

Ocean colour remote sensing at high latitudes

Marcel Babin

Takuvik Joint International Laboratory
Laval University (Canada) & CNRS (France)

www.takuvik.ulaval.ca

Why a lecture on this specific topic?

- Increasing attention is paid to what's going on there
- The use of ocean color remote sensing faces specific challenges at high latitudes
 - Low Sun elevation
 - High cloudiness
 - Sea ice
 - Peculiar optical properties
 - Peculiar phytoplankton properties
 - Vertical distribution
 - Optical properties
 - Physiology

Outline

1. Ocean colour remote sensing in polar seas

- 1.1. Ocean and sea ice in Arctic and Antarctic: relevant features
- 1.2. Seawater optical properties
- 1.3. Retrieval of ocean properties from ocean colour
 - 1.3.1. Atmospheric corrections
 - 1.3.2. Contamination of the signal by sea ice
 - 1.3.3. Retrieval of IOPs and AOPs, and biogeochemically relevant variables
- 1.4. Availability of data as favoured by polar orbits and limited by elevated cloudiness

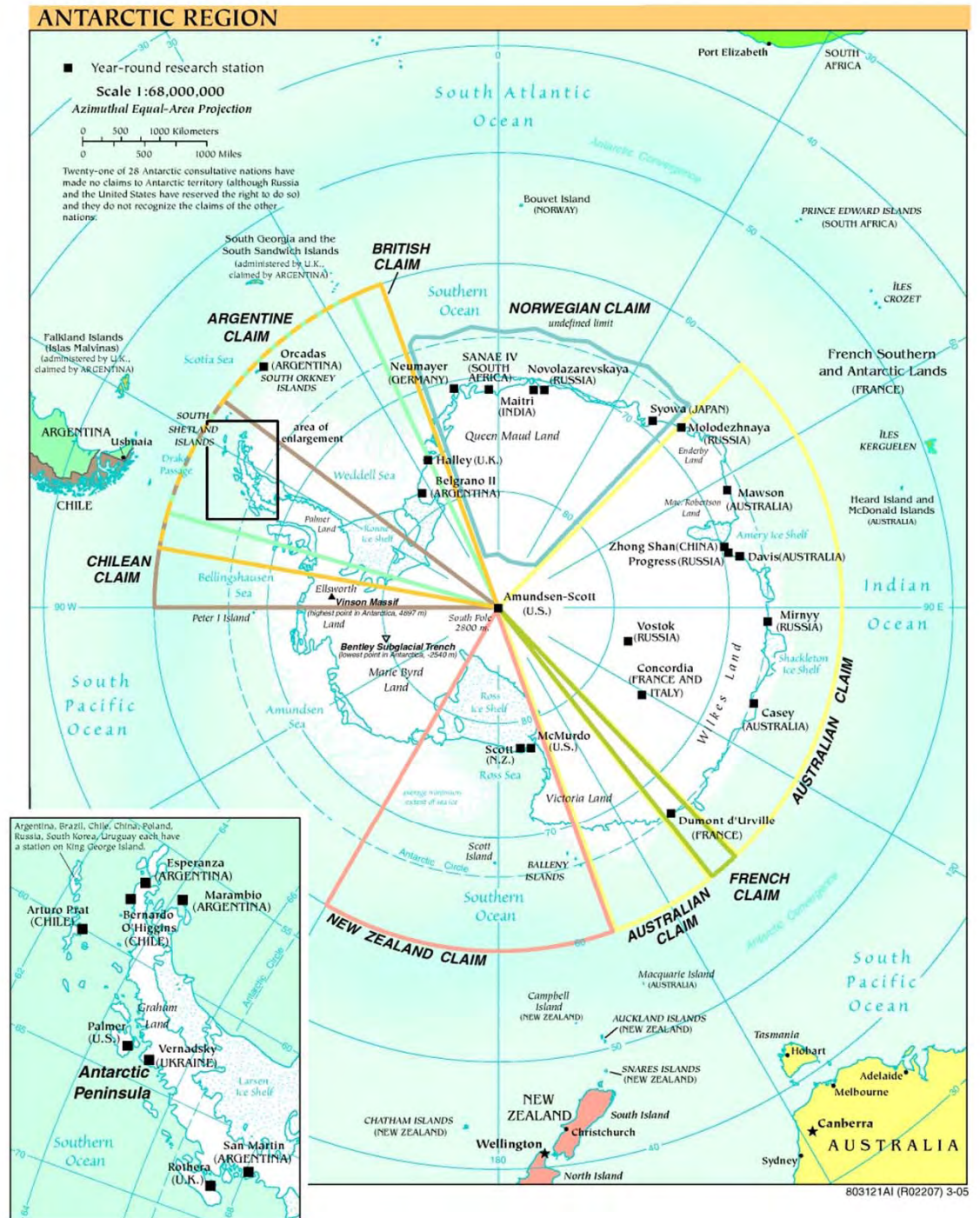
2. Primary production estimates from OC in polar seas

- 2.1. PP model and validation: an example
- 2.2. Results from PP models

1. Ocean colour remote sensing in polar seas

1.1. Ocean and sea ice in Arctic and Antarctic:
relevant features

Geography



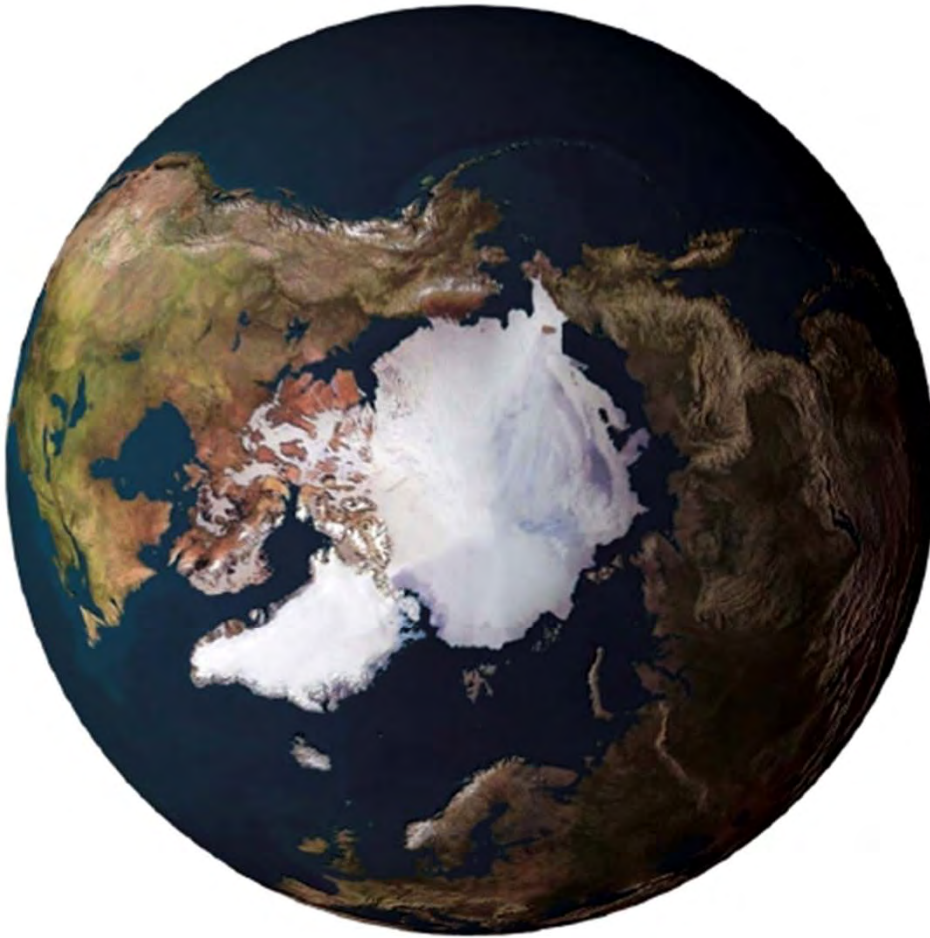
Geography



Geography

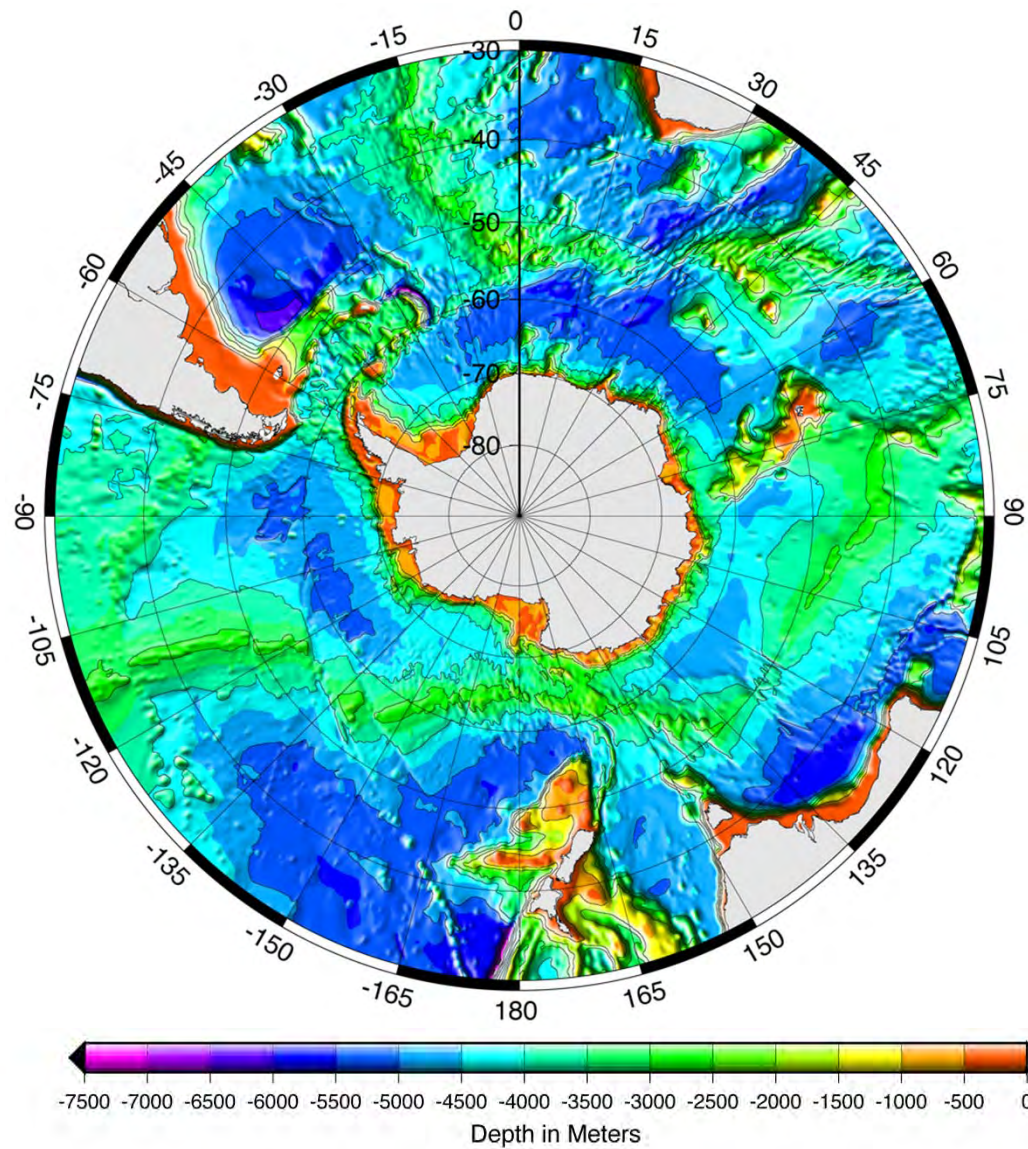


Arctic vs. Antarctic



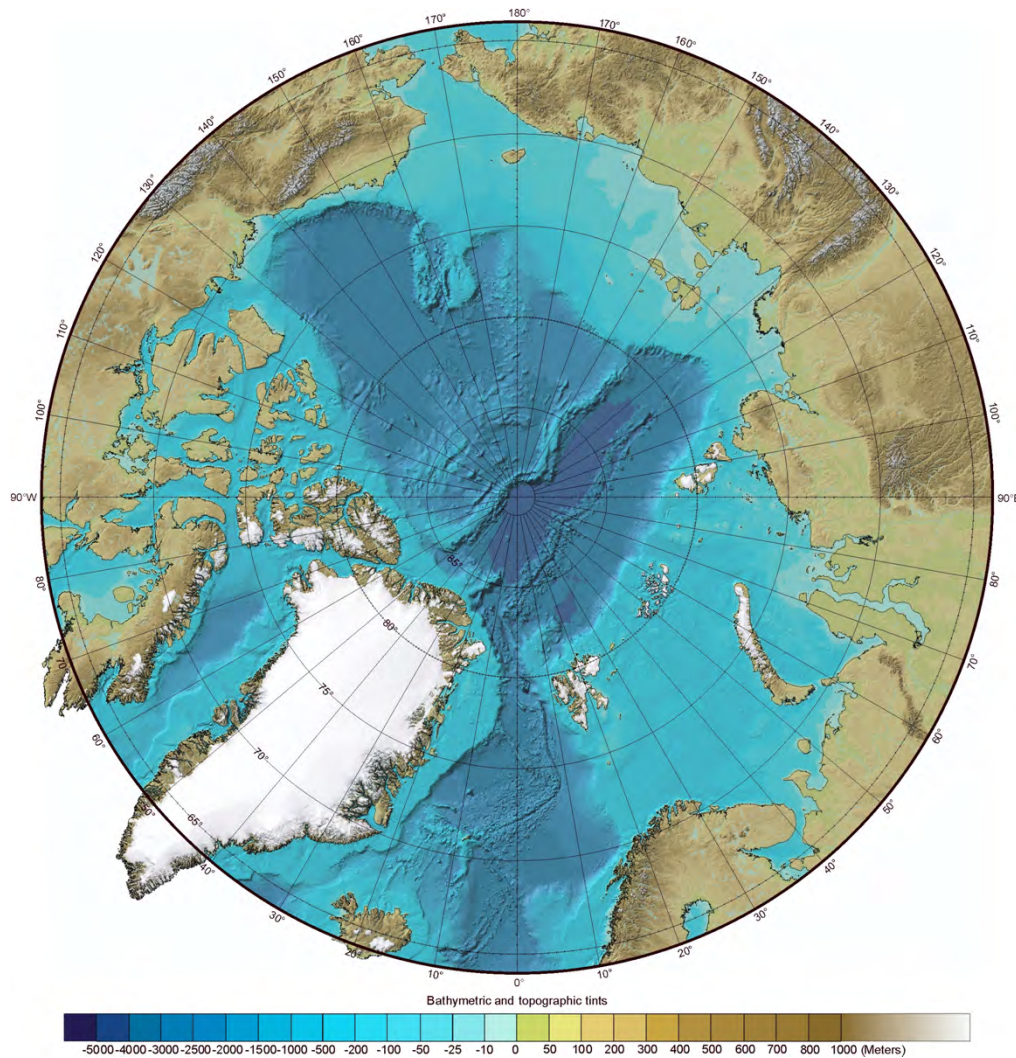
<http://www.cnes.fr>

Bathymetry



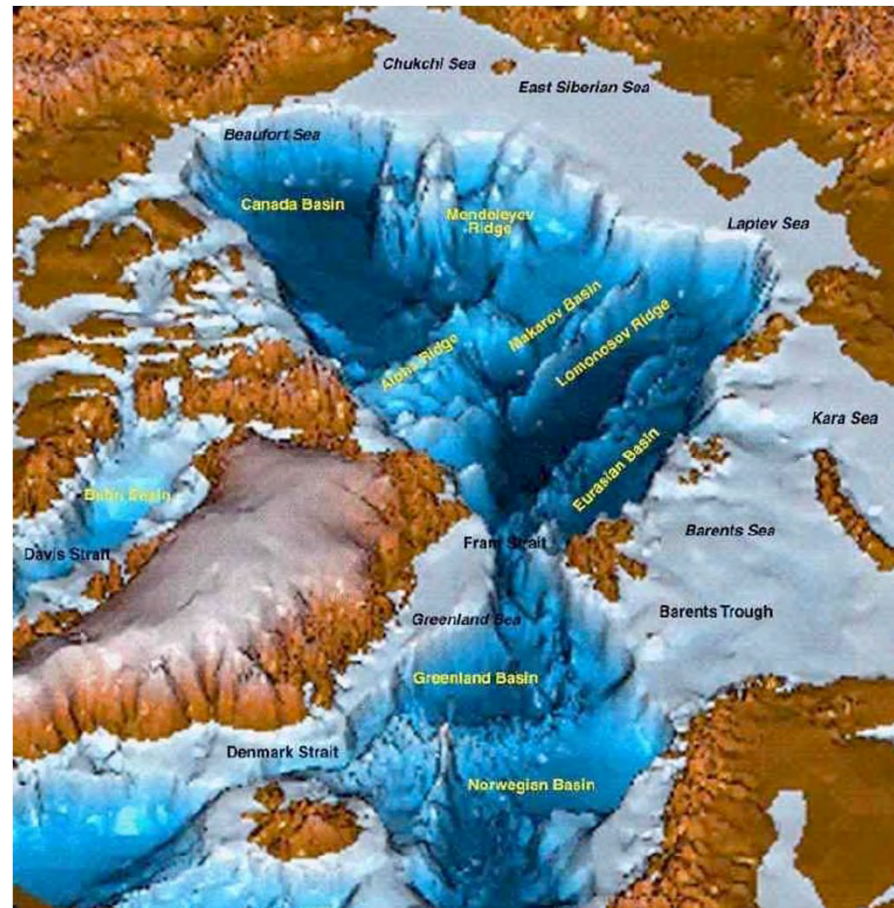
<http://www.ngdc.noaa.gov/mgg/bathymetry/apb/>

Bathymetry

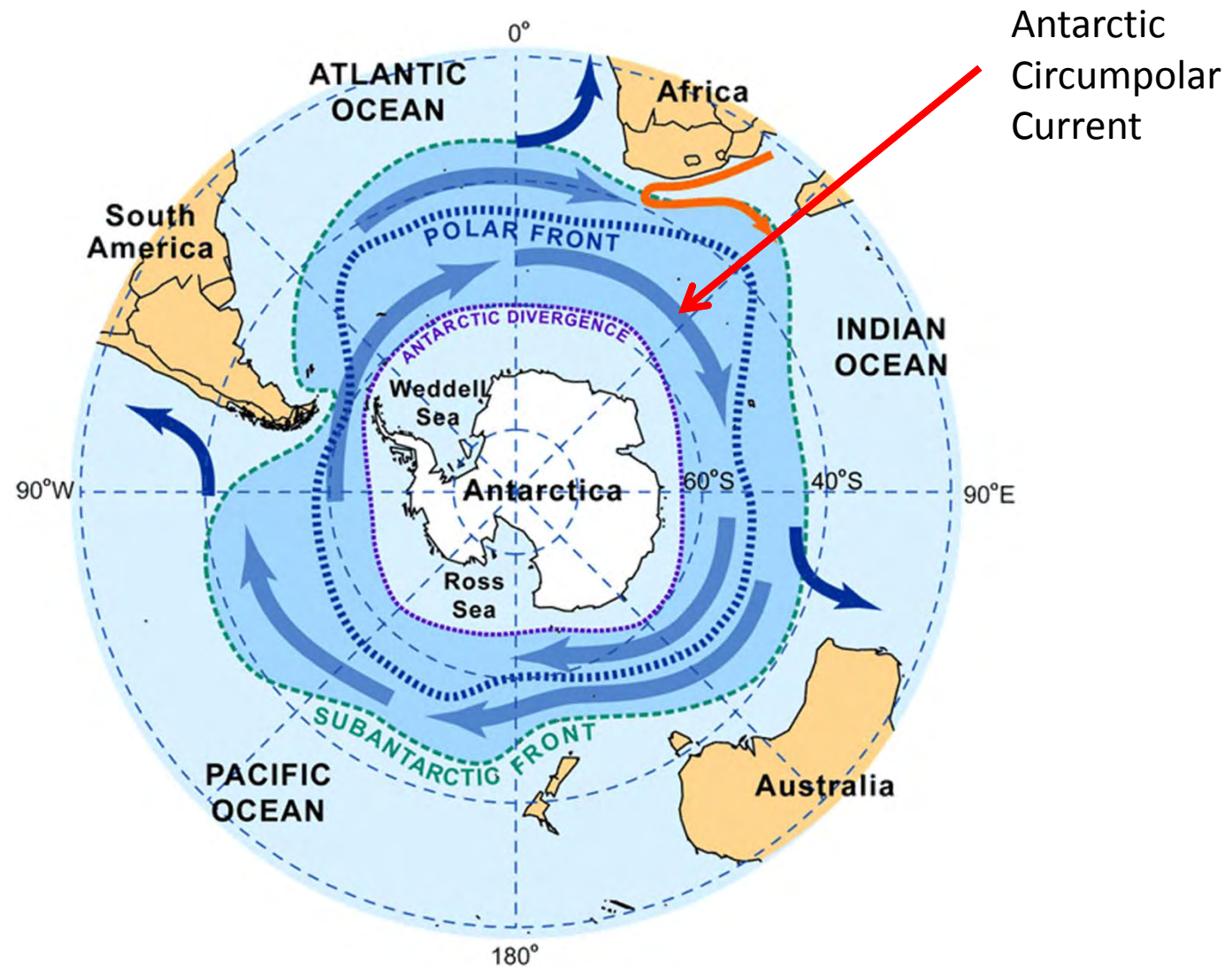


Shelves occupy more than 50% of the Arctic Ocean area

Bathymetry



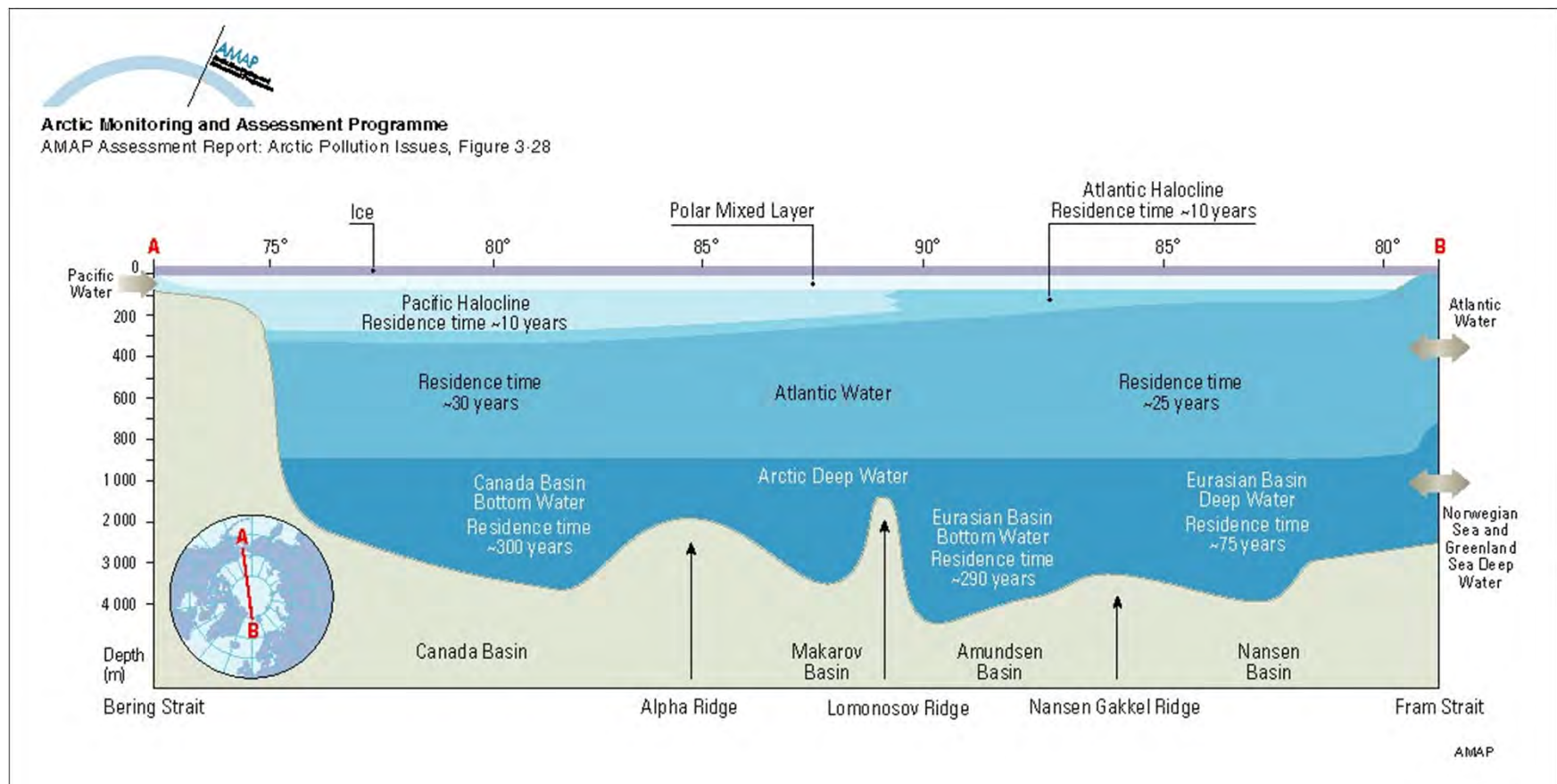
Circulation



Circulation



Water masses



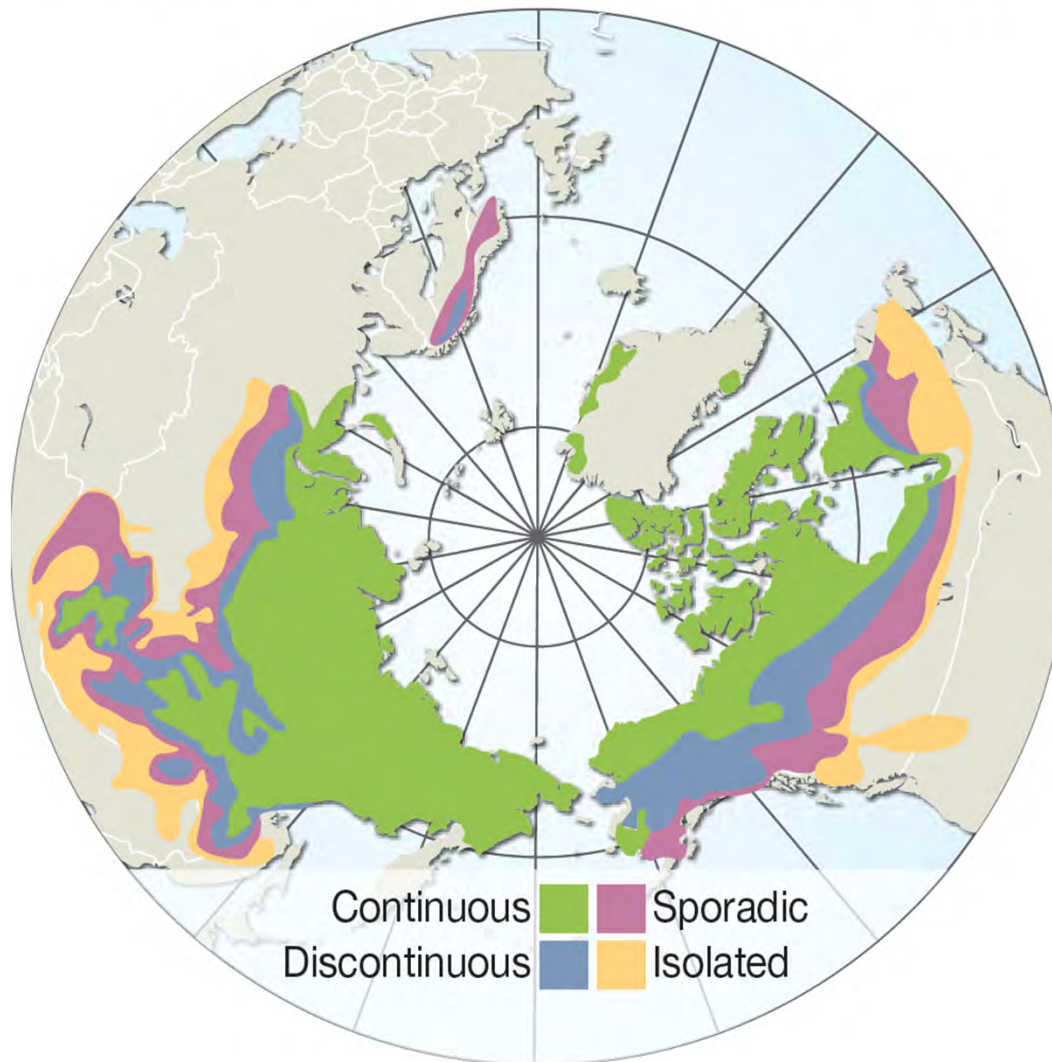
Freshwater discharges



Arctic Ocean = 1% of Global Ocean

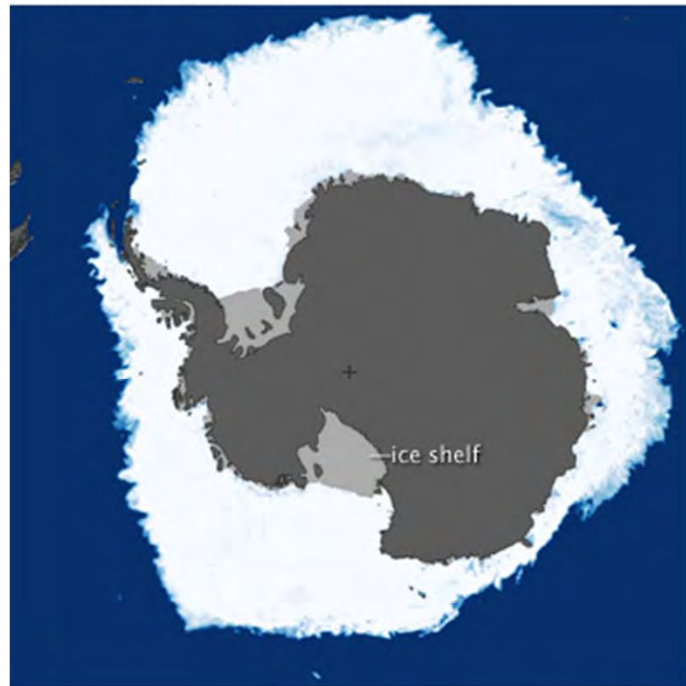
Arctic Ocean receives 10% of global freshwater discharge

Observed Distribution of permafrost types

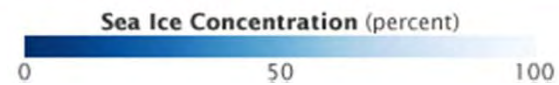


Sea ice

Antarctic Maximum (September 4, 2008)

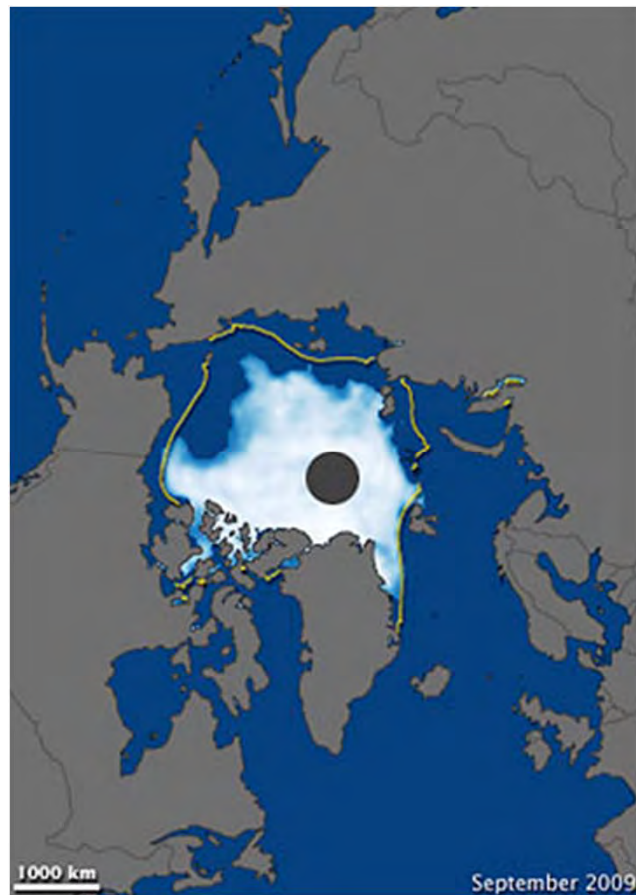


Antarctic Minimum (February 20, 2009)



<http://earthobservatory.nasa.gov>

Sea ice

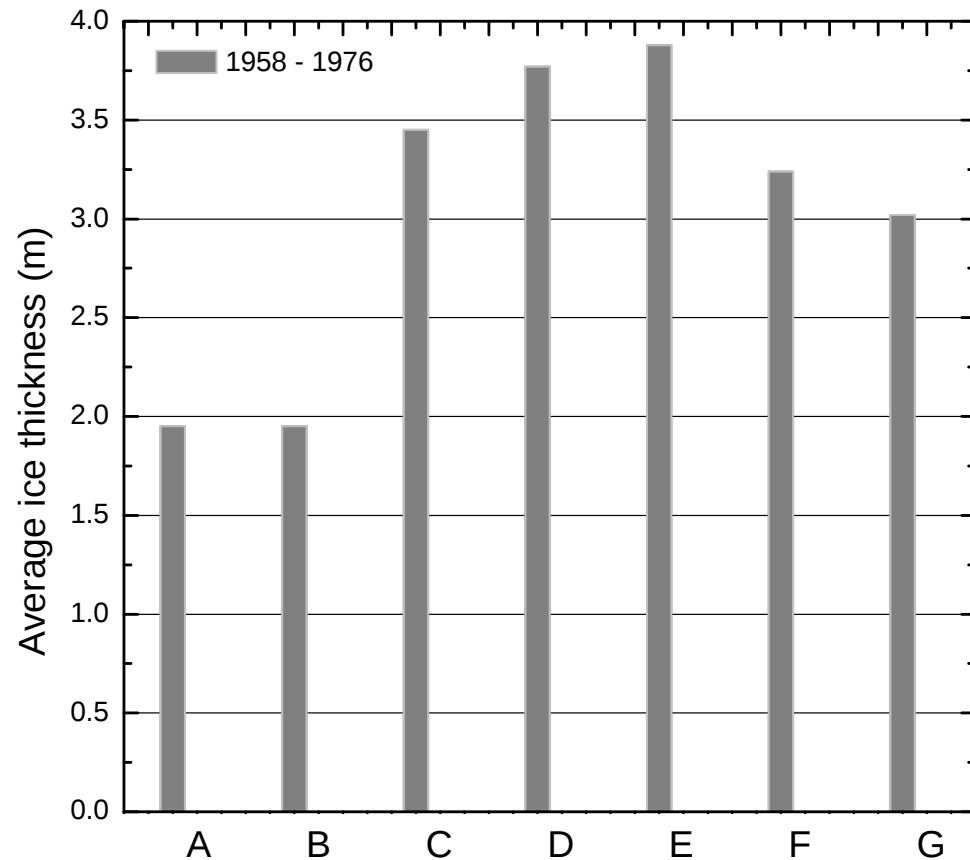


Sea ice thickness: then

Summer ice thickness from 1958-1976

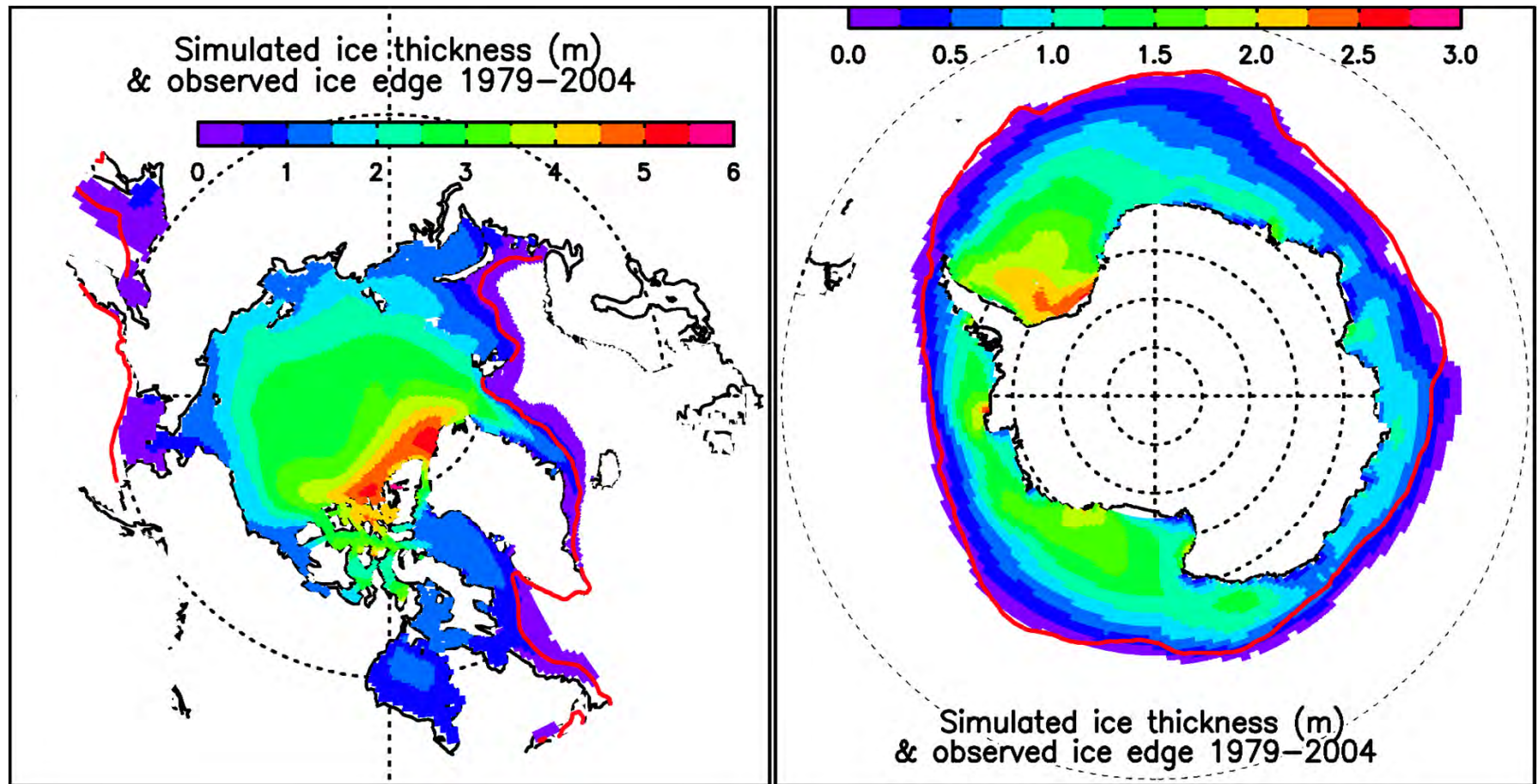


- Analysis of submarine data
- Rothrock et al.

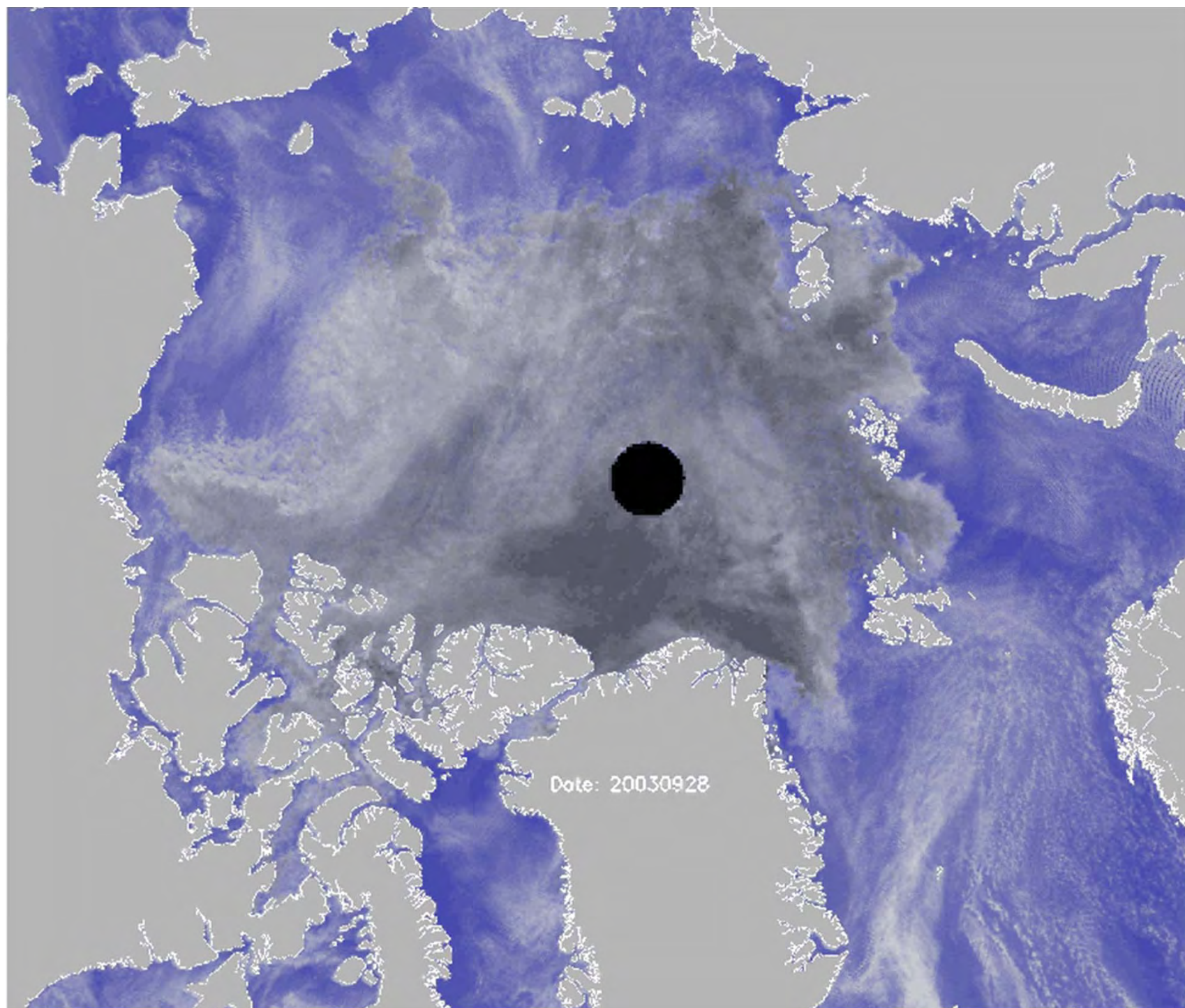


Thick ice everywhere ... even in summer

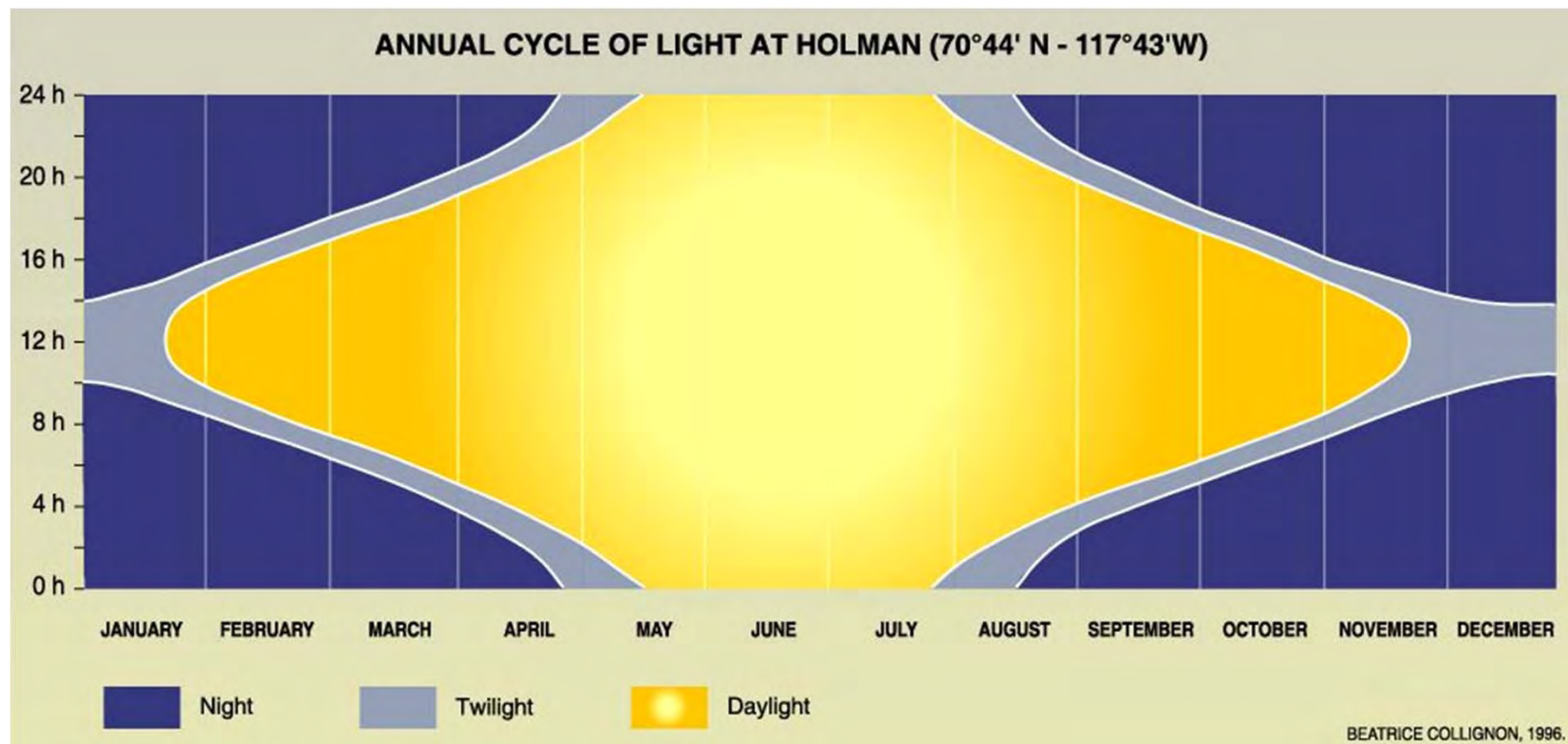
Courtesy of Don Perovich



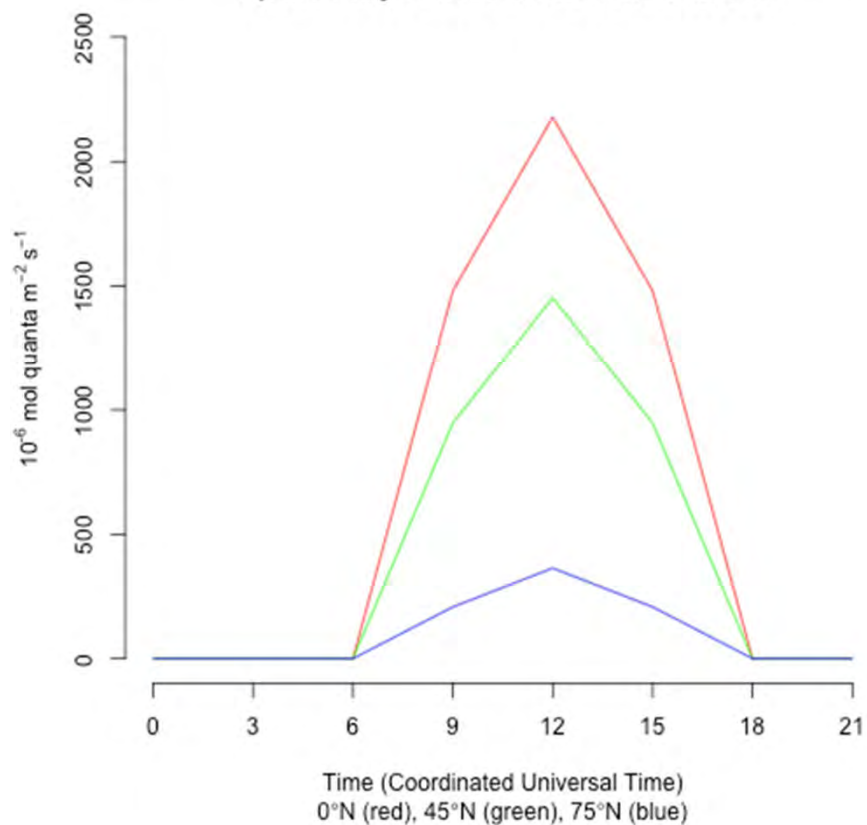
http://psc.apl.washington.edu/zhang/Global_seaice



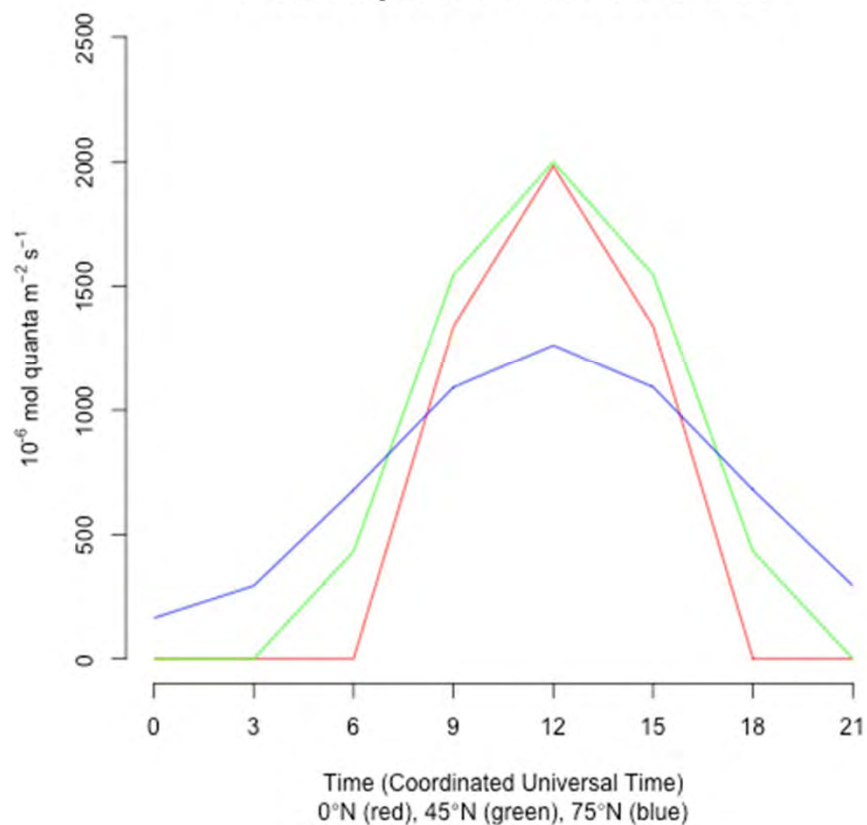
Light cycle

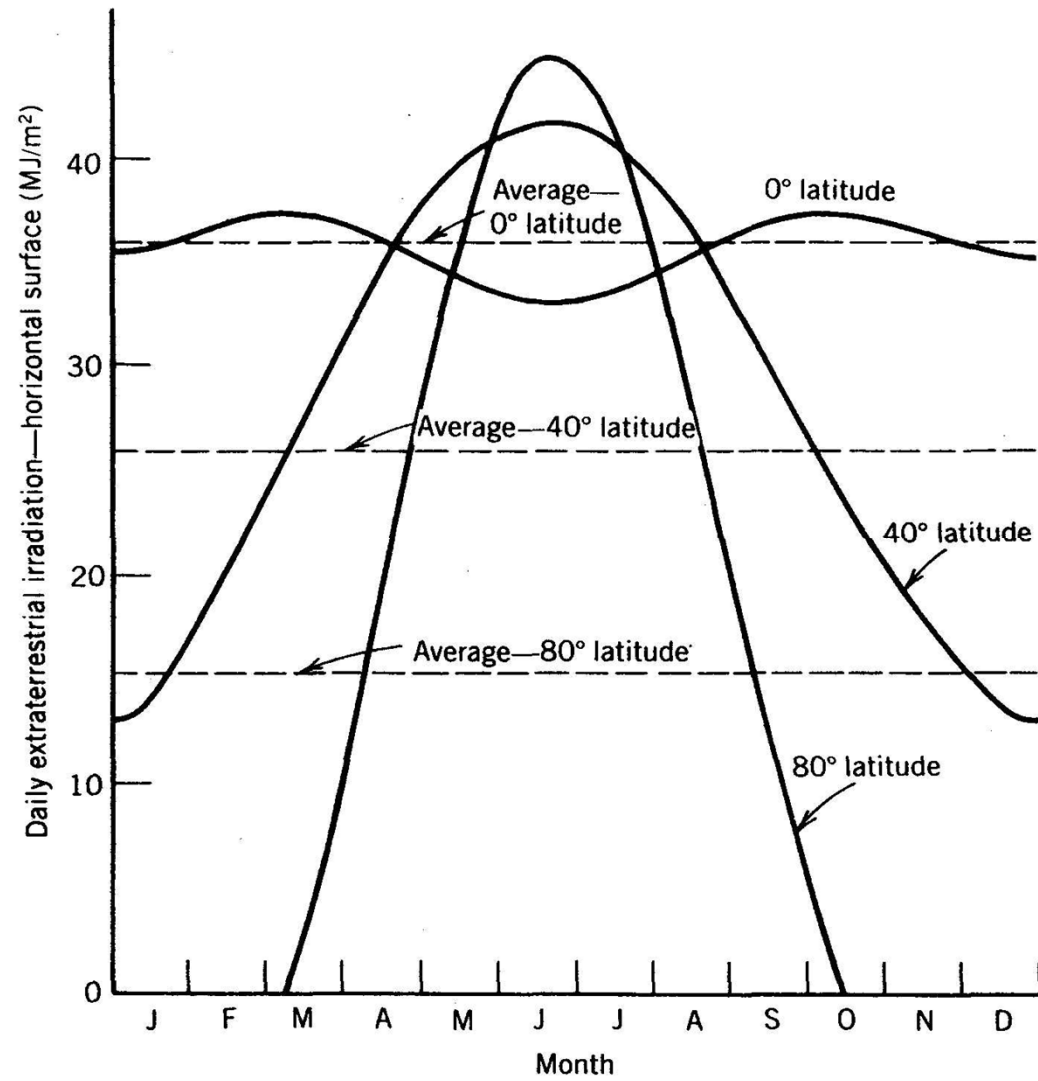


Diurnal variation of downward irradiance of PAR just below the surface water if that day was cloud-free at 0E on the spring equinox day of 2009 at different latitudes

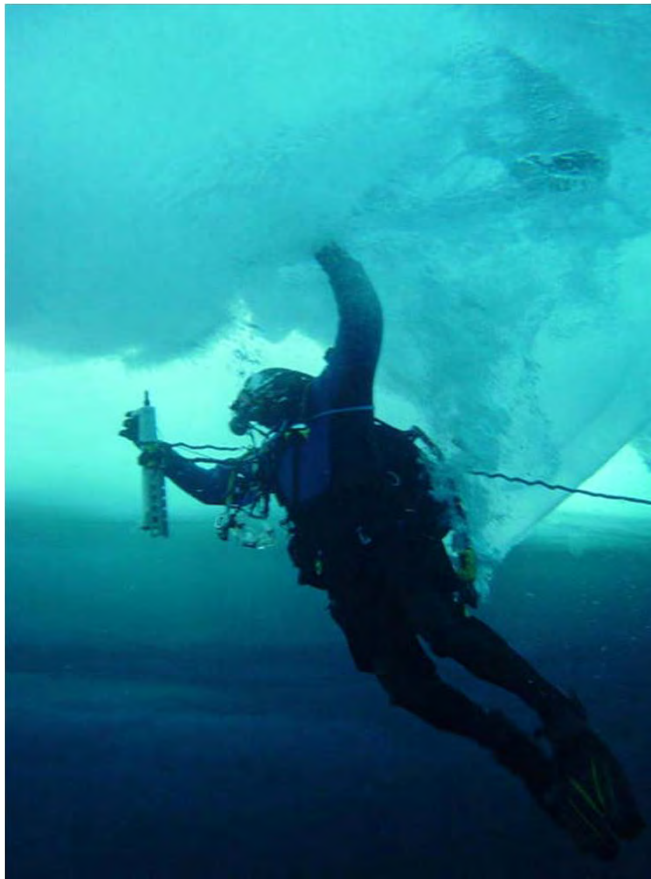


Diurnal variation of downward irradiance of PAR just below the surface water if that day was cloud-free at 0E on the summer solstice day of 2009 at different latitudes

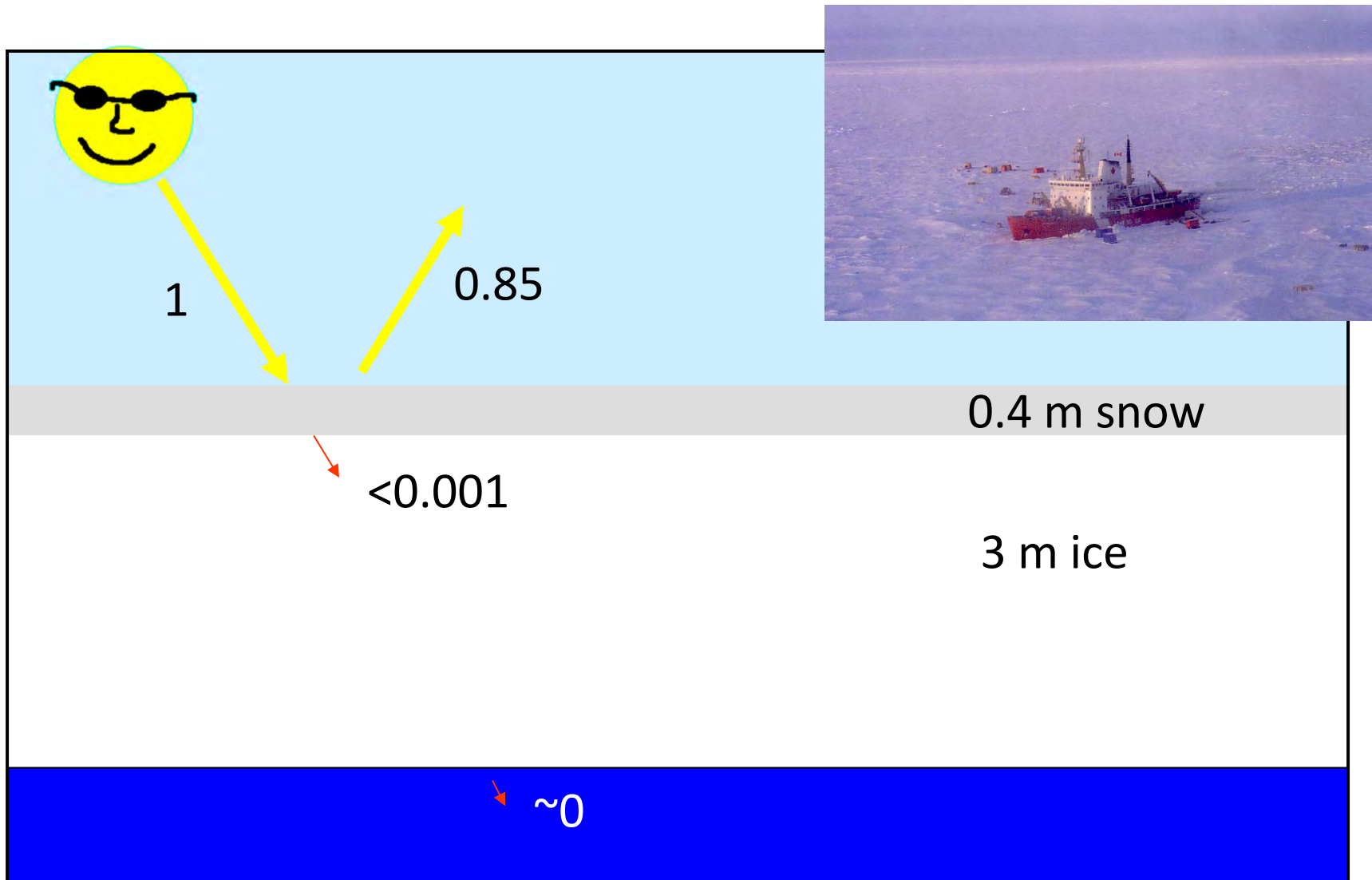




Light under sea ice



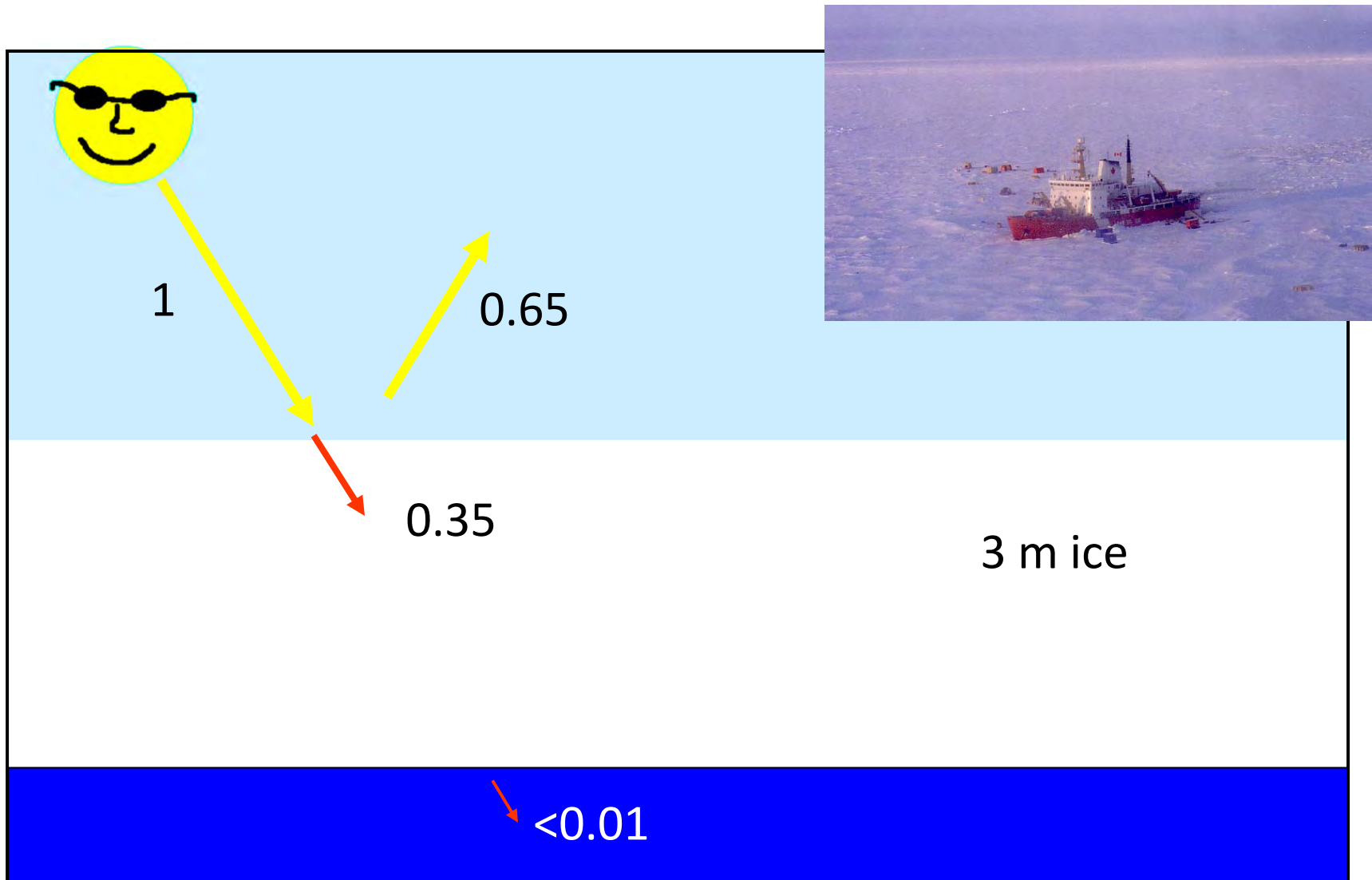
Back then: Assume a simple slab



Snow covered, thick, multiyear ice

Courtesy of Don Perovich

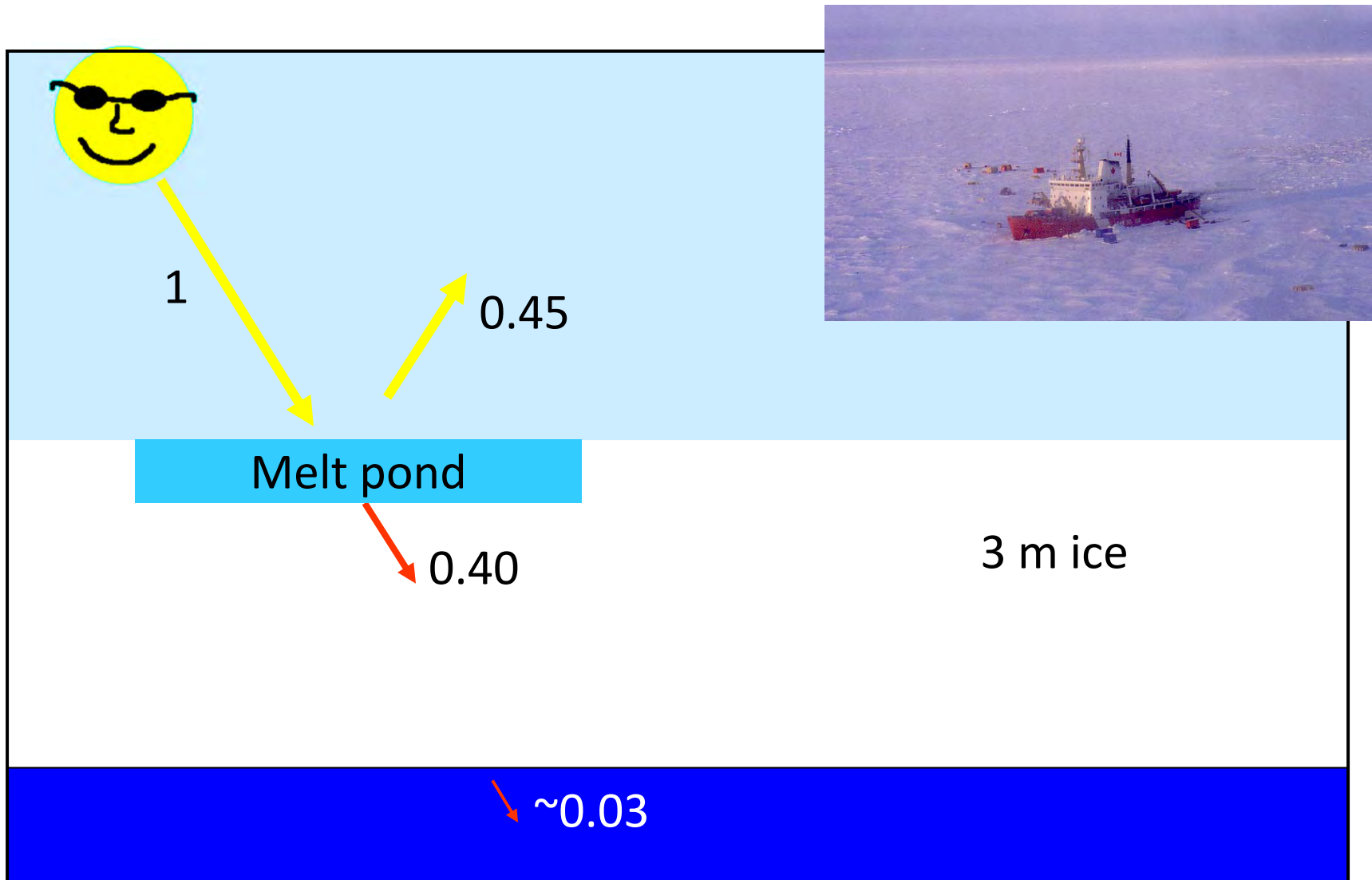
Back then: Assume a simple slab



Let the snow melt

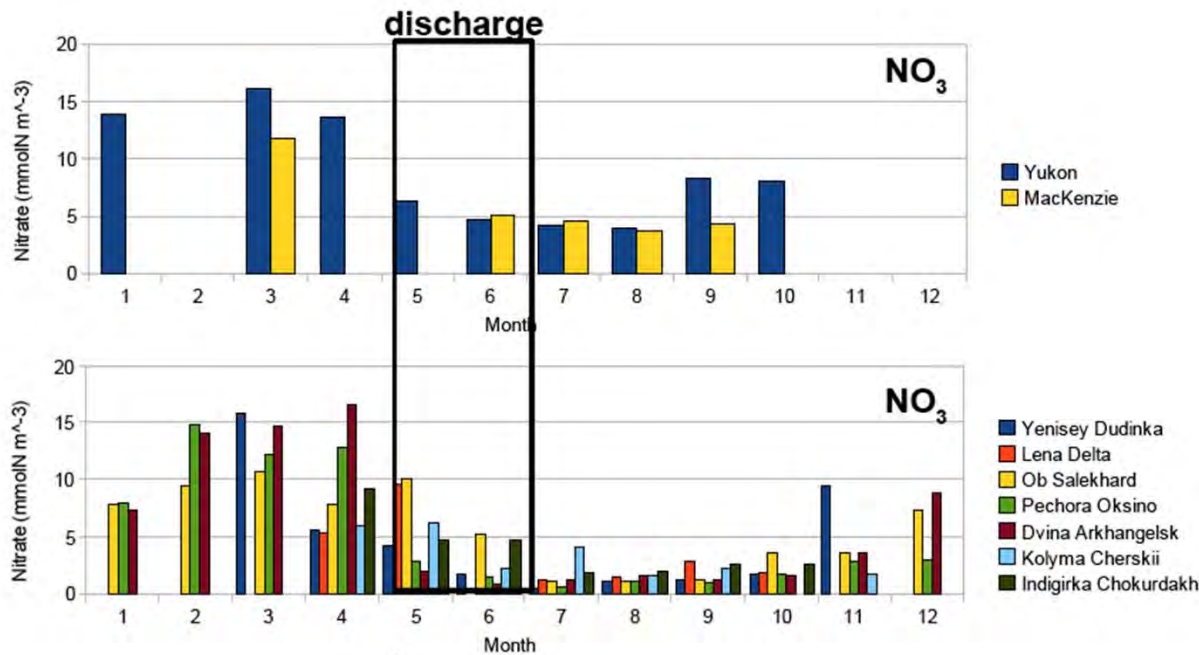
Courtesy of Don Perovich

Back then: Assume a simple slab



Let the snow melt and add a pond

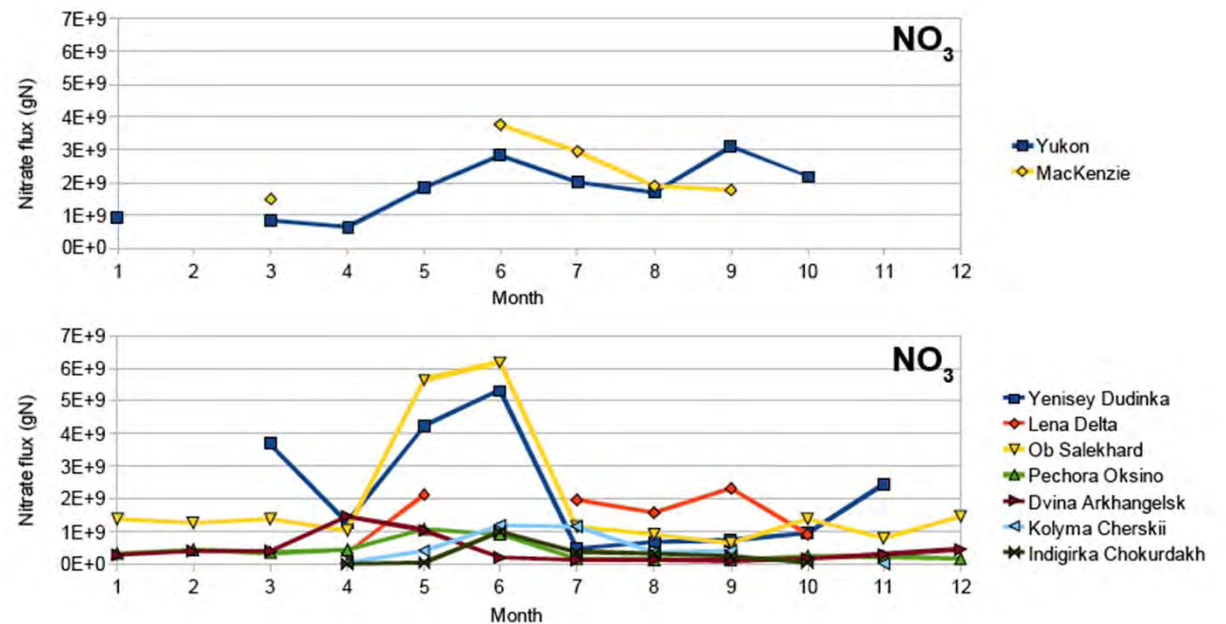
Courtesy of Don Perovich



A monthly climatology of nitrate, phosphate, silicate, DON et DOC concentration was built-up:

