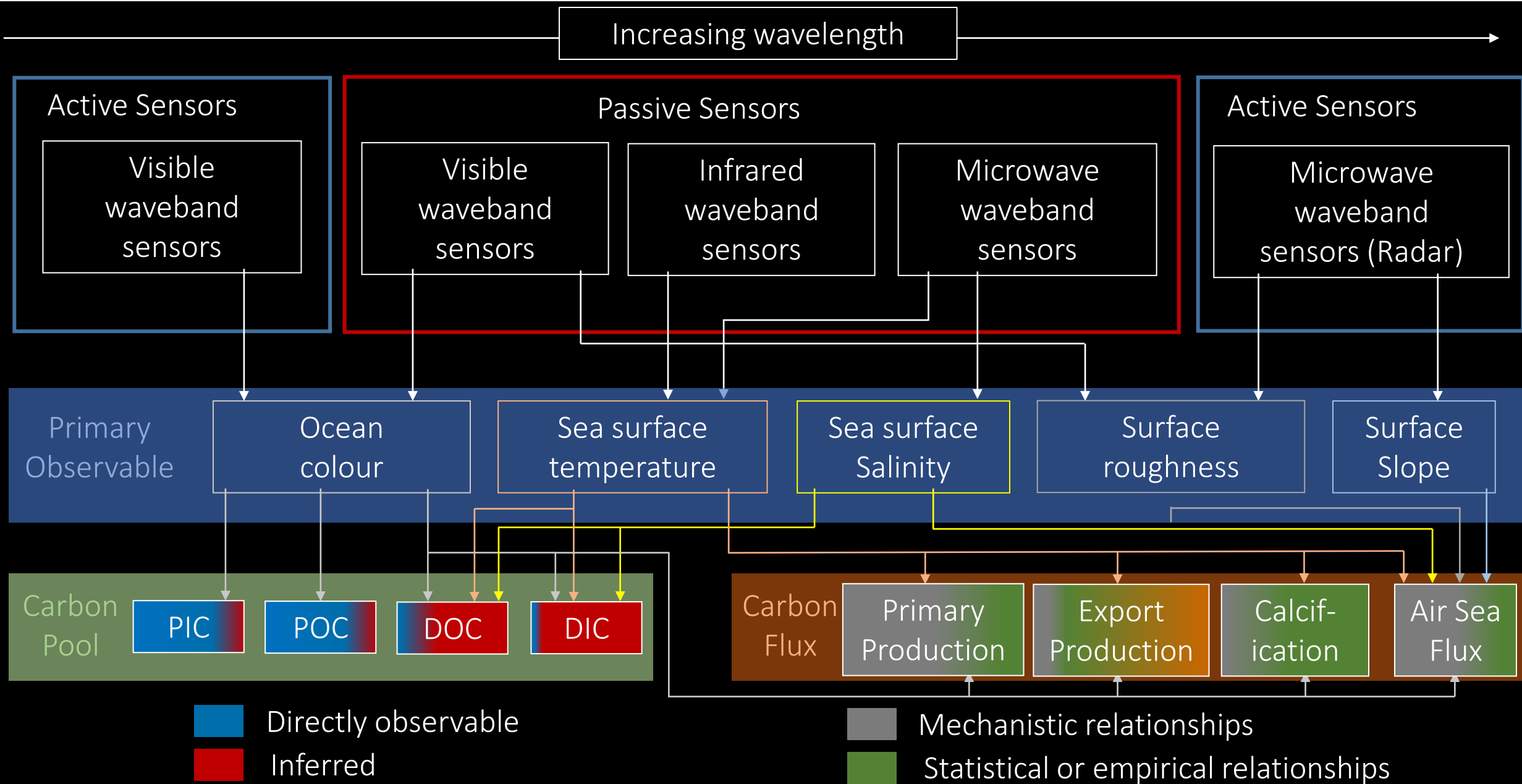
PARTICULATE

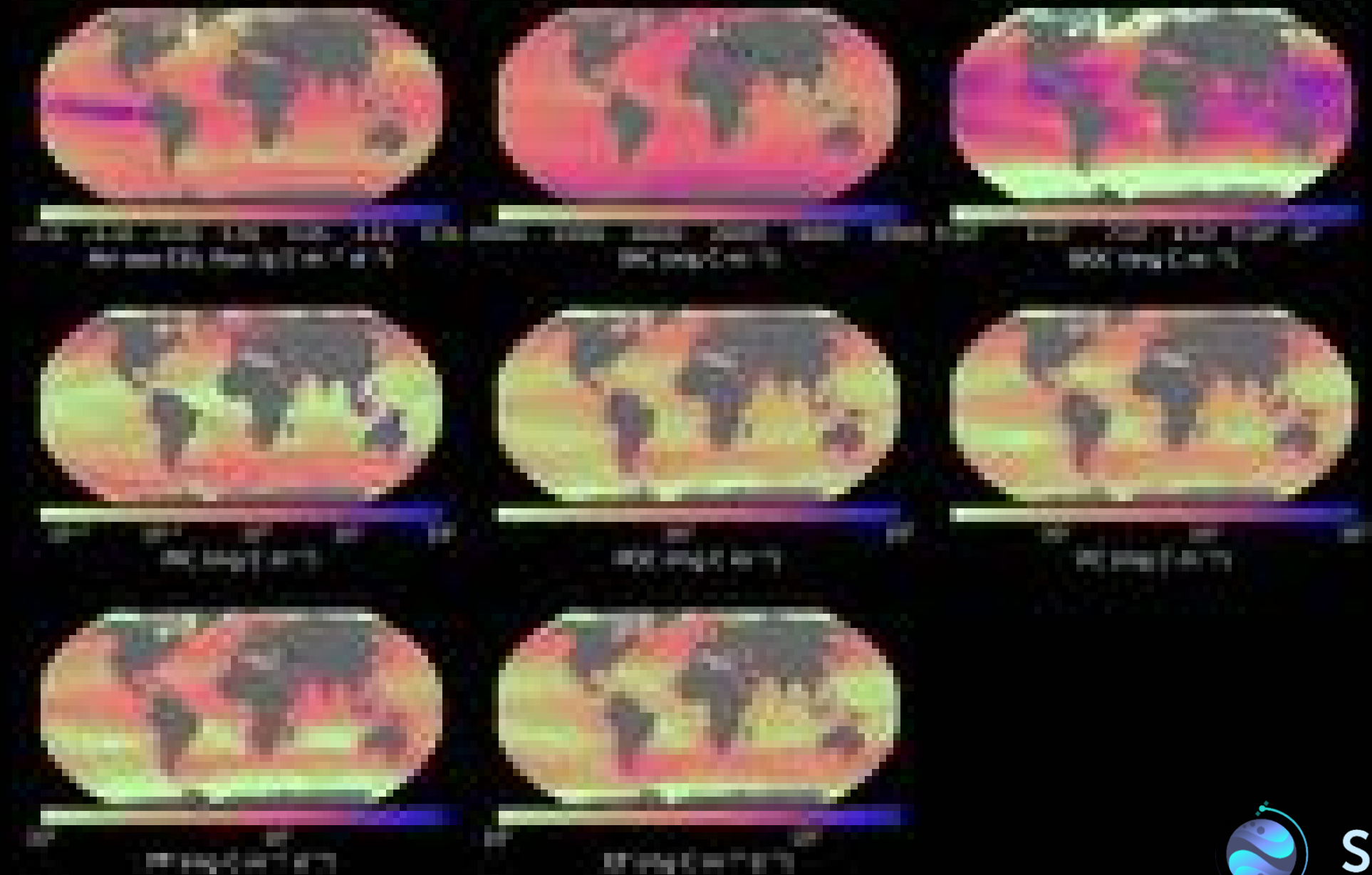


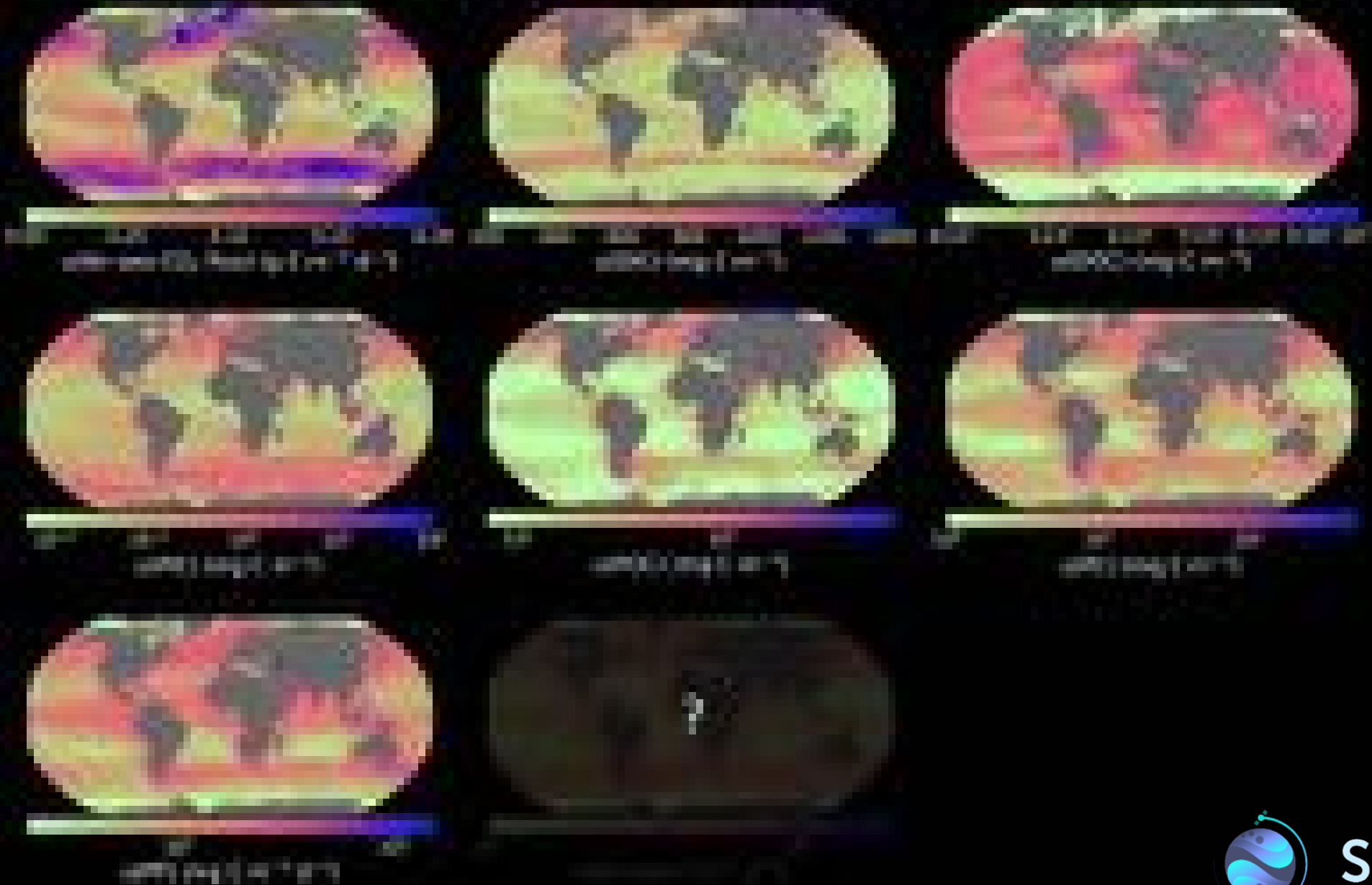
DISSOLVED

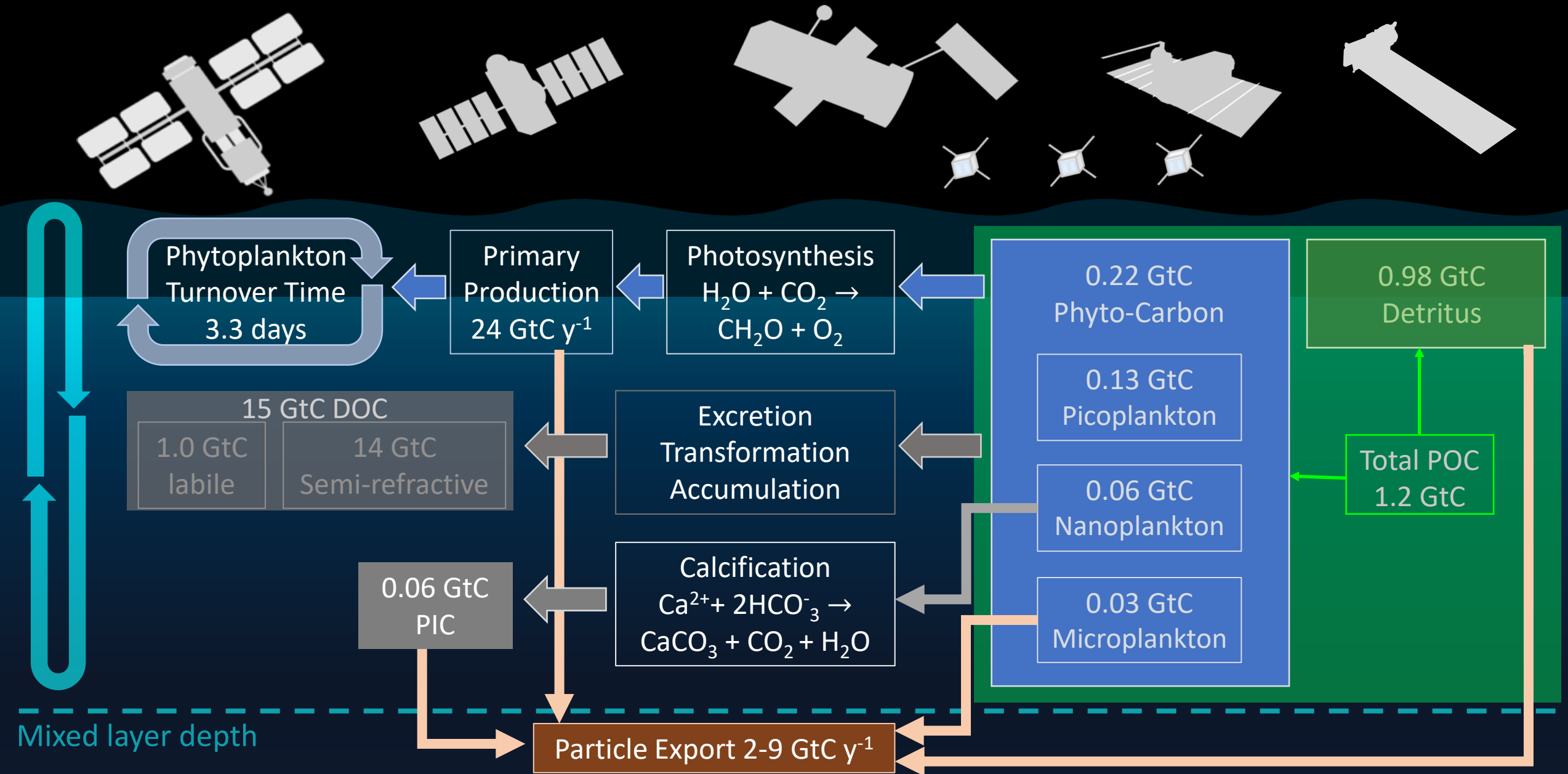


- Advance and standardise analytical methods for improved measurement of total POC, PIC and DO, and techniques to separate living and non-living POC. (<https://ioccg.org/what-we-do/ioccg-publications/ioccg-protocols/>)
- Development of optical sensors (e.g. light-scattering polarisation, hyperspectral backscattering)
- Expanding network of autonomous platforms
- Advancements in mechanistic and statistical approaches (e.g. machine learning) for relating optical signals to pools of carbon





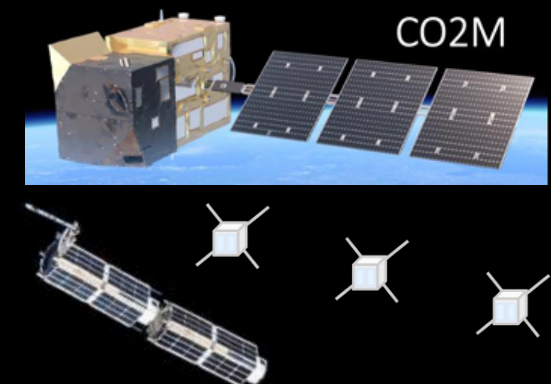
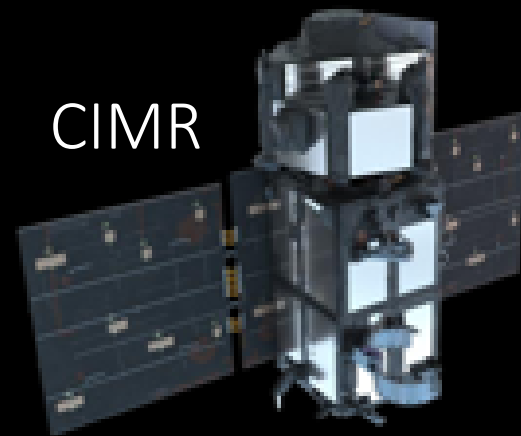
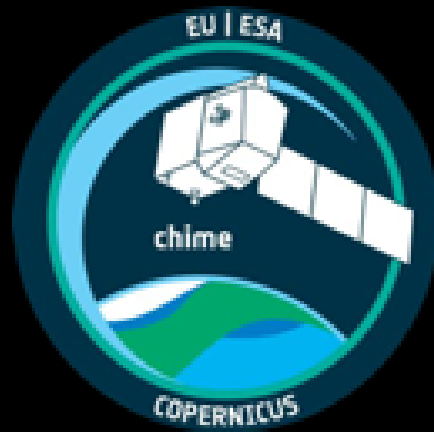
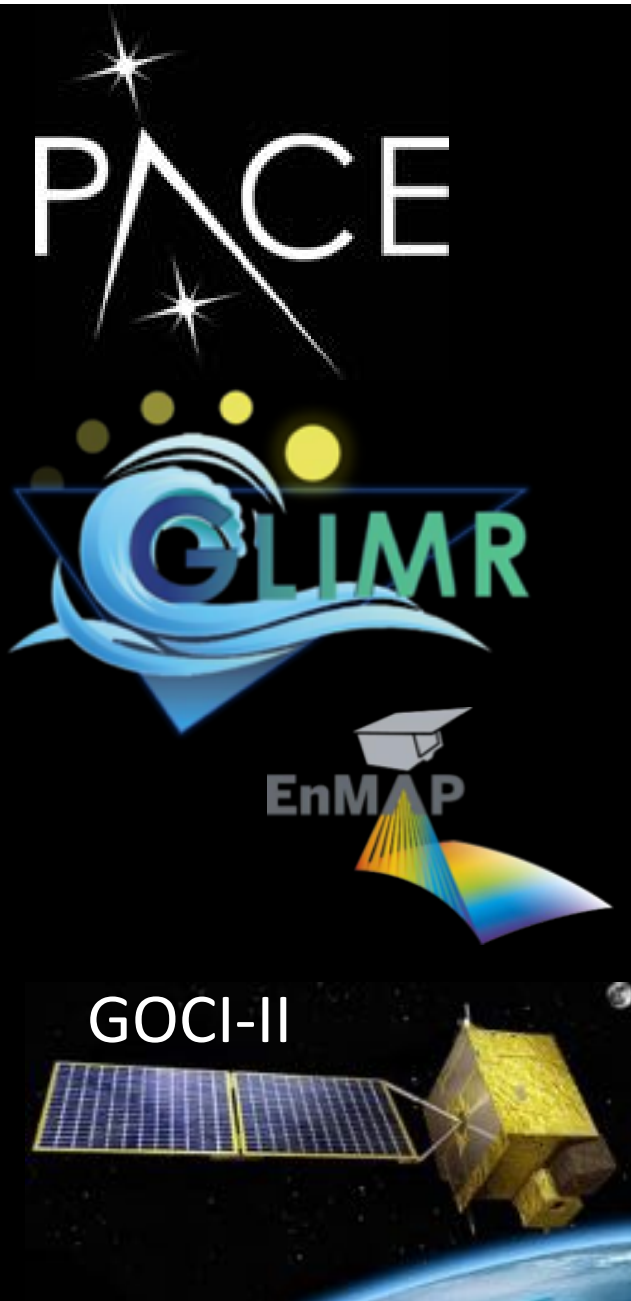




Harness recent and emerging satellites



Behrenfeld et al. (2023) <https://doi.org/10.3390/rs15235567>

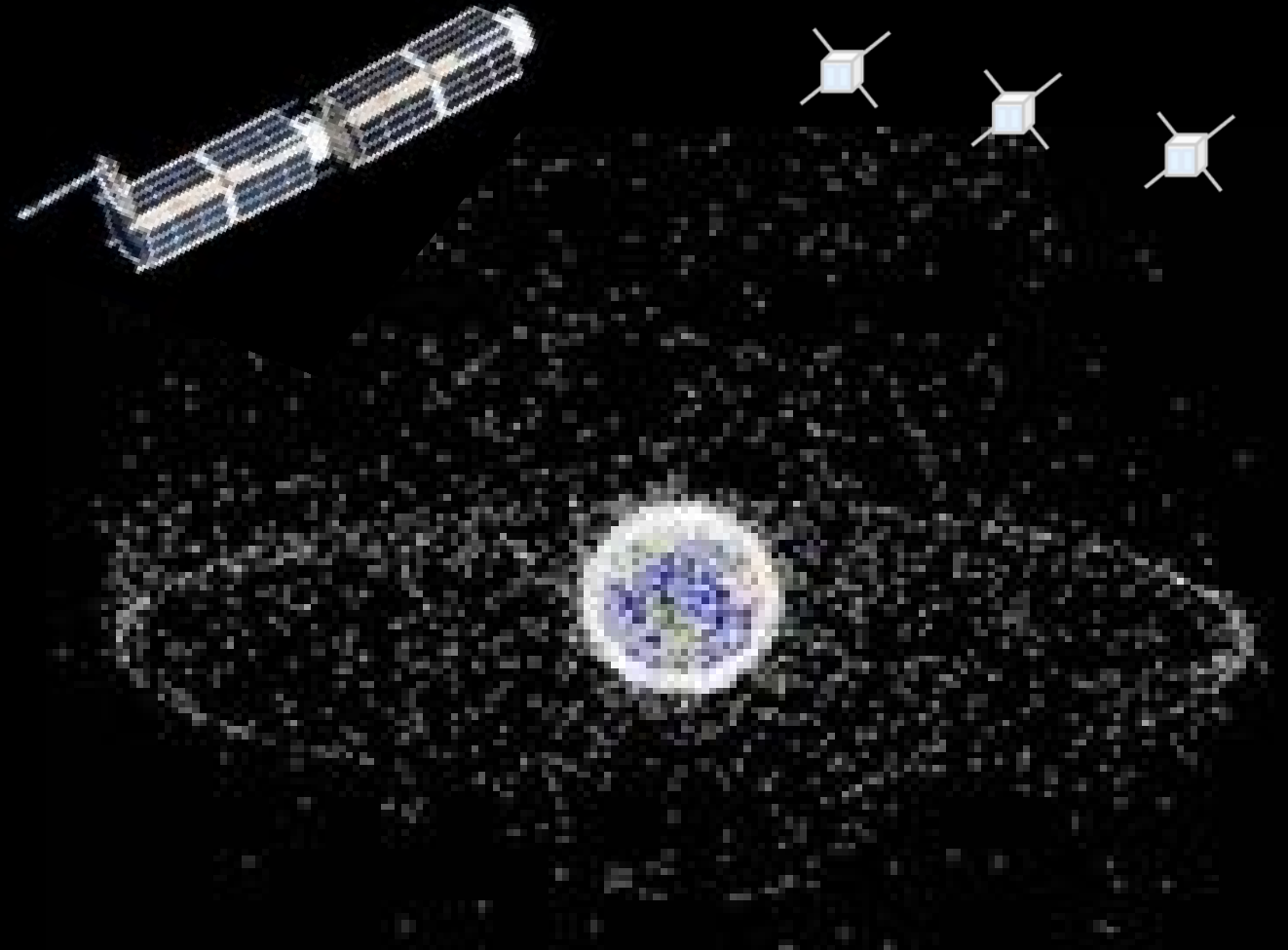


Small and commercial satellites



Bresnahan et al. (2024)

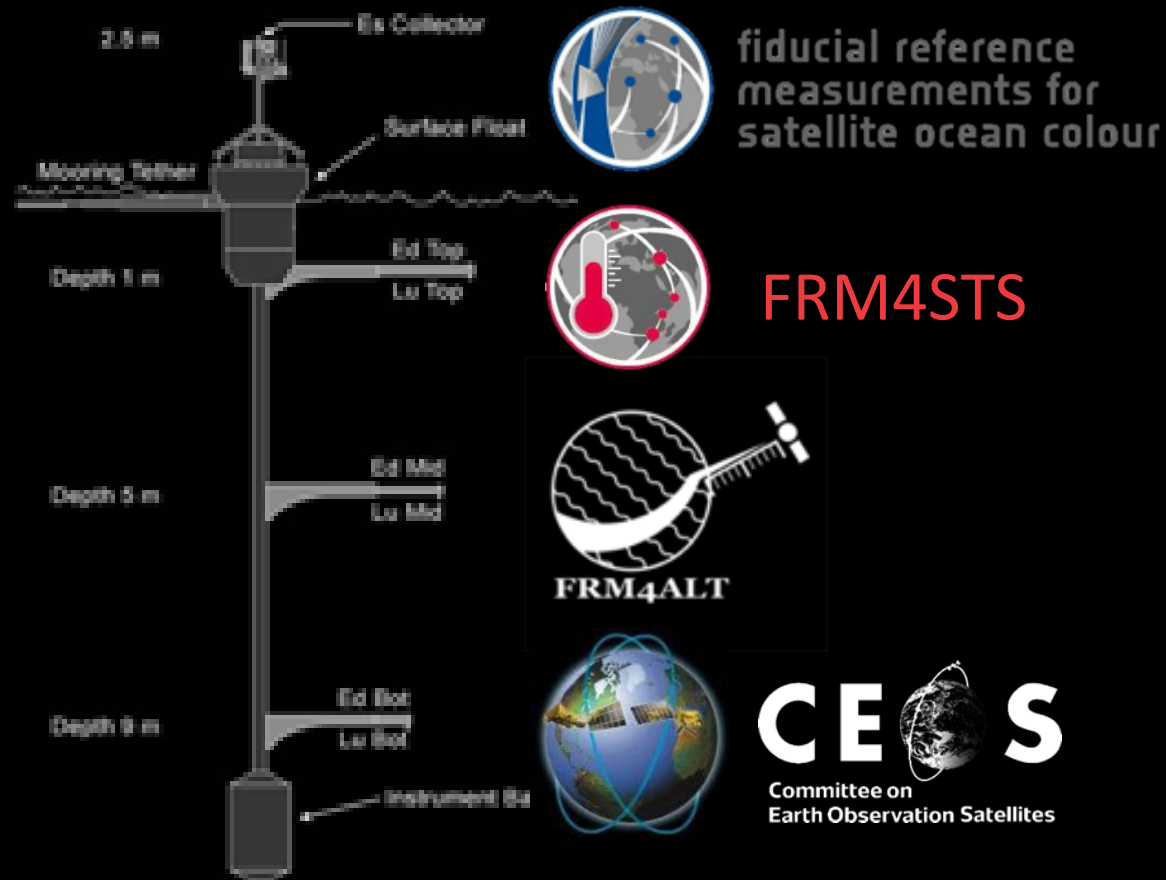
<https://doi.org/10.1038/s41597-024-04076-4>



Martellato et al. (2022)

<https://doi.org/10.3390/rs14092066>

Cal/Val of primary observables



Essential role of ships

PERSPECTIVE

The essential role of large research vessels in marine ecosystem observations and ocean sustainability



Satellites will address critical science priorities for quantifying ocean carbon

Jamie D Shutler^{1*}, Rik Wanninkhof², Philip D Nightingale³, David K Woolf⁴, Dorothee CE Bakker⁵, Andy Watson⁶, Ian Ashton¹, Thomas Holding¹, Bertrand Chapron⁷, Yves Quilfen⁷, Chris Fairall⁸, Ute Schuster⁶, Masakatsu Nakajima⁹, and Craig J Donlon¹⁰

Shutler et al. (2020) <https://doi.org/10.1002/fee.2129>

Sensing the ocean biological carbon pump from space: A review of capabilities, concepts, research gaps and future developments

Robert J.W. Brewin^{a,b,*}, Shubha Sathyendranath^{b,c}, Trevor Platt^b, Heather Bouman^d, Stefano Ciavatta^{b,c}, Giorgio Dall'Olmo^{b,c}, James Dingle^b, Steve Groom^b, Bror Jönsson^b, Tihomir S. Kostadinov^e, Gemma Kulk^b, Marko Laine^f, Victor Martínez-Vicente^b, Stella Psarra^g, Dionysios E. Raitos^h, Katherine Richardsonⁱ, Marie-Hélène Rio^j, Cécile S. Rousseaux^{k,l}, Joe Salisbury^m, Jamie D. Shutler^a, Peter Walker^b

Brewin et al. (2021) <https://doi.org/10.1016/j.earscirev.2021.103604>

**ANNUAL
REVIEWS**

Annual Review of Marine Science

Quantifying the Ocean's Biological Pump and Its Carbon Cycle Impacts on Global Scales

David A. Siegel,¹ Timothy DeVries,¹ Ivona Cetinić,^{2,3} and Kelsey M. Bisson⁴



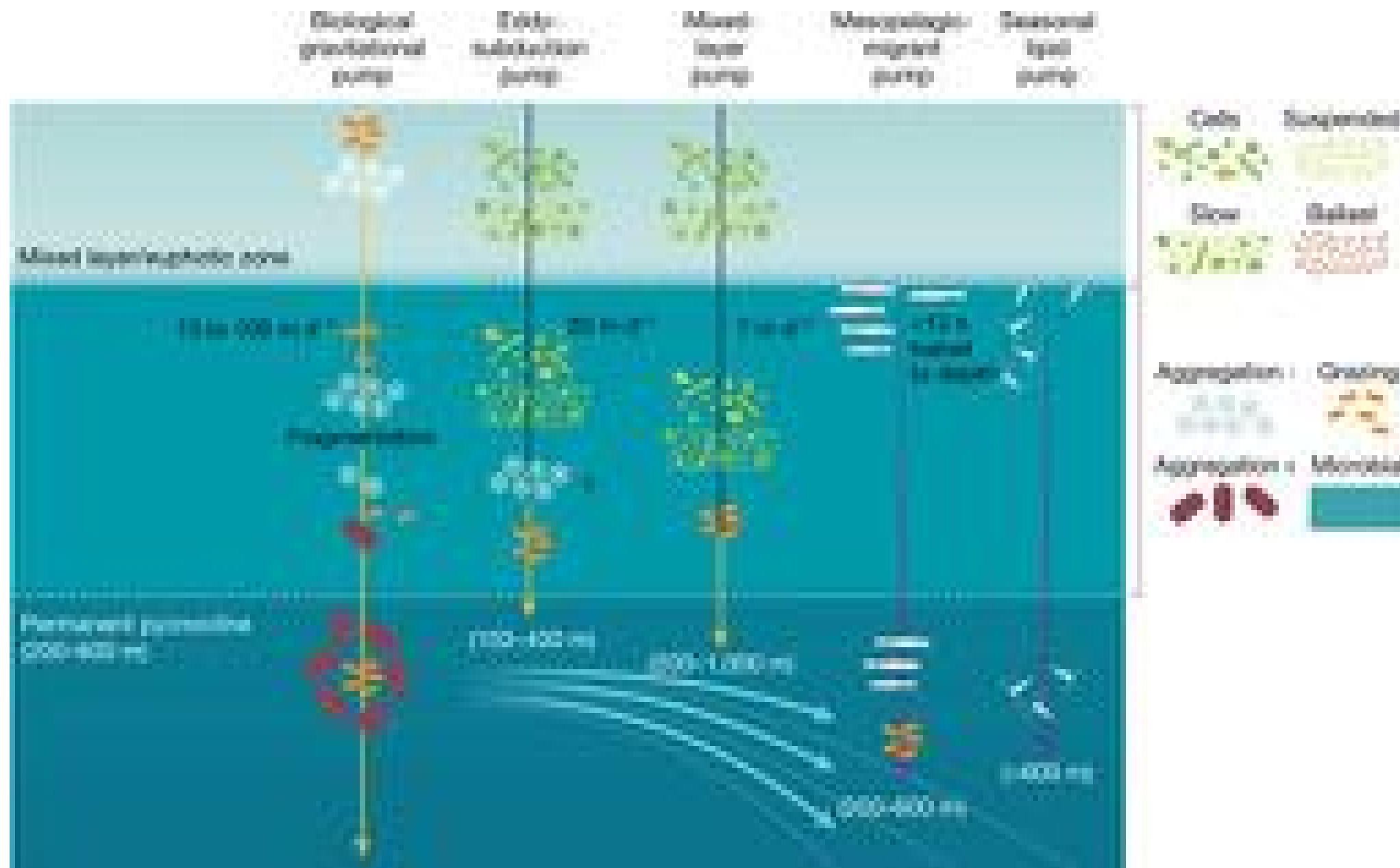
Earth-Science Reviews
Supports open access.

Special Issue
Aquatic carbon stocks and fluxes: The big picture from remote sensing
(first volume 22 April 2021)

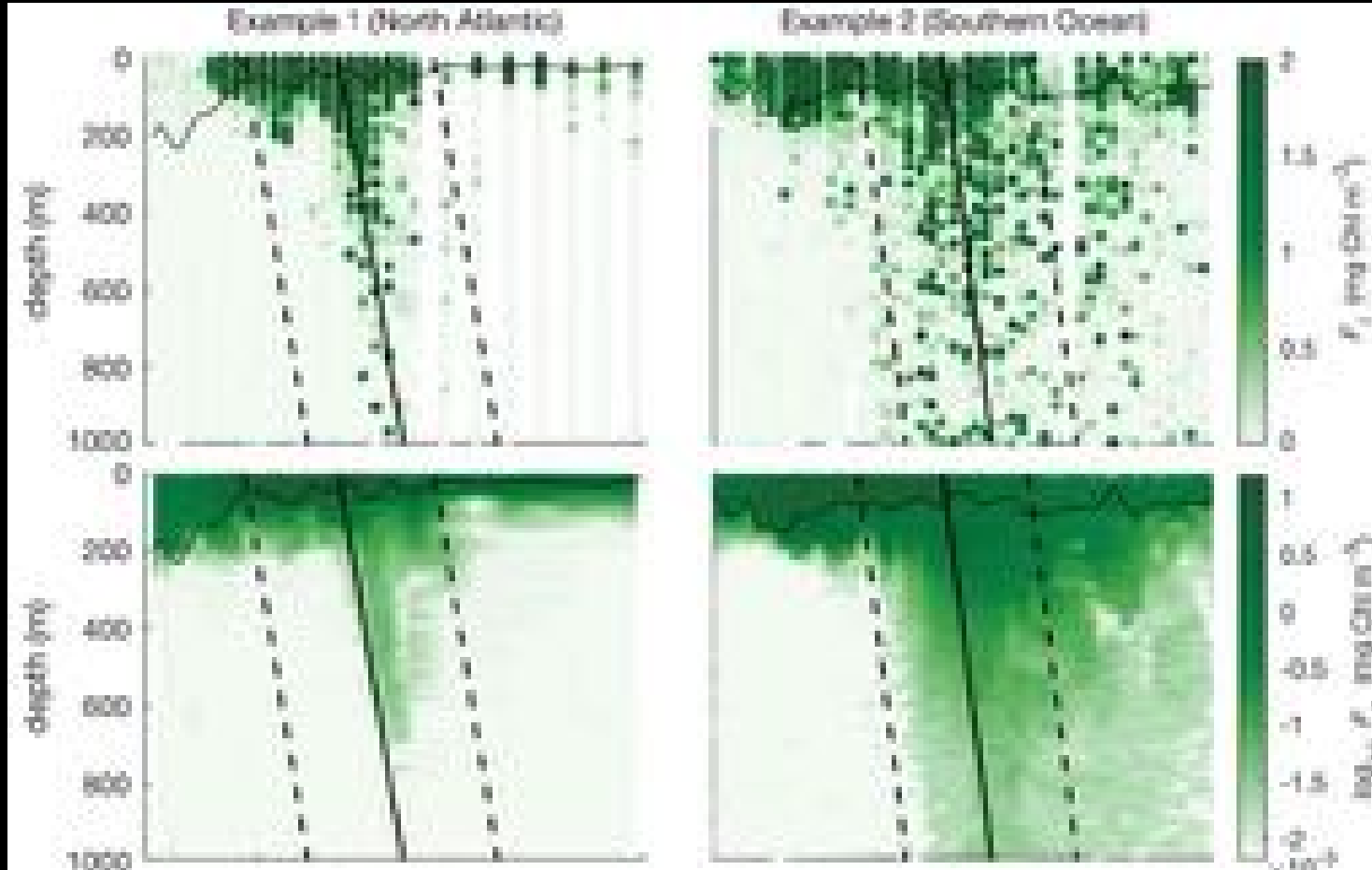
Aquatic landscapes contribute 80% of the habitable space on Earth. In these watery realms, microscopic and macroscopic organisms conduct half of the biosphere's photosynthesis, physical and chemical processes remove billions of tons of carbon dioxide from the atmosphere, and complex food webs both produce and consume climate-sensitive compounds, some of which are transported into deep reaches of the ocean while others can enter the atmosphere to influence a variety of processes. In recent decades, our understanding of carbon cycling, biogeochemistry, ecology, and interactions of global aquatic systems with the atmosphere has greatly improved through observations using airborne and satellite remote sensing systems. With these observations, we have realized a revolution in our quantitative assessment of carbon stocks and fluxes in aquatic systems. While a rich scientific literature exists detailing these advances, we lack a compiled synthesis of developments, findings, and future directions that is accessible both to researchers in aquatic sciences and the broader scientific community. The synthesis manuscripts in this collection summarize our current understanding of carbon stocks and fluxes in aquatic systems, and explore directions for future advancements, with a particular focus on globally relevant topics that rely on the use of remote sensing data, including integration with in-situ measurements and models. This

<https://www.sciencedirect.com/special-issue/10JHBMJVSNW>

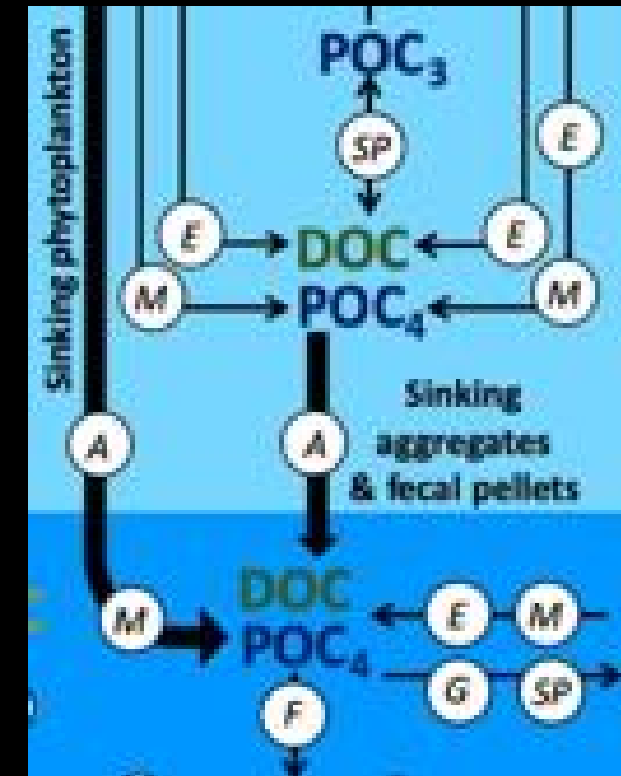
Siegel et al. (2023) <https://doi.org/10.1146/annurev-marine-040722-115226>

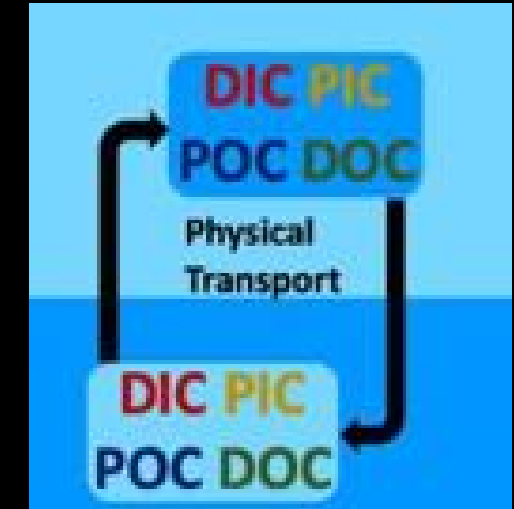
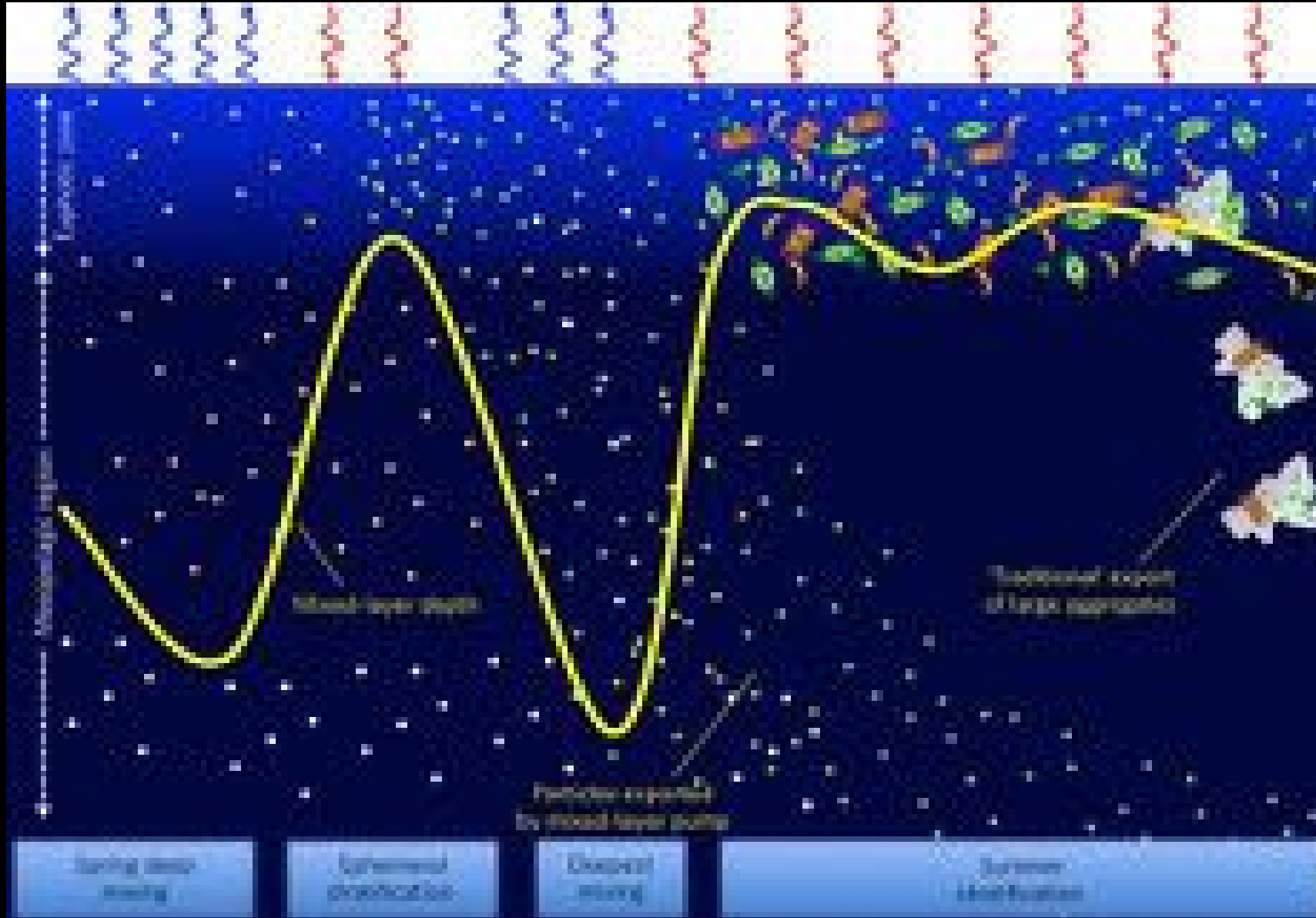


Boyd et
al. (2019)
<https://doi.org/10.1038/s41586-019-1098-2>

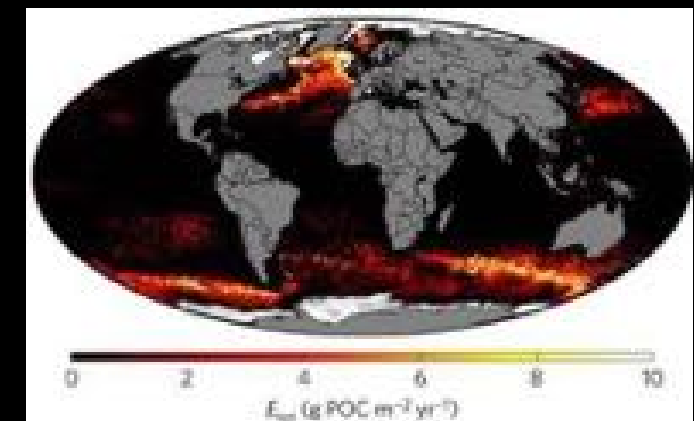


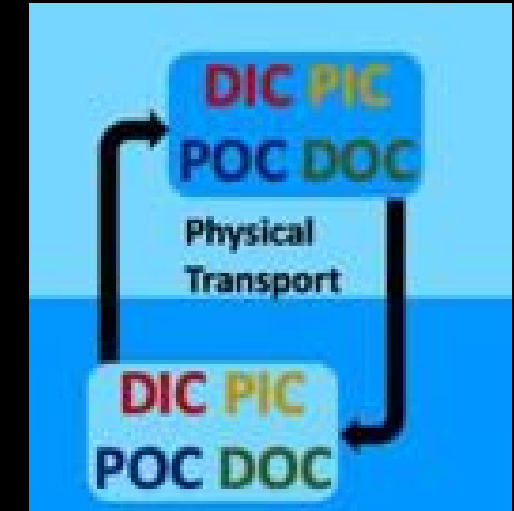
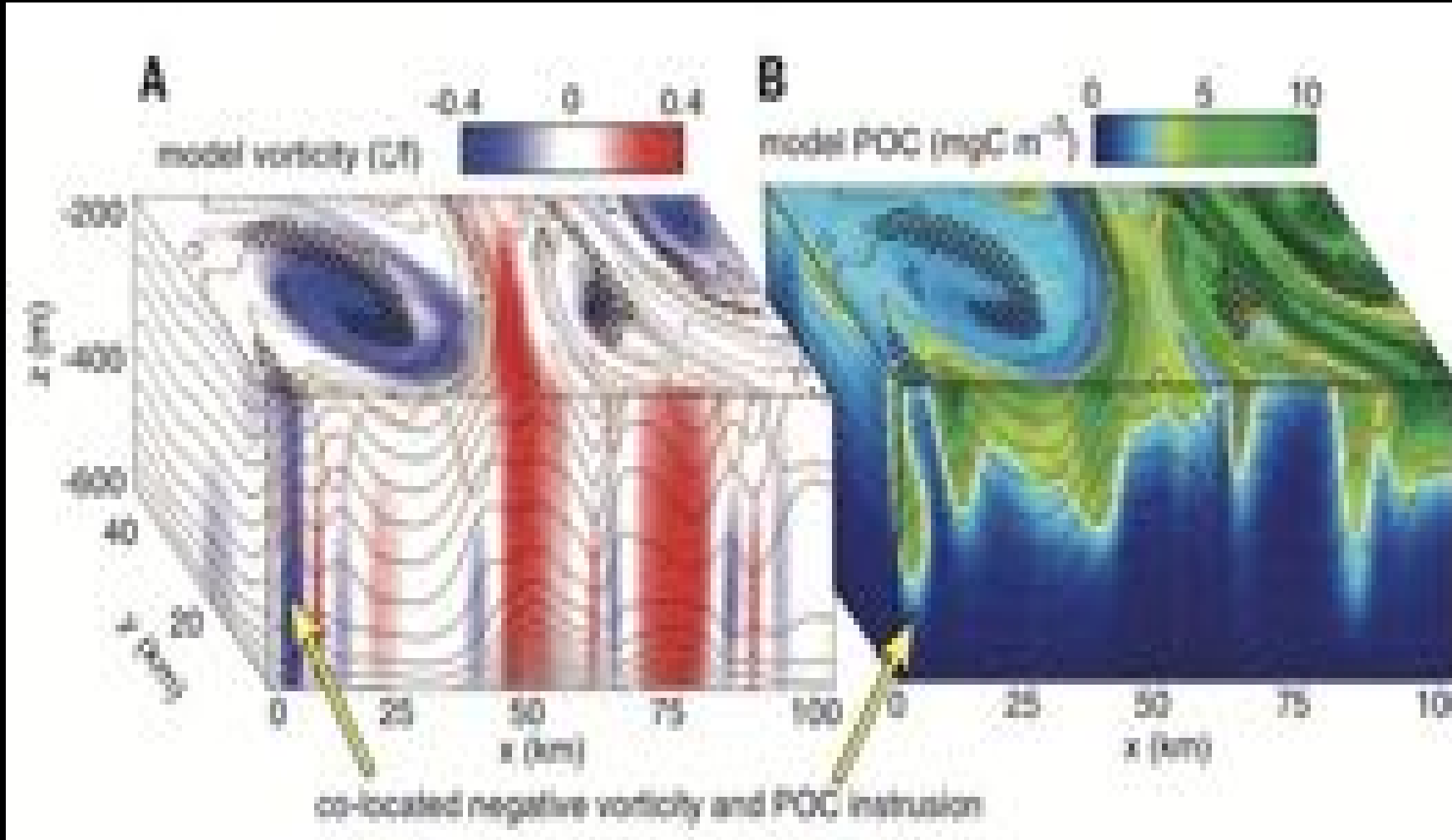
Briggs et al. (2020) <http://doi.org/10.1126/science.aay1790>





Mixed-layer pump is responsible for a global flux of $0.1 - 0.5 \text{ Pg C yr}^{-1}$



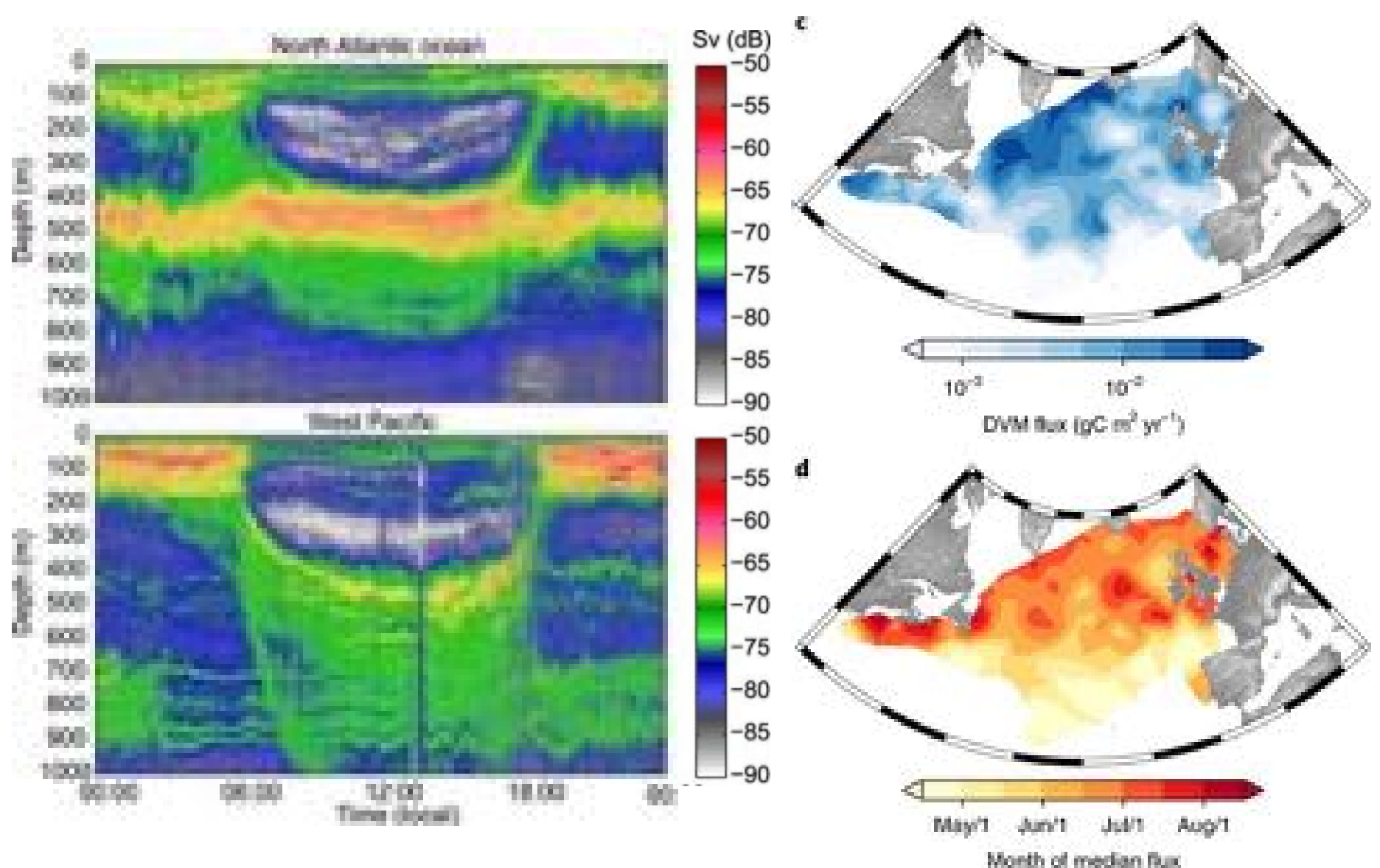


Export of POC through subduction can contribute as much as $\frac{1}{2}$ the total springtime export of POC from the highly productive subpolar oceans



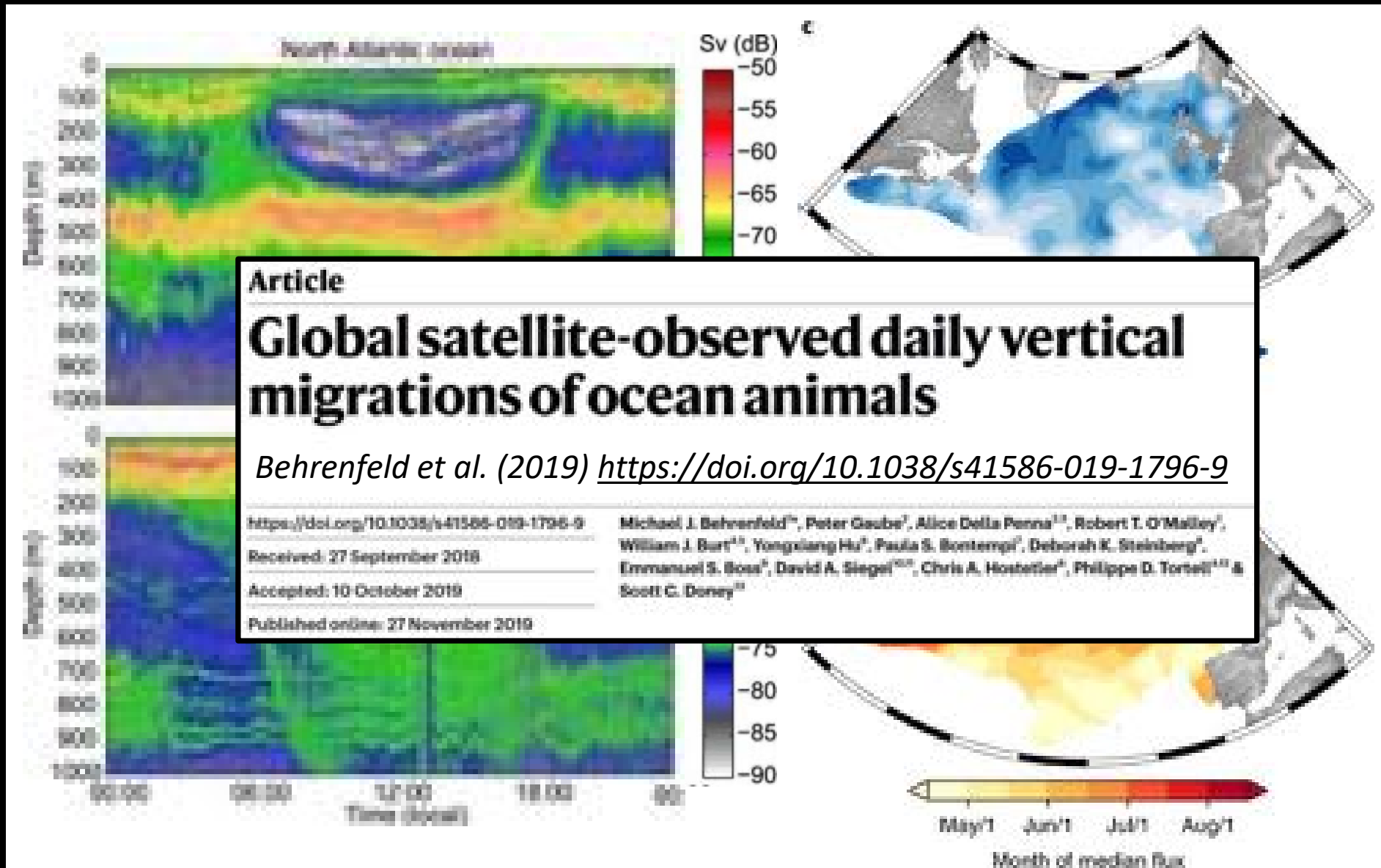
Omand et al. (2015) <http://doi.org/10.1126/science.1218740>





Mesopelagic-migrant pump injecting particles (fast sinking zooplankton faecal pellet) to greater depth

Vertical migration

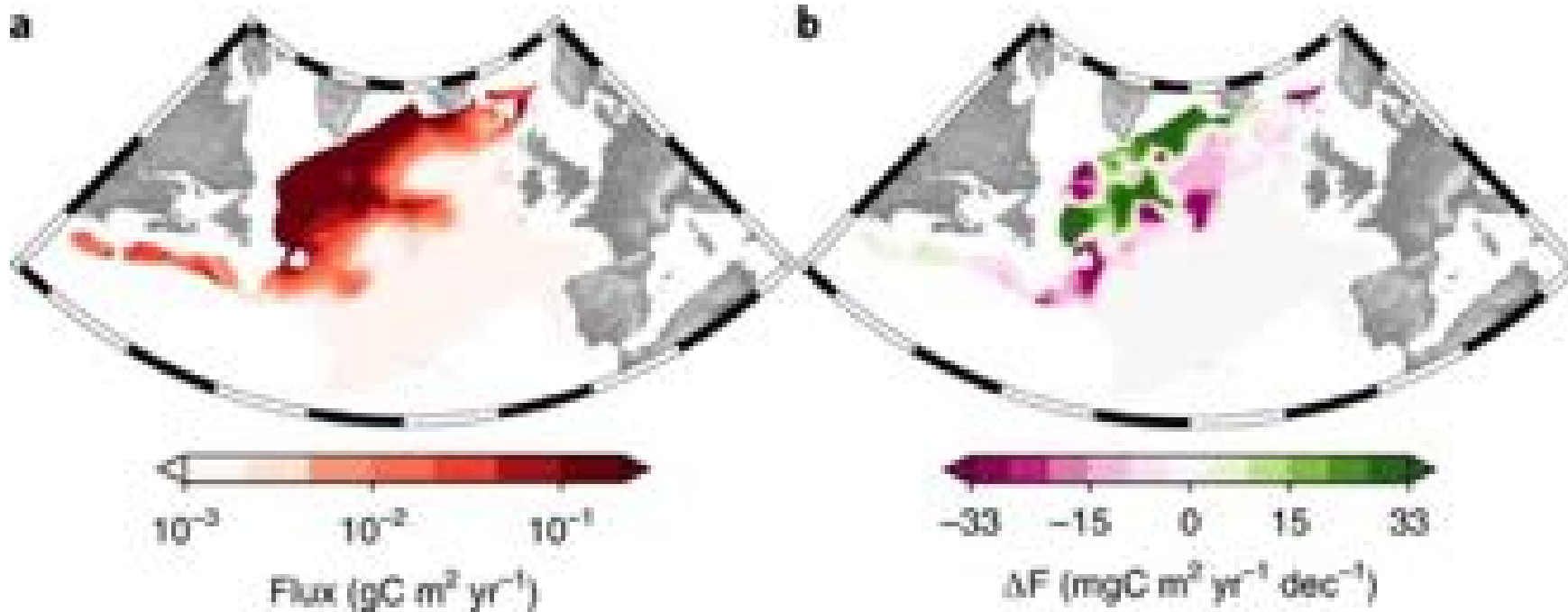


Mesopelagic-migrant pump injecting particles (fast sinking zooplankton faecal pellet) to greater depth



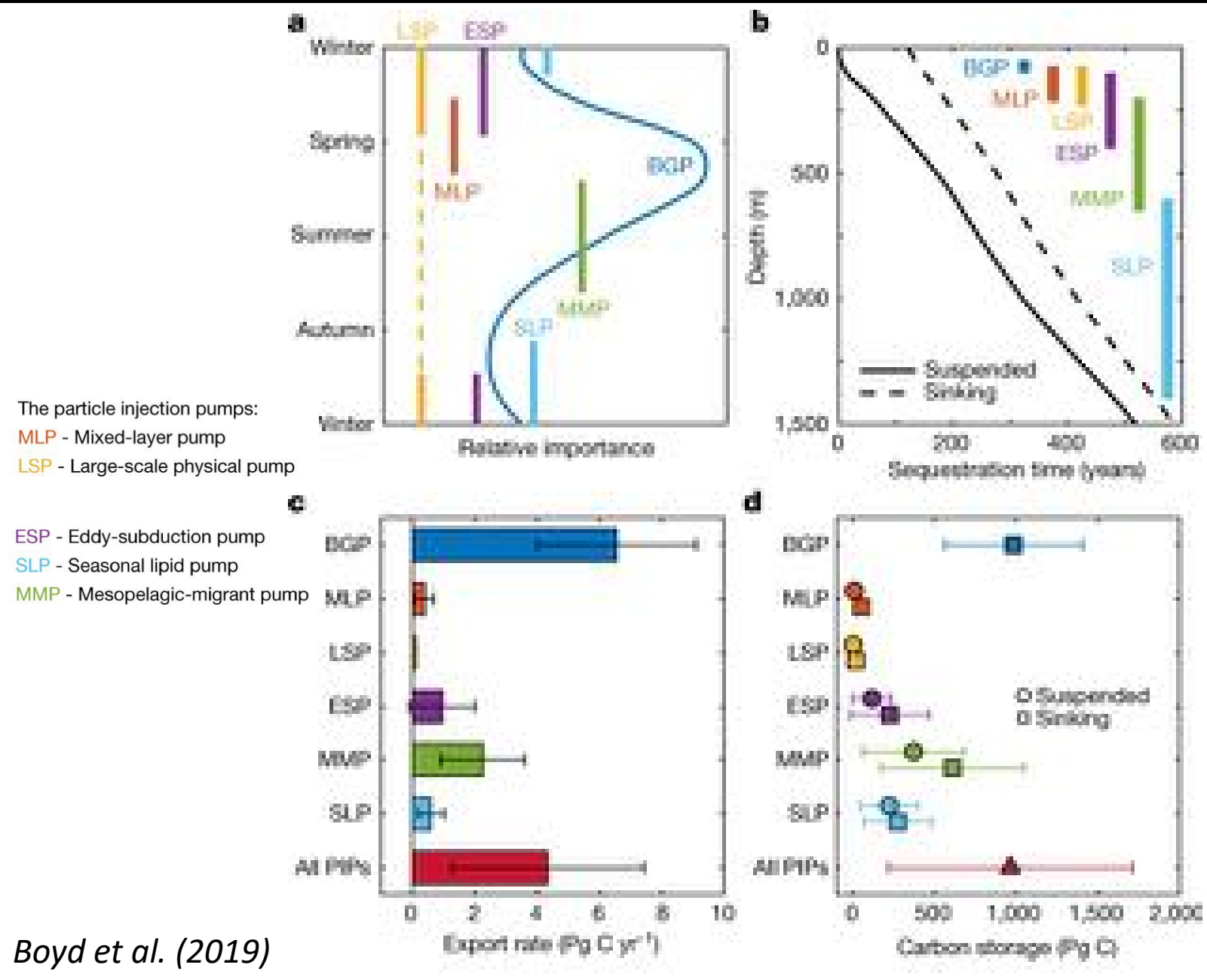
Klevjer et al. (2016) <https://doi.org/10.1038/srep19873>

Brun et al. (2019) <https://doi.org/10.1038/s41559-018-0780-3>



Hibernation of copepods at depths between 600 m and 1,400 m gives rise to a 'seasonal lipid pump', where copepods catabolize carbon-rich lipids that they accumulated in the upper layers during summer; shunting carbon to depth.

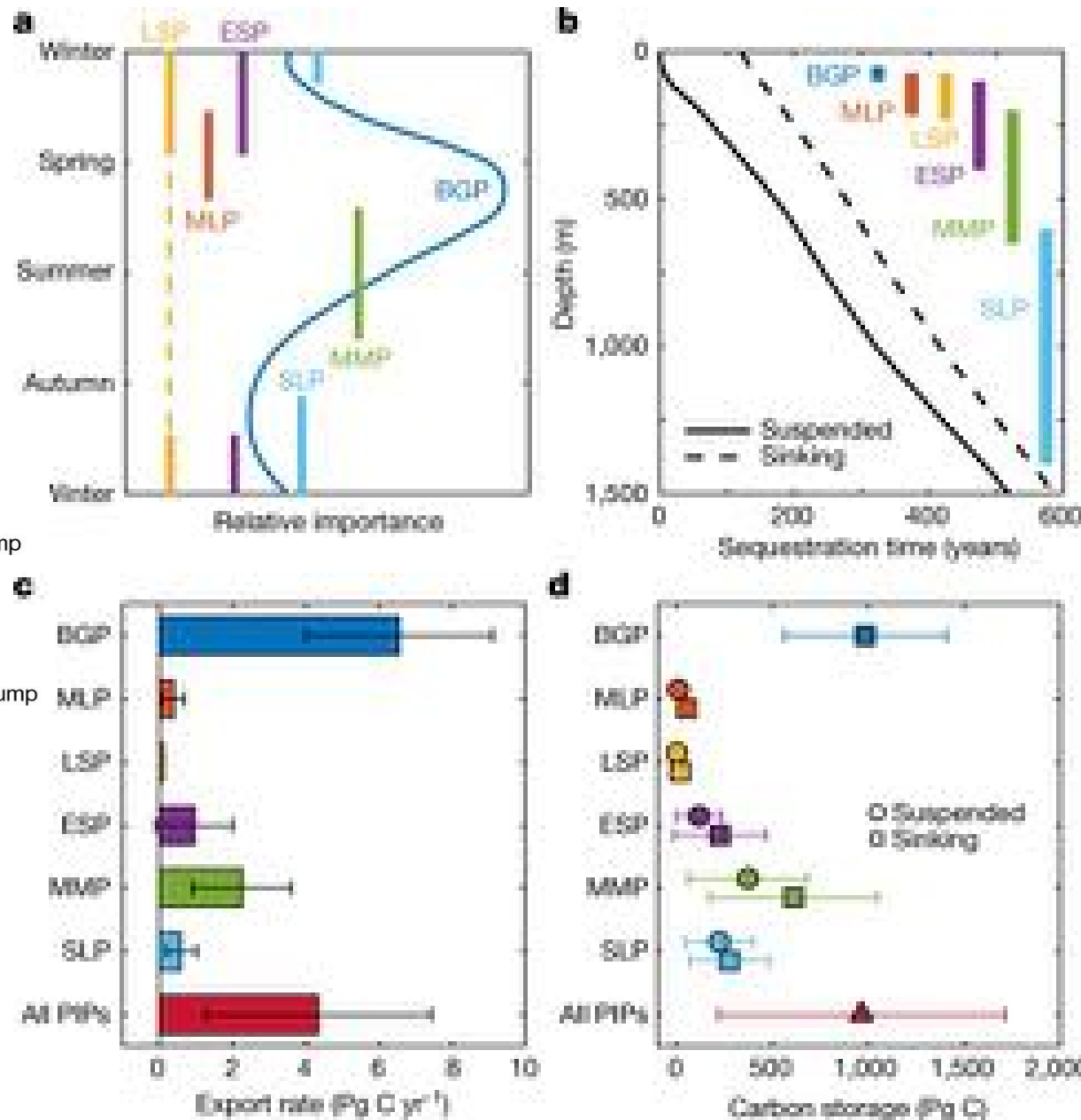




Boyd et al. (2019)

These 'particle injection pumps' probably sequester as much carbon as the gravitational pump, helping to close the carbon budget and motivating further investigation into their environmental control

Boyd et al. (2019)
<https://doi.org/10.1038/s41586-019-1098-2>



The particle injection pumps:

MLP - Mixed-layer pump

LSP - Large-scale physical pump

ESP - Eddy-subduction pump

SLP - Seasonal lipid pump

MMP - Mesopelagic-migrant pump

Boyd et al. (2019)

Ocean Biological Carbon pump exports between 9-13 Gt C y⁻¹

Laws et al. (2011)

<https://doi.org/10.4319/lom.2011.9.593>

If the OBCP were to be switched off, it has been estimated that atmospheric CO₂ concentrations would be 50% higher than they are today.

Parekh et al. (2006)

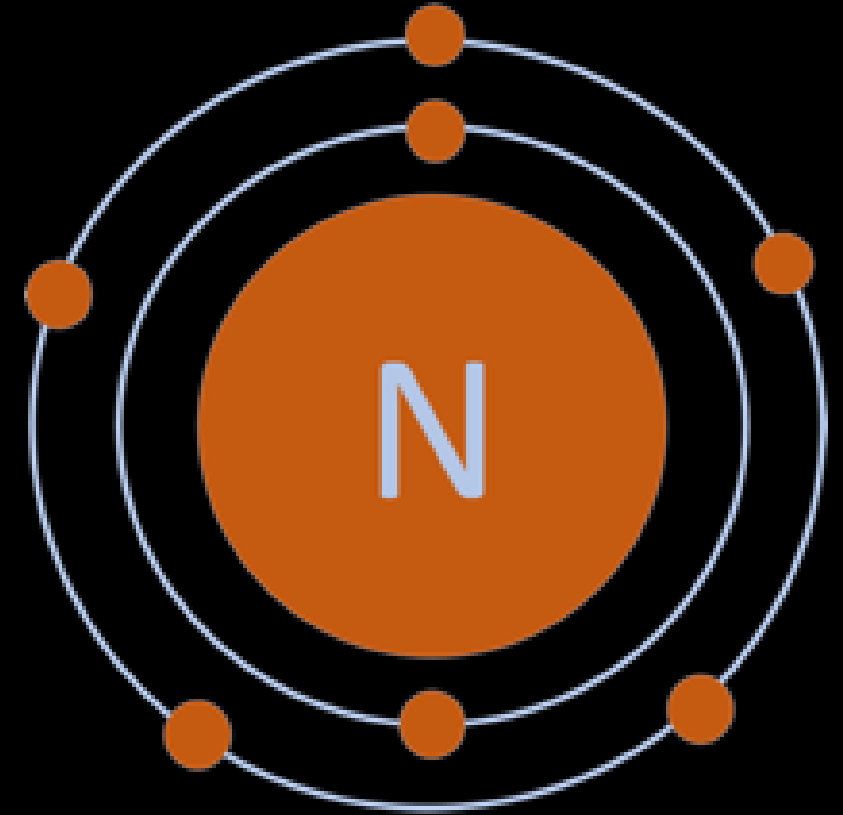
<https://doi.org/10.1029/2005GL025098>

Nitrogen (though incredibly abundant) is often a limiting nutrient for the growth for photosynthetic organisms. This is because the particular form of nitrogen found in air and that dissolves in water —nitrogen gas (N_2)—cannot be assimilated by most organisms.

Most algae and microbes require biologically available, or “fixed” forms of nitrogen, such as ammonia (NH_3^-) or ammonium (NH_4^+).

Fixed nitrogen is one of the most important growth-limiting nutrients for photosynthetic organisms (primary producers) such as algae. Nitrogen can also serve as an energy source or as an oxidant for marine bacteria and archaea.

The ocean absorbs nitrogen gas (N_2) from the atmosphere. It is the most abundant form of nitrogen in the ocean



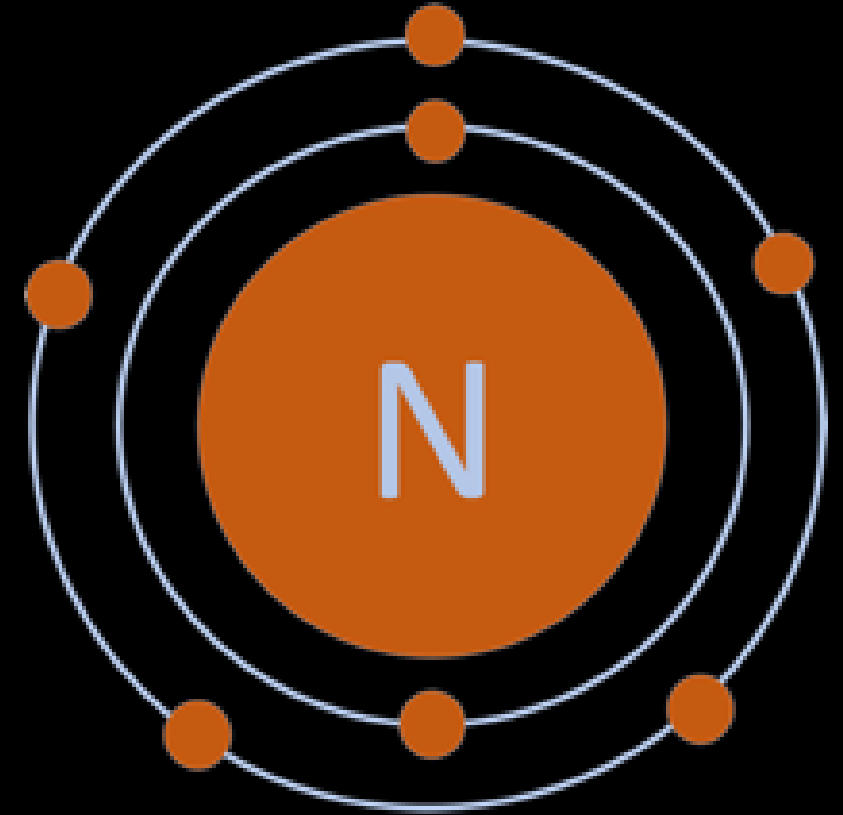
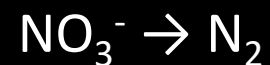
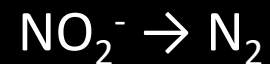
Nitrogen fixation is a process by which nitrogen in the water is converted into ammonium



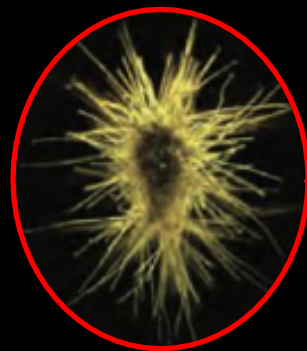
Nitrification is the biological oxidation of ammonium to nitrite (NO_2^-) and nitrate (NO_3^-)



Denitrification is the microbial conversion of nitrate or nitrite into nitrogen gas (N_2).

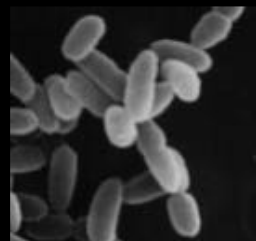


Trichodesmium



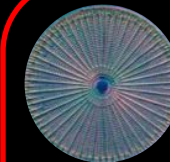
(Abby Helthoff, WHOI, IOCCG Report 15)

Bacteria

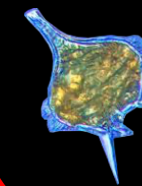


Credit:Kaden11a

Other Phytoplankton



Diatoms
(Carolina Biological Supply Company)



Dinoflagellates
(Tintinnidguy)

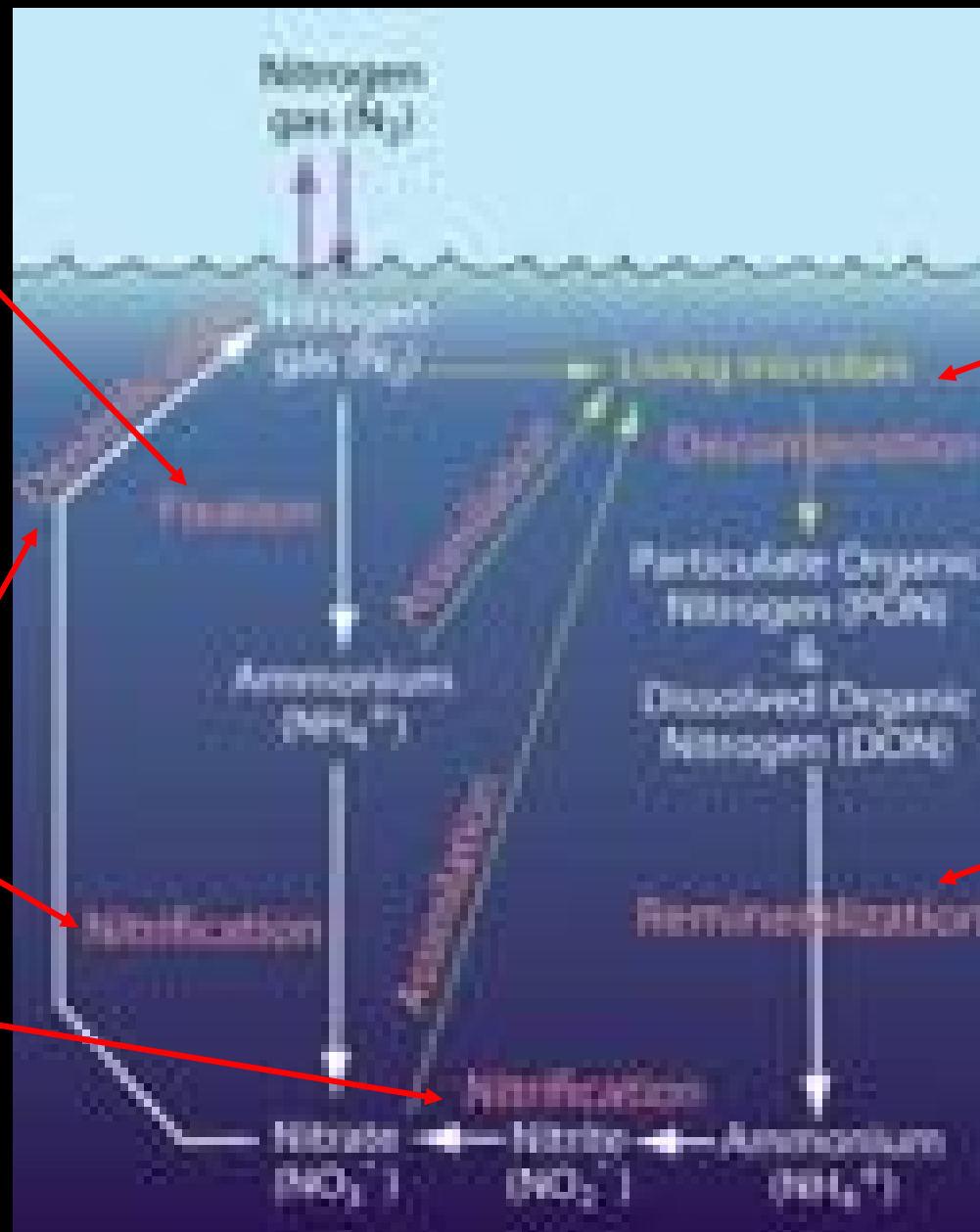
Zooplankton & Bacteria

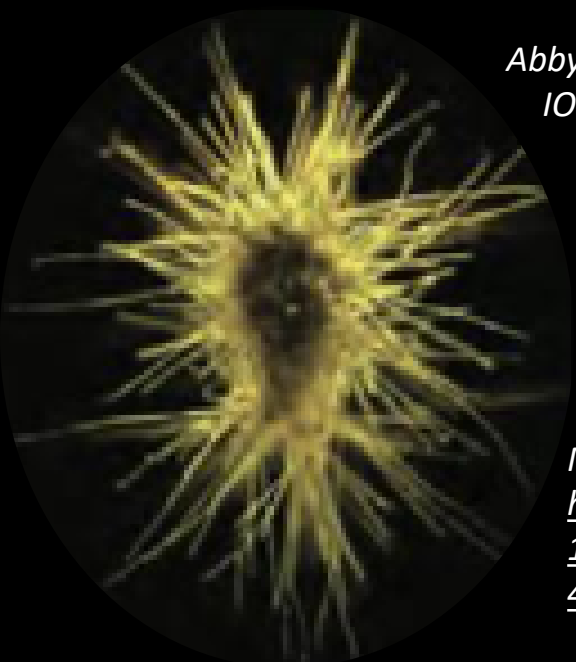


Credit:
Plymouth
Marine
Laboratory



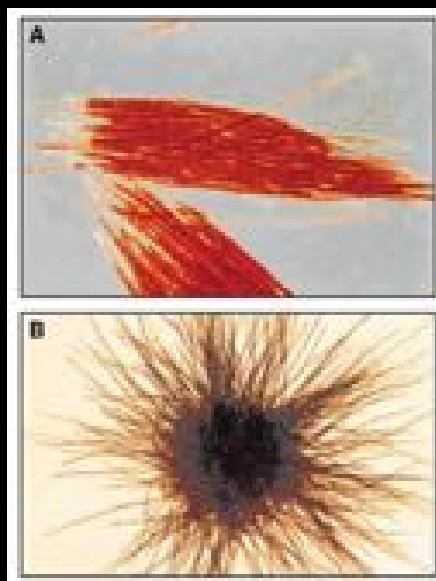
Credit:
Kaden11a





Abby Helthoff, WHOI,
IOCCG Report 15

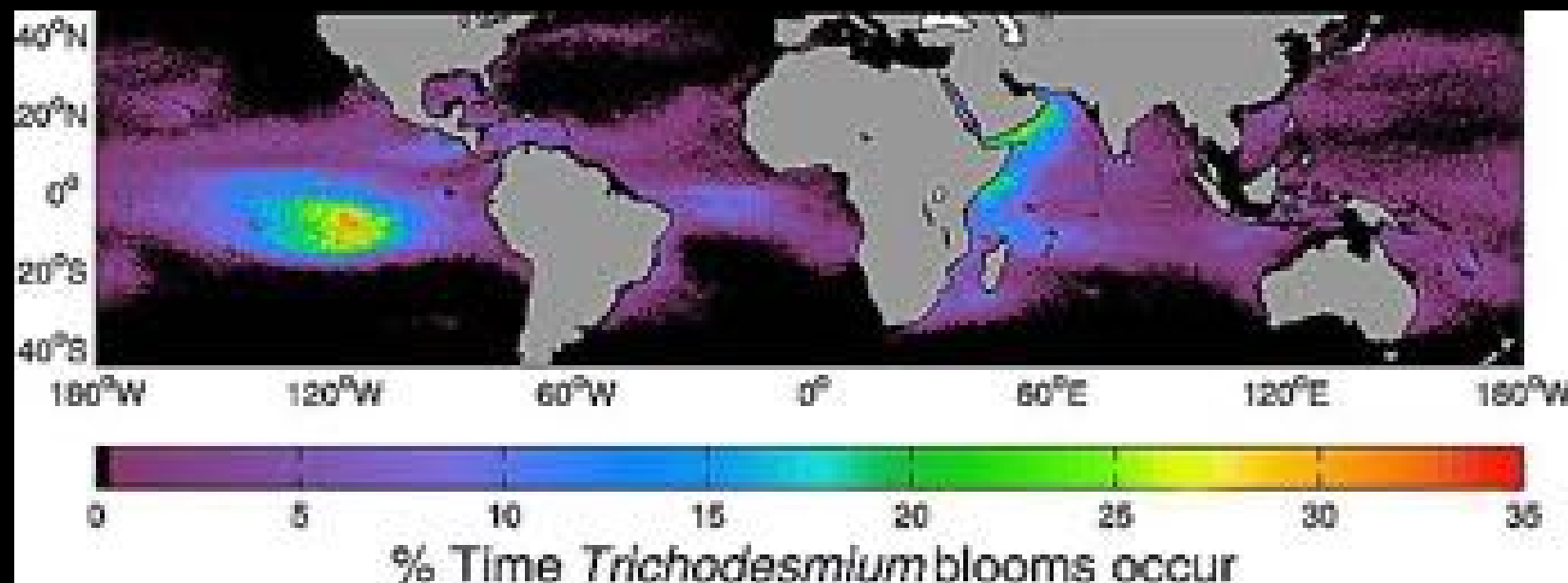
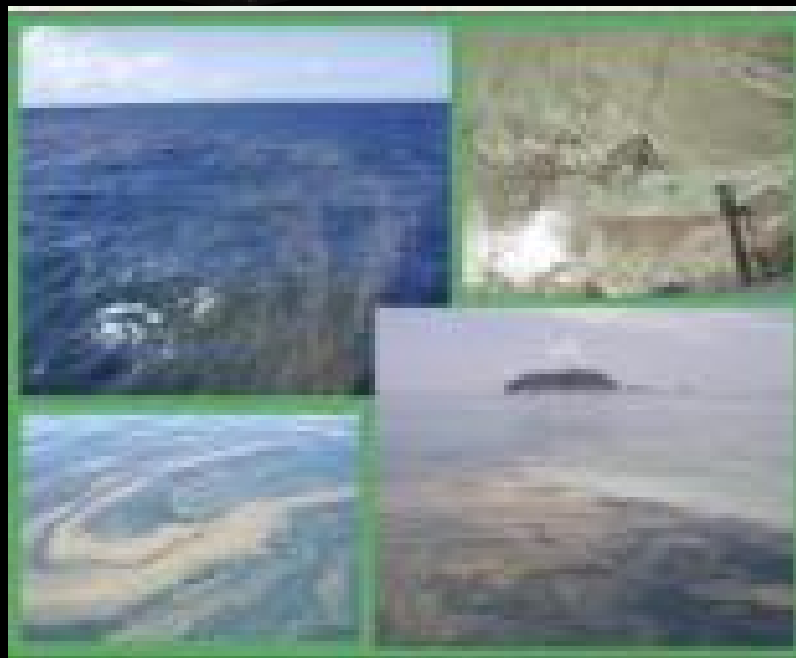
McKinna (2015)
<https://doi.org/10.1016/j.pocean.2014.12.013>



Trichodesmium, a Globally Significant Marine Cyanobacterium

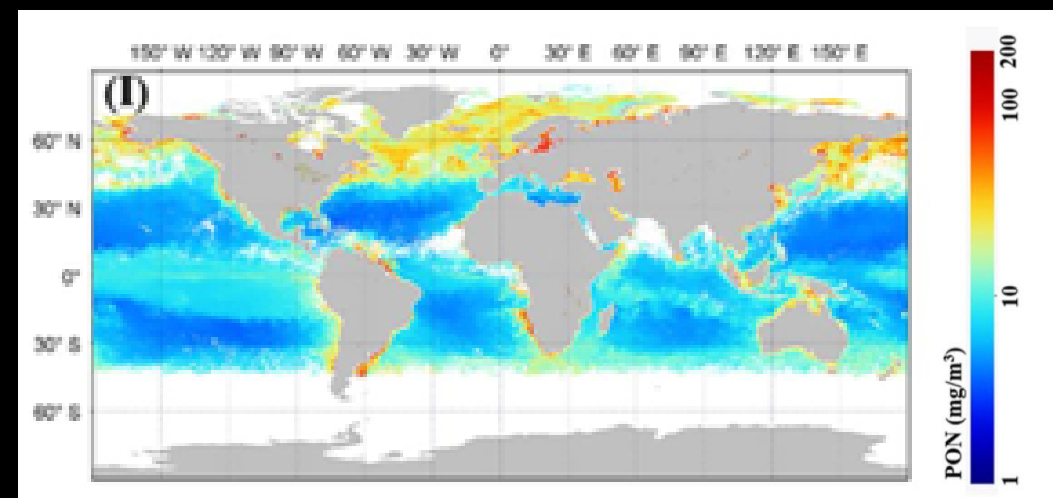
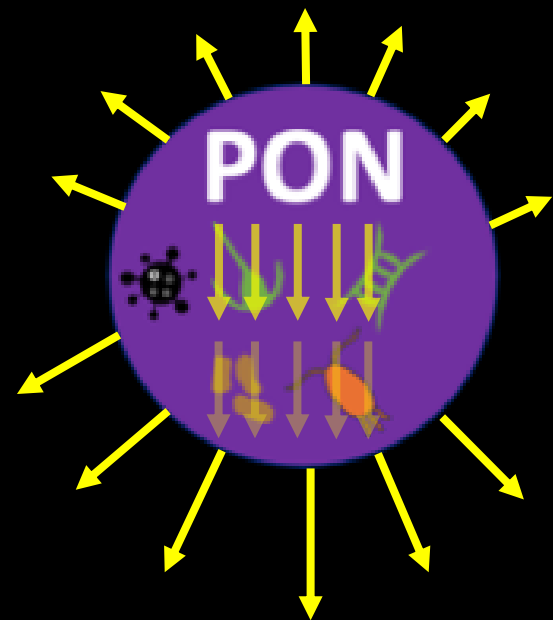
Douglas G. Capone, Jonathan P. Zehr, Hans W. Paerl,
Birgitta Bergman, Edward J. Carpenter

Capone et al. (1997) <https://doi.org/10.1126/science.276.5316.1221>

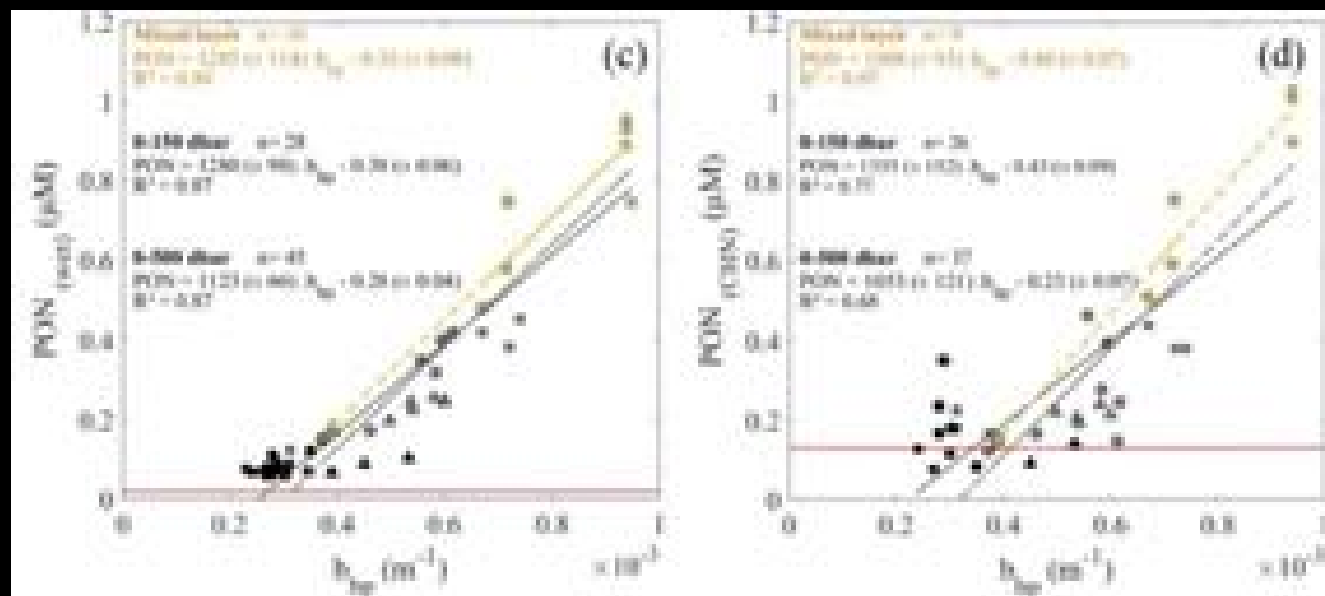


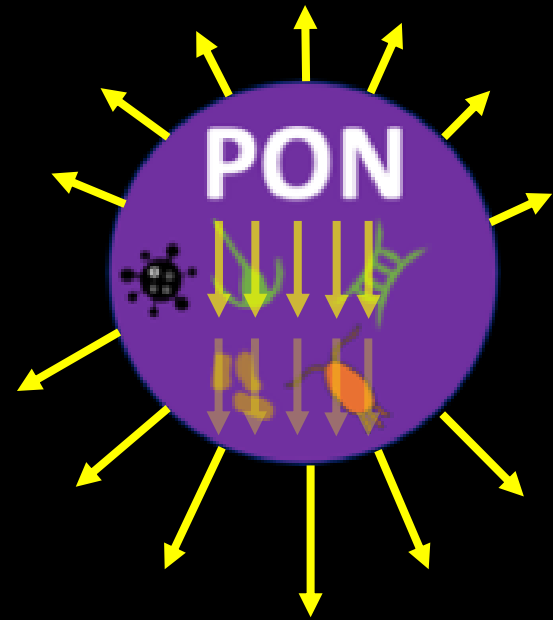
Subramaniam et al. (2001) [https://doi.org/10.1016/S0967-0645\(01\)00096-0](https://doi.org/10.1016/S0967-0645(01)00096-0)

Westberry and Siegel (2006). <https://doi.org/10.1029/2005GB002673>

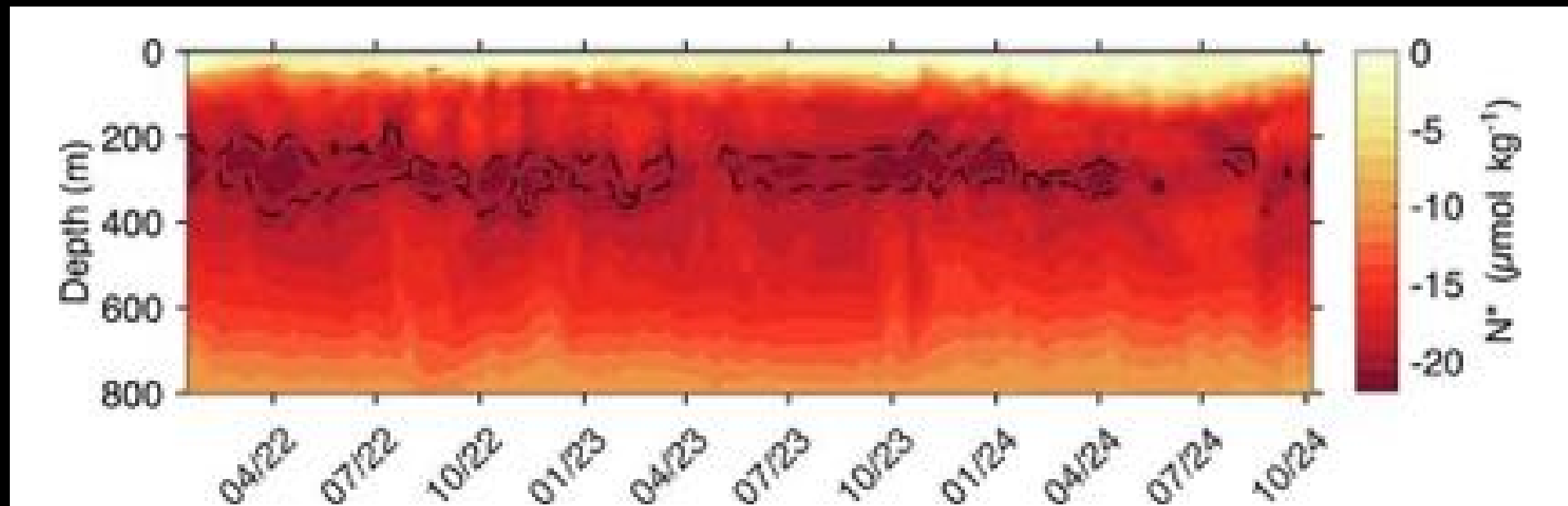


Wang et al. (2022) <https://doi.org/10.3389/fmars.2022.943867>





N^* indicates the deviation from Redfield stoichiometry, with more negative values reflecting progressive fixed nitrogen loss. Black dashed contour marks the $N^* = -20 \mu\text{mol kg}^{-1}$ threshold.



Bif et al. (2026) <https://doi.org/10.1038/s43247-026-03410-5>

GLACIAL-INTERGLACIAL CO₂ CHANGE: THE IRON HYPOTHESIS

John N. Martin

Moss Landing Marine Laboratories
Moss Landing, California

Abstract. Several explanations for the 200 to 300 ppm glacial/interglacial change in atmospheric CO₂ concentrations deal with variations in southern ocean phytoplankton productivity and the related use or release of major plant nutrients. An hypothesis is presented herein in which arguments are made that new productivity in today's southern ocean (7.4×10^{13} g yr⁻¹) is limited by iron deficiency, and hence the phytoplankton are unable to take advantage of the excess surface nitrate/phosphate that, if used, could result in total southern ocean new production of 2.3×10^{15} g C yr⁻¹. As a consequence of Fe-limited new productivity, Holocene interglacial CO₂ levels (preindustrial) are as high as they were during the last interglacial (~280 ppm). In contrast, atmospheric dust Fe supplies were 50 times higher during the last glacial maximum (LGM). Because of this Fe enrichment, phytoplankton growth may have been greatly enhanced, larger amounts of upwelled nutrients may have been used, and the resulting stimulation of new productivity may have contributed to the LGM drawdown of atmospheric CO₂ to levels of less than

a Global atmospheric CO₂



b Iron flux to the atmosphere



*“Give me
half a
tanker of
iron and I'll
give you
the next ice
age”*

*John
Martin*

Martin et al. (1990)

<https://agupubs.onlinelibrary.wiley.com/doi/epdf/10.1029/PA005i001p00001>

Stoll et al.(2020) <https://doi.org/10.1038/d41586-020-00393-x>

Considered among
the biggest
achievements in
biological
oceanography in
the 20th century
(Barber & Hilting
2000)

Iron Hypothesis

The iron issue is an example of classic science progress:

- a. There was a nagging question.
- b. A tentative explanation was advanced.
- c. Available data did not support the explanation.
- d. The data, however, were suspect.
- e. A technical (analytical) breakthrough was made.
- f. The new data suggested an hypothesis.
- g. The hypothesis was tested and confirmed.
- h. Textbooks had to be revised.

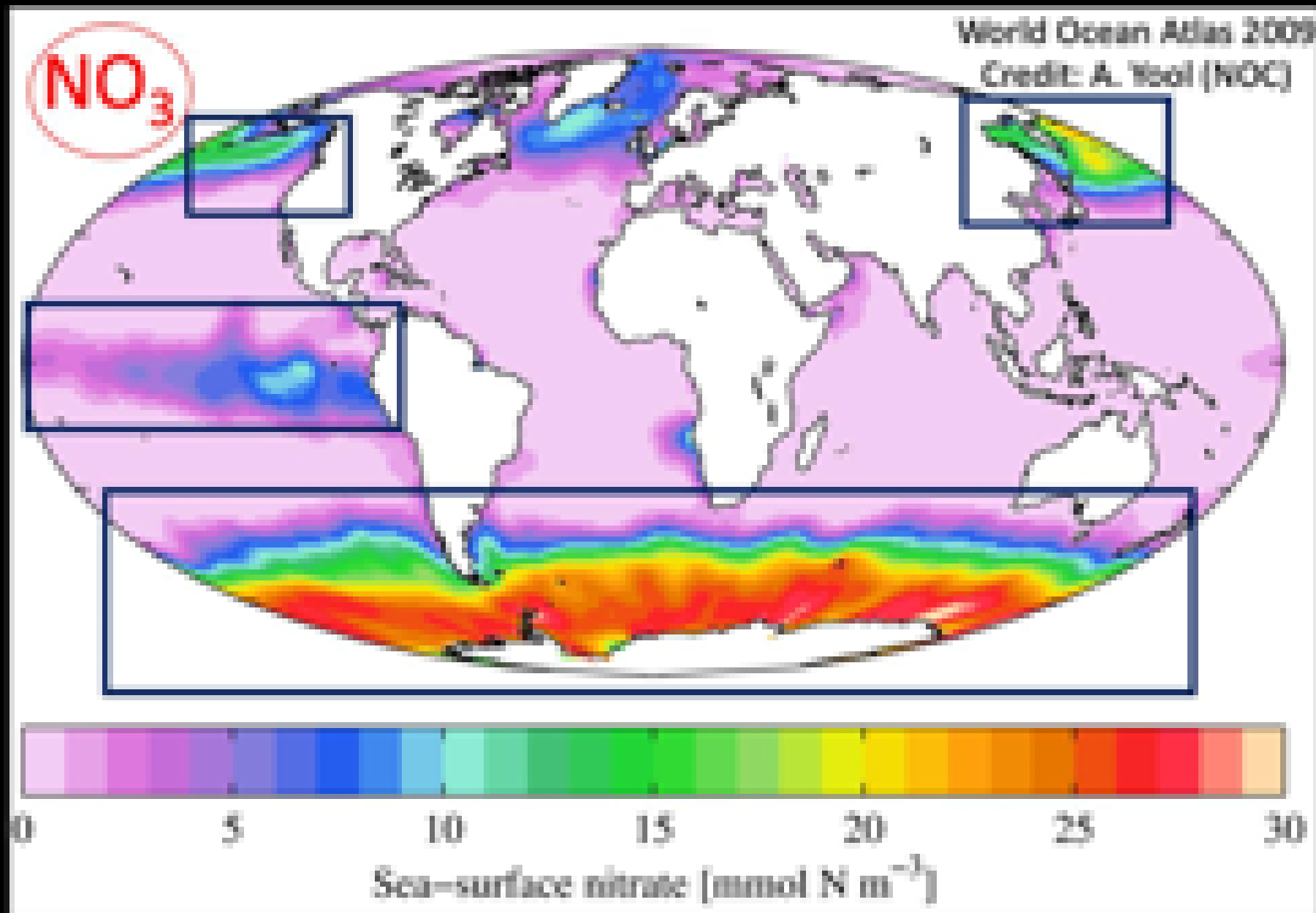
For the iron issue the nagging question was: Why do excess plant nutrients persist in the surface ocean in certain regions such as the Antarctic, equatorial Pacific, and North-east Pacific? For 50 years there had been speculation that iron limitation might be a factor, but measurements showed there was abundant iron in seawater.

John Martin set out to improve the analytical chemistry of iron, and when he had done so, he found that iron was much less abundant in the ocean than previously thought.

*Alongside
discovery of
hydrothermal
vents and seeing
the ocean for
the first time
though satellite
ocean colour!*

Barber & Hilting (2000)

https://www.ncbi.nlm.nih.gov/books/NBK208814/pdf/Bookshelf_NBK208814.pdf



**Micro-
nutrients**

Fe

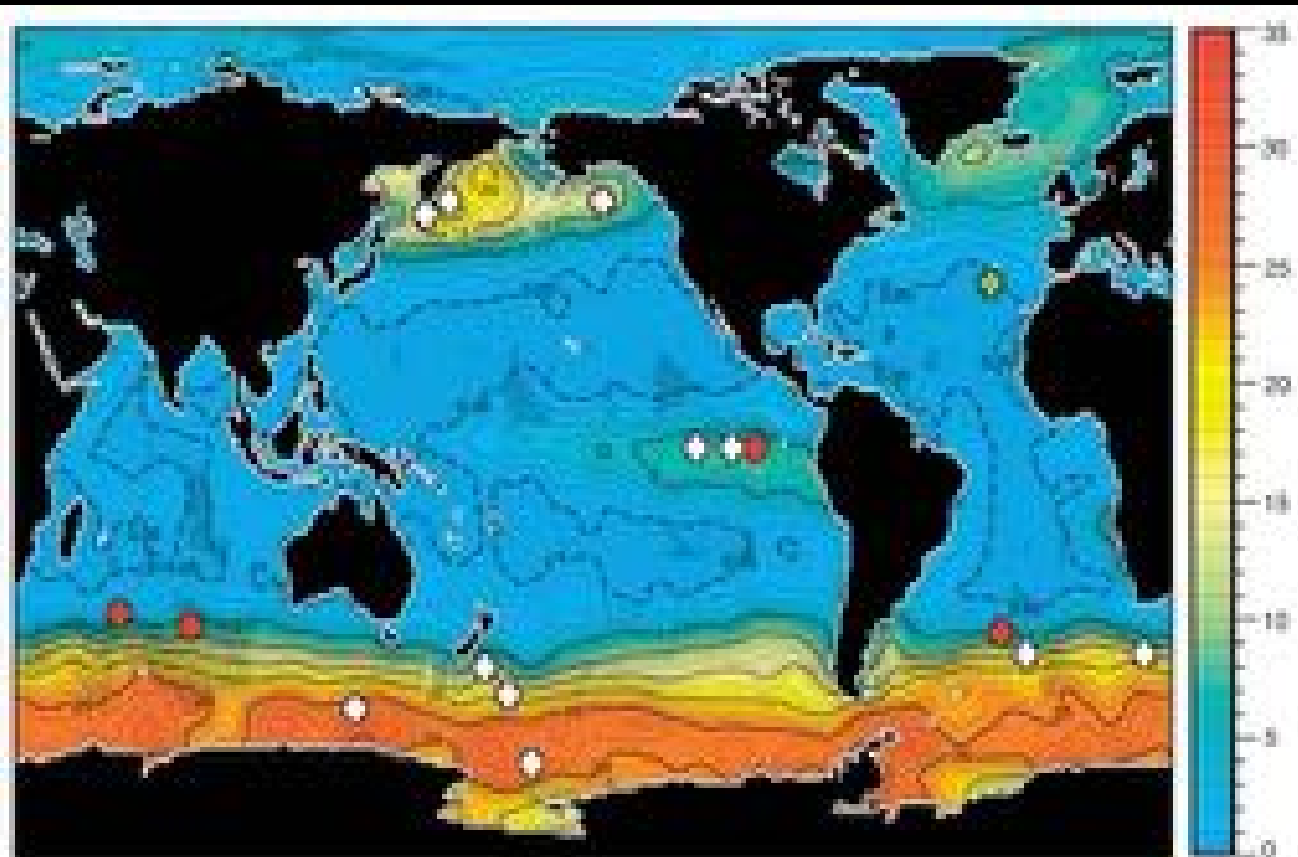


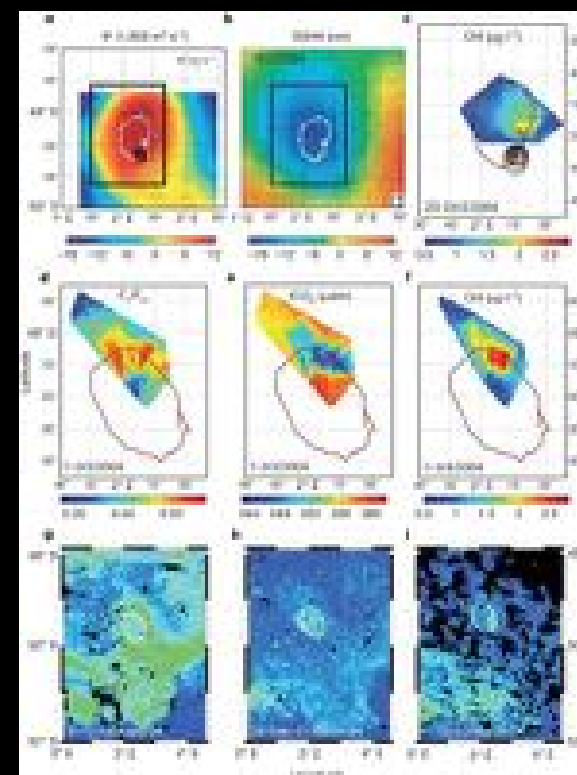
Fig. 1. Annual surface mixed-layer nitrate concentrations in units of $\mu\text{mol liter}^{-3}$ (M), with approximate site locations of FeAOs (white crosses), FeNOs (red crosses), and a joint Fe and P enrichment study of the subtropical IRLC Atlantic Ocean (FeP; green cross). FeAOs shown are SEEDS I and II (northwest Pacific; same site but symbols are offset), SERIES (northeast Pacific), IronEX I and II (equatorial Pacific; IronEX II is to the left), IronEX and EIFEX (Atlantic polar waters; EIFEX is directly south of Africa), SOFEX (polar waters south of Australia), SOFEX-S (polar waters south of New Zealand), SOFEX-N (subpolar waters south of New Zealand), and SAGE (subpolar waters nearest to New Zealand). FeNO sites shown are the Galapagos Plume (equatorial Pacific), Antarctic Polar Front (polar Atlantic waters), and the Crozet and Kerguelen plateaus (Indian sector of Southern Ocean; Crozet is to the left of Kerguelen). For the geographical positions of the FeAOs, see IRLC. FeP investigated whether N-fixing phytoplankton are simultaneously limited by Fe and P; see Table 1.

ARTICLE

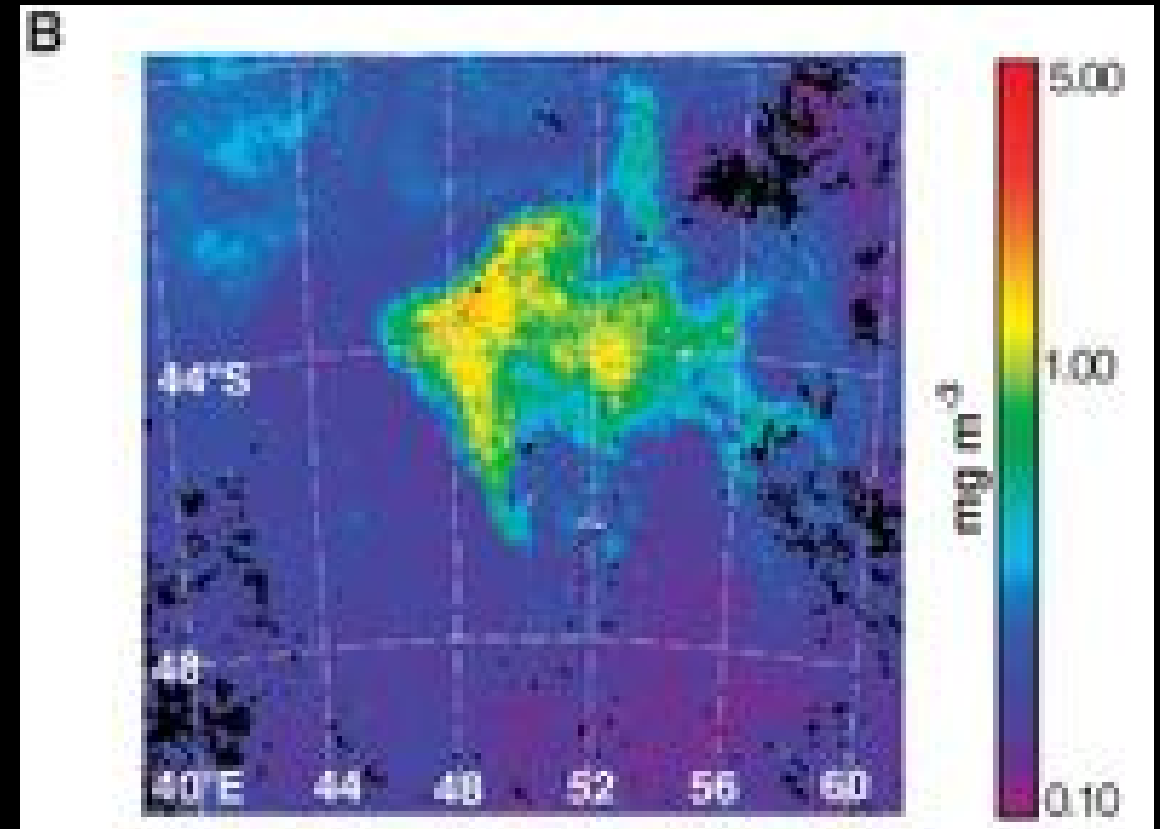
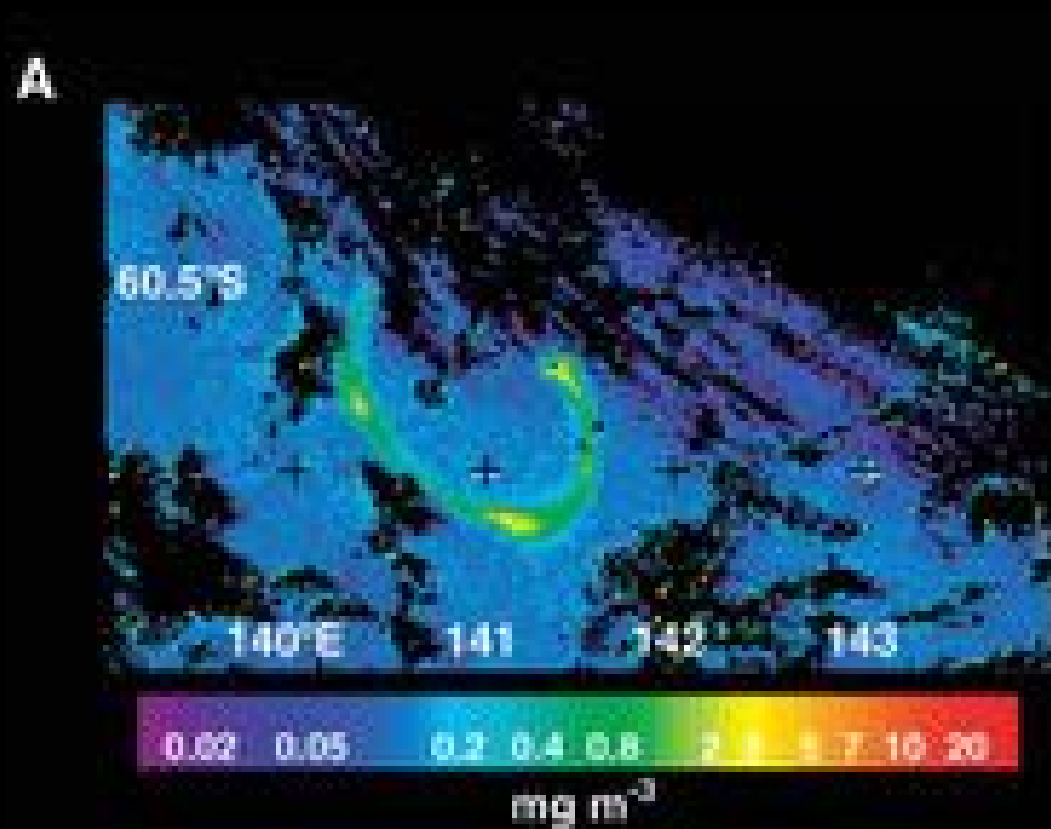
doi:10.1038/nature11229

Deep carbon export from a Southern Ocean iron-fertilized diatom bloom

Victor Smetacek^{1,2*}, Christine Klaus^{3*}, Volker H. Strass¹, Philipp Assmy^{1,2}, Marina Montresor⁴, Boris Ojczewski^{1,2}, Nicolas Savoye^{5,6}, Adrian Webb⁷, Francesco d'Ovidio⁸, Jesús M. Arrieta^{9,10}, Ulrich Barthmann^{11,12}, Richard Bellerby^{13,14}, Gry Mline Berg¹⁵, Peter Cassa^{16,17}, Santiago Gonzalez¹⁸, Joachim Henjes^{1,19}, Gerhard J. Herndl^{20,21}, Linn J. Hoffmann²², Harry Leach²³, Martin Losch¹, Matthew M. Mills²⁴, Craig Neill^{1,25}, Ilka Peeken^{1,26}, Rüdiger Röttgers²⁷, Oliver Sachs^{1,28}, Eberhard Sauter¹, Malke M. Schmidt²⁹, Hil Schwarze^{1,26}, Anja Tschirggen¹ & Dieter Wolf-Gladrow¹

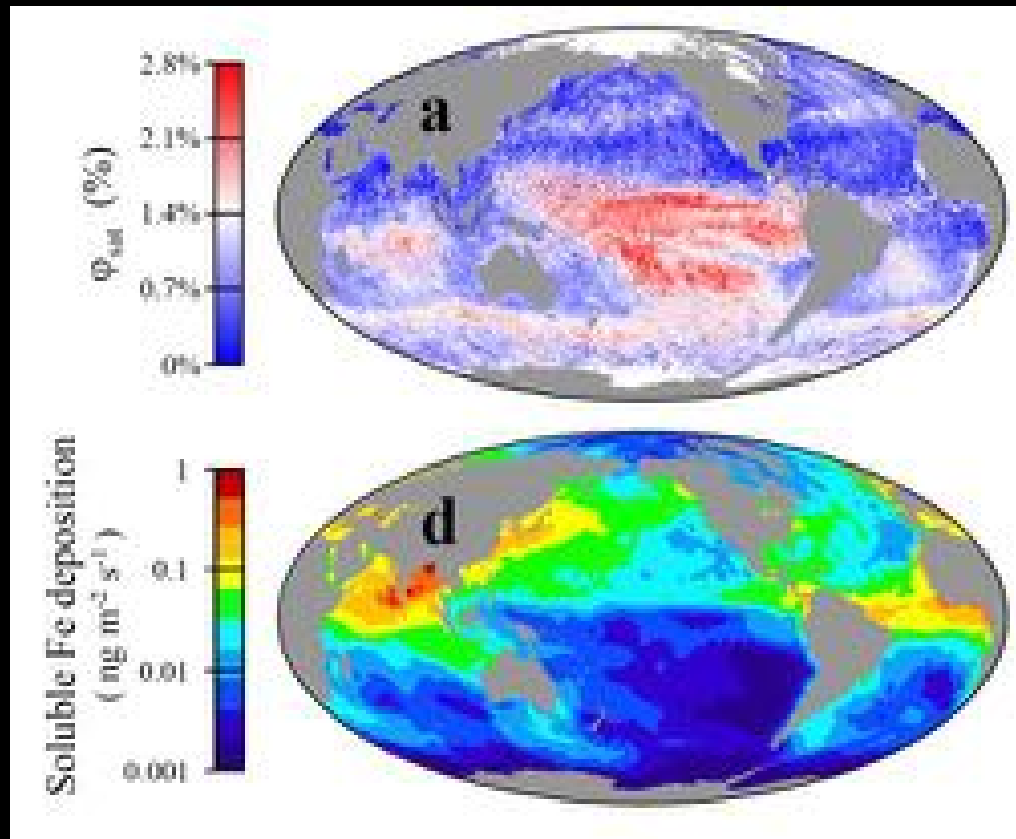


“at least half the bloom biomass sank far below a depth of 1,000 metres and that a substantial portion is likely to have reached the sea floor”



- (A) Satellite image of a purposeful in situ Southern Ocean FeAX (SOIREE)
- (B) An FeNX near Crozet within the HNLC Southern Ocean, where naturally occurring blooms are evident from remote sensing

Passive and active fluorescence used as an optical measure of iron stress in phytoplankton



Behrenfeld et al. (2009)

<https://doi.org/10.5194/bg-6-779-2009>

Article

Persistent equatorial Pacific iron limitation under ENSO forcing

<https://doi.org/10.1038/s41586-023-06439-0>

Received: 13 January 2022

Accepted: 14 July 2023

Thomas J. Browning^{1,2}, Mak A. Saito³, Shungudzemawo P. Garaba⁴, Xuechao Wang¹, Eric P. Achterberg⁵, C. Mark Moore⁶, Anja Engel¹, Matthew R. Moliver⁷, Dawn Moran⁸, Daniela Voss⁹, Oliver Ziehm^{10,11} & Alessandro Tagliabue⁸

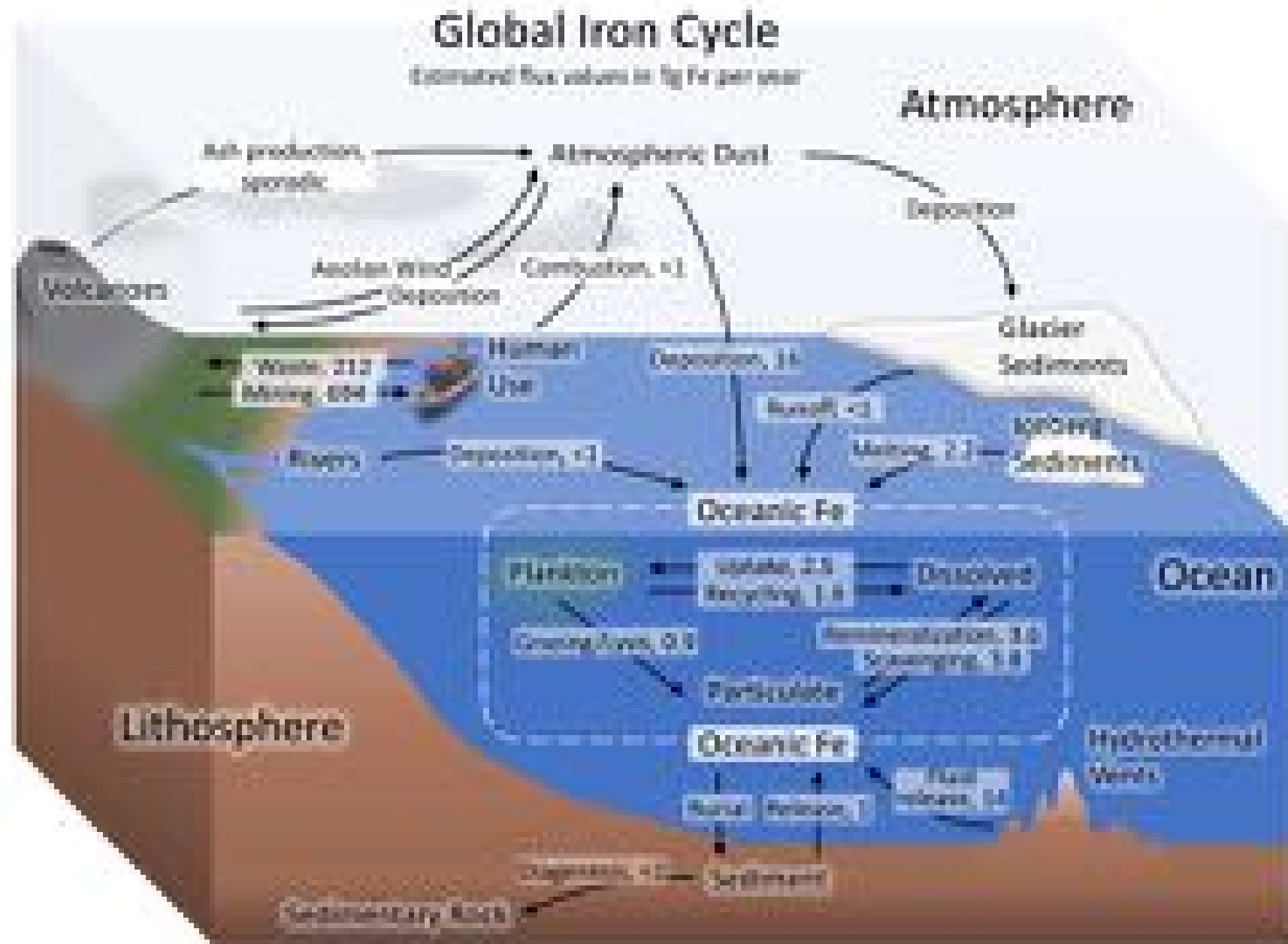
PRIMARY PRODUCTIVITY

Multidecadal trend of increasing iron stress in Southern Ocean phytoplankton

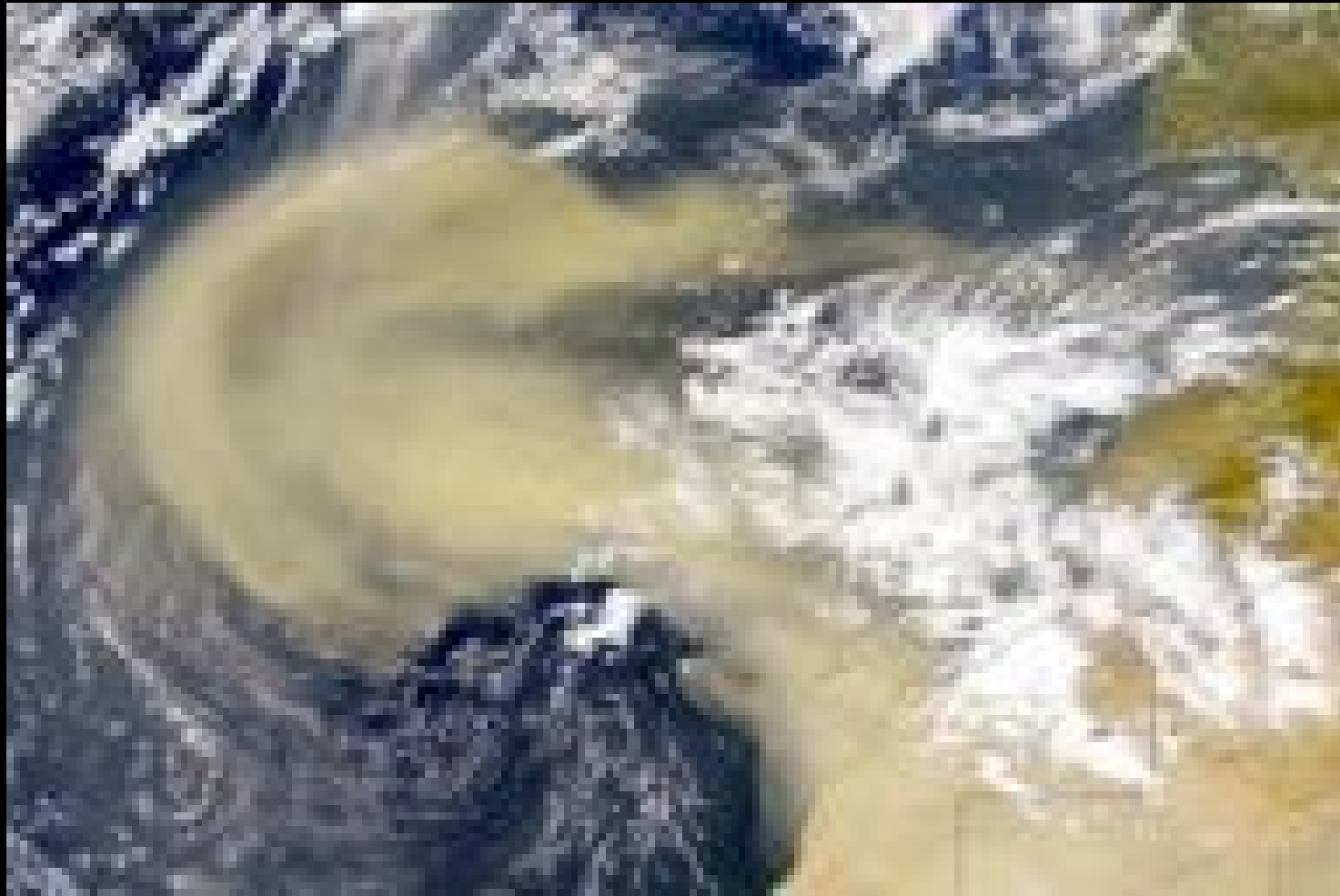
Thomas J. Ryan-Keogh^{1,2}†, Sandy J. Thomalla^{1,2}†, Pedro M. S. Monteiro^{1,3}, Alessandro Tagliabue⁴

Browning et al. (2023) <https://doi.org/10.1038/s41586-023-06439-0>

Ryan-Keogh et al. (2023) <https://doi.org/10.1126/science.abl5237>

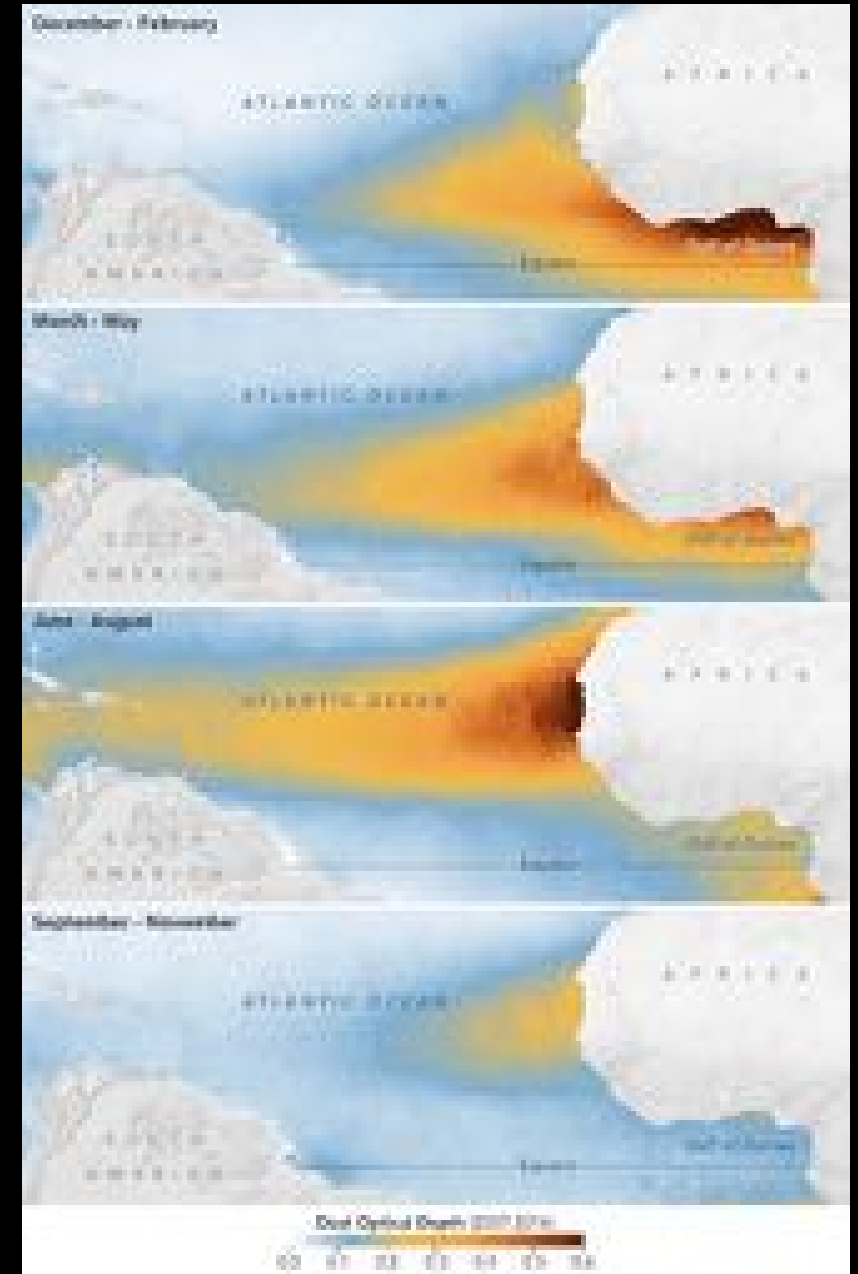


Credit
BeccaMGM
https://en.wikipedia.org/wiki/Iron_cycle#/media/File:Iron_cycle7.png



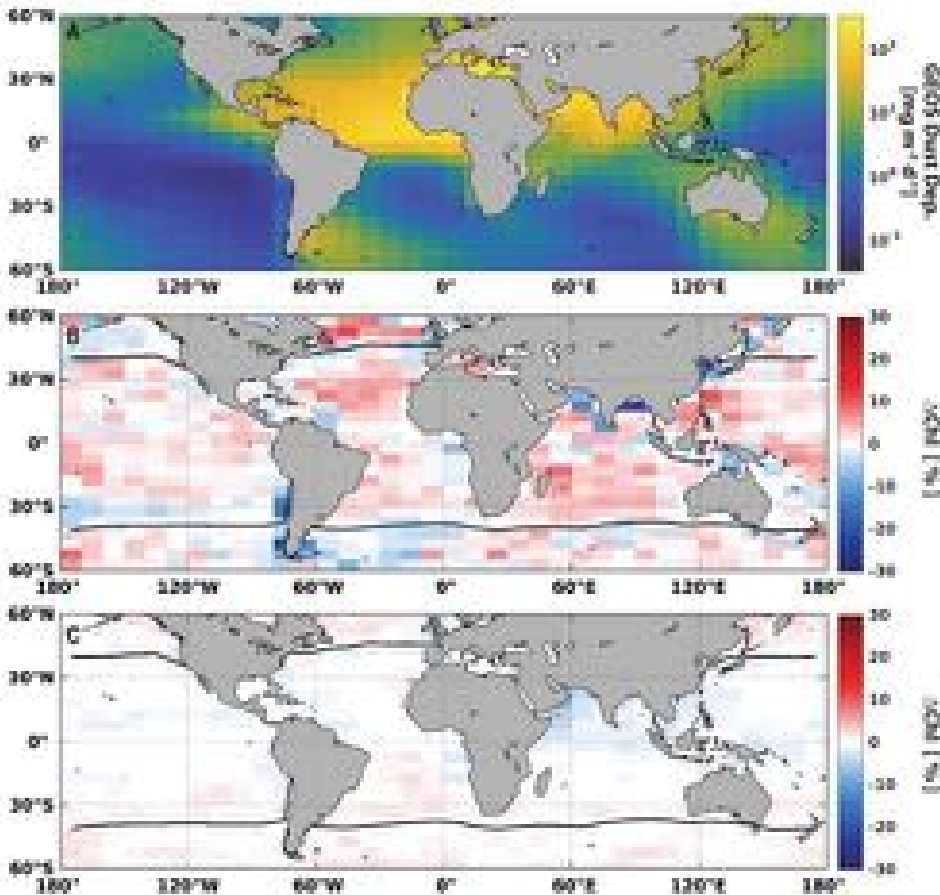
When winds lift particles up, they can spread across thousands of miles, circling the planet and fertilizing key parts of oceans with critical nutrients like iron.

Credit <https://www.nesdis.noaa.gov/content/dirt-atmospheric-dust>
<https://earthobservatory.nasa.gov/images/146122/a-dusty-journey>



MARINE PRODUCTIVITY

Atmospheric nourishment of global ocean ecosystems

T. K. Westberry^{1*}, M. J. Behrenfeld¹, Y. R. Shi^{2,3}, H. Yu³, L. A. Remer^{2,4}, H. Bian^{2,3}

Westberry et al. (2023)

<https://doi.org/10.1126/science.abq5252>PNAS
nexus

PNAS Nexus, 2024, 3, pgae386

<https://doi.org/10.1093/pnasnexus/pgae386>

Advance access publication 1 October 2024

Research Report

An exceptional phytoplankton bloom in the southeast Madagascar Sea driven by African dust deposition

John A. Gittings¹, Giorgio Dall'Omo², Weiye Tang³, Joan Lloret⁴, Fatma Jebri⁵, Eleni Livanou⁶, Francesco Nencioli⁷, Sofia Dammaraki⁸, Jason Theodorou⁹, Robert J. W. Kerwin¹⁰, Meric Srokosz¹¹, Nicolas Cassar^{12*} and Dionysios E. Raitsos^{13*}

Gittings et al (2024)

<https://doi.org/10.1093/pnasnexus/pgae386>

Article

Widespread phytoplankton blooms triggered by 2019–2020 Australian wildfires

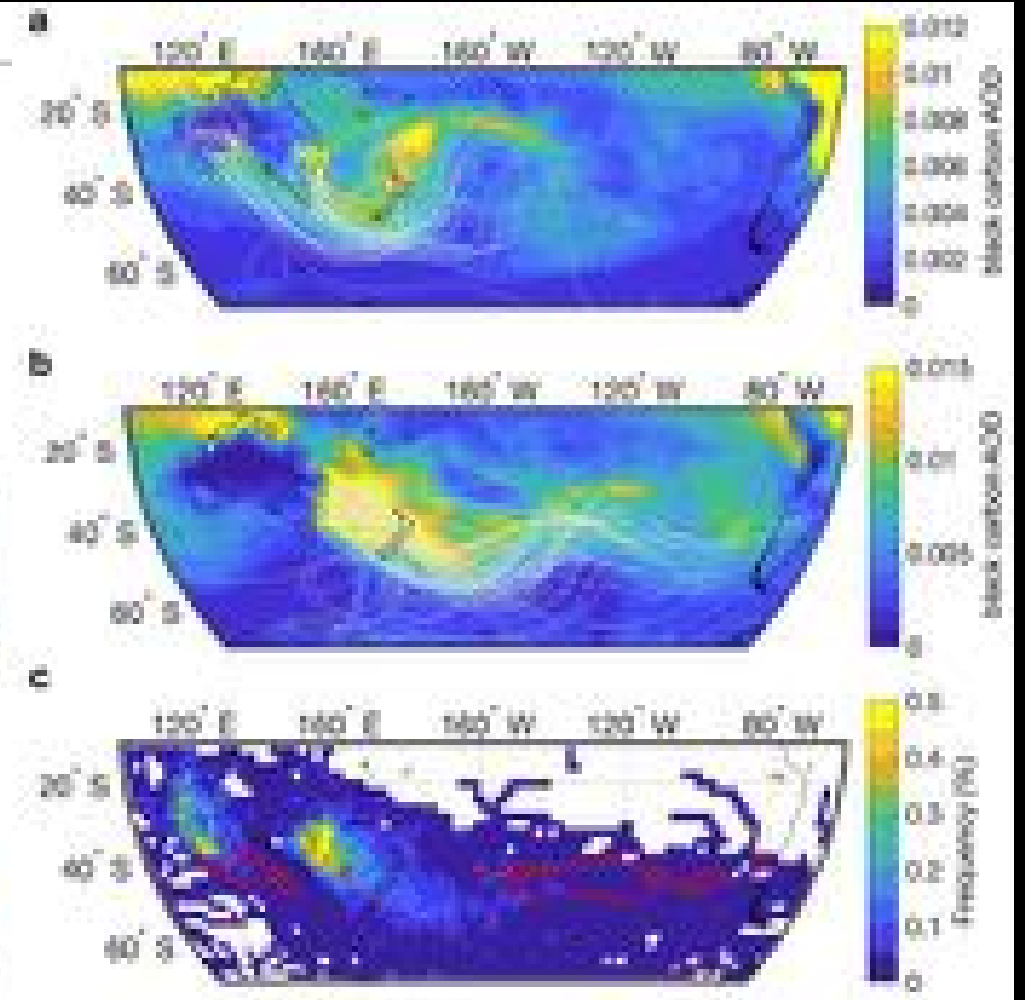
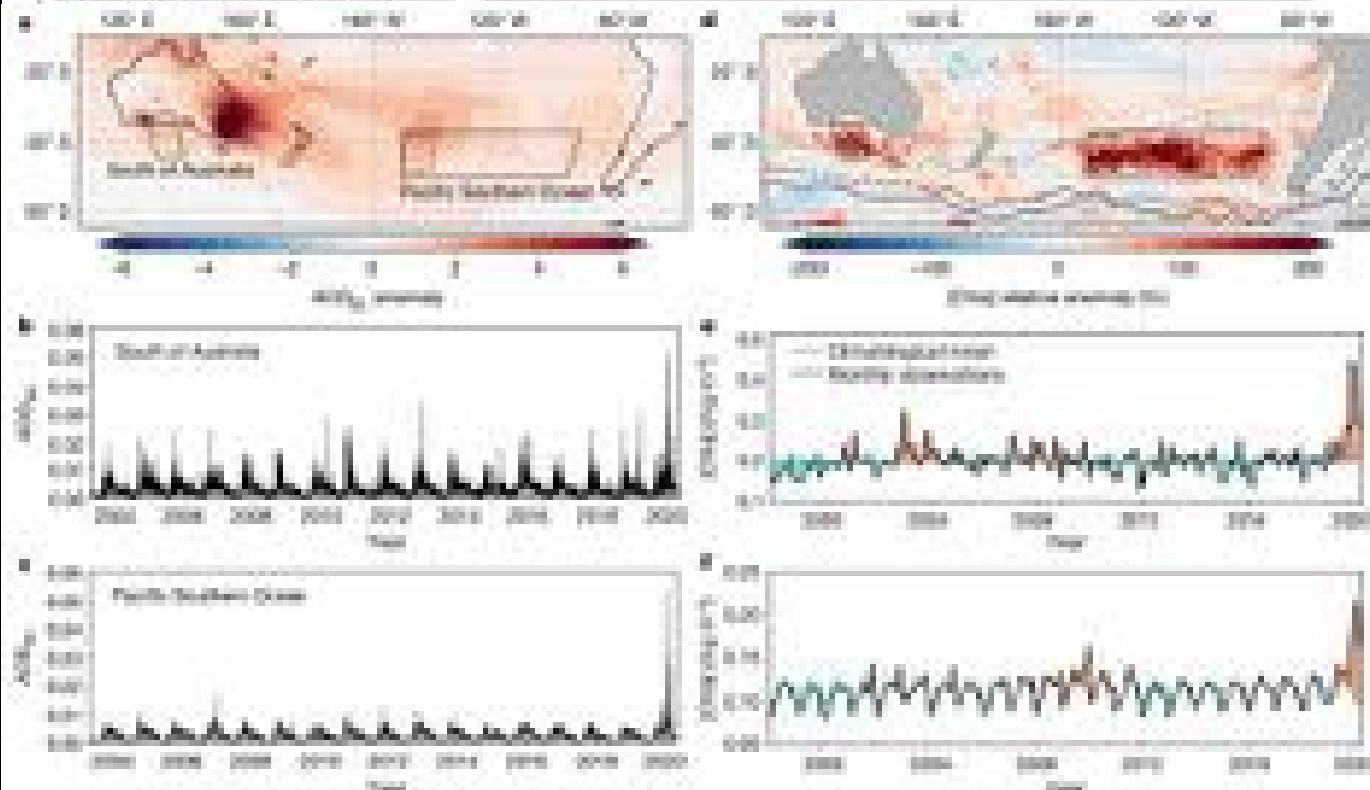
<https://doi.org/10.1038/s41586-021-03805-8>

Received: 25 August 2020

Accepted: 5 July 2021

Published online: 15 September 2021

Wenli Tang^{1,2*}, Joan Lloret^{1,2*}, Jakob Weiss^{3,4}, Morgane M. G. Perron⁵, Sara Bassat⁶, Zuchuan Li⁷, Shubha Sathyendranath⁸, Thomas Jackson⁹, Estrella Sanz Rodríguez⁶, Bernadette C. Proenca⁶, Andrew R. Bowie¹⁰, Christina Schallenberg¹¹, Peter G. Strutton¹², Richard Matear¹³ & Nicolas Cassar^{14,15}



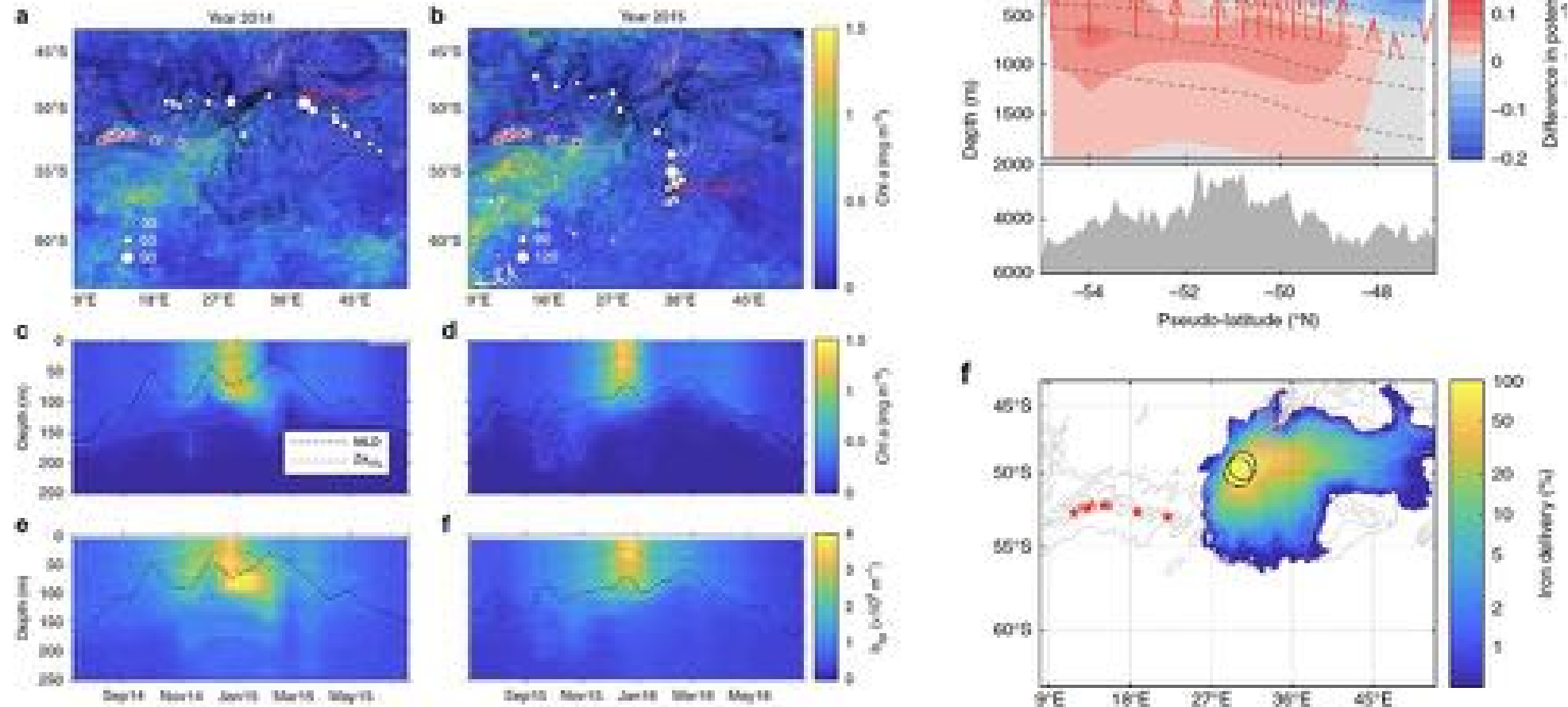
Tang et al. (2021)
<https://doi.org/10.1038/s41586-021-03805-8>

ARTICLE

<https://doi.org/10.1038/s41467-019-09973-6>

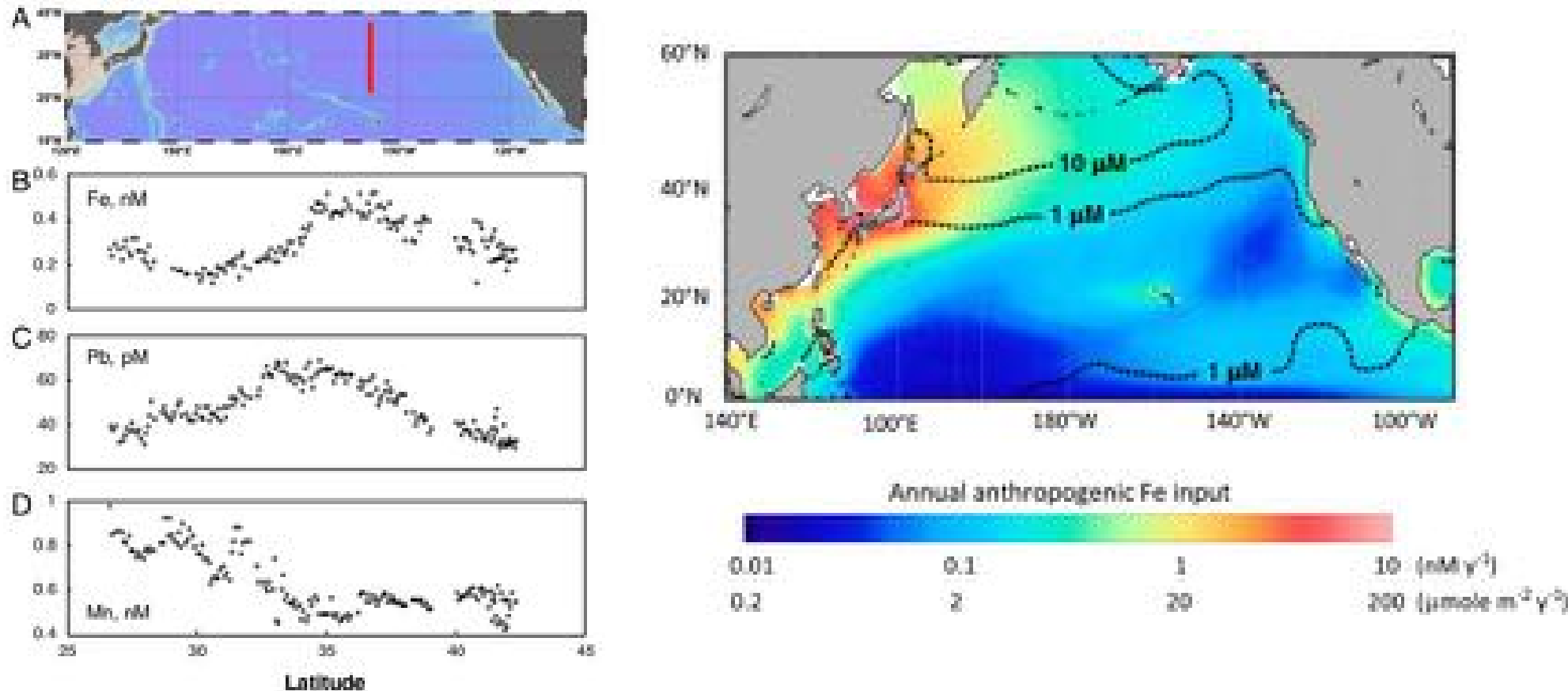
OPEN

Hydrothermal vents trigger massive phytoplankton blooms in the Southern Ocean



Anthropogenic Asian aerosols provide Fe to the North Pacific Ocean

Paulina Pinedo-González^{a,b,1}, Nicholas J. Hawco^{a,c}, Randelle M. Bundy^d, E. Virginia Armbrust^d, Michael J. Follows^e, B. B. Cael^f, Angelicque E. White^c, Sara Ferrón^c, David M. Karl^c, and Seth G. John^a



Anthropogenic Fe contributes 21-59% of dissolved Fe measured between 35° and 40°N

Lecture outline

1. Introduction to marine biogeochemical cycles
2. The Water cycle
3. The Redfield Ratio
4. The Carbon cycle
5. The Ocean Biological Carbon Pump
6. The Nitrogen cycle
7. The Iron cycle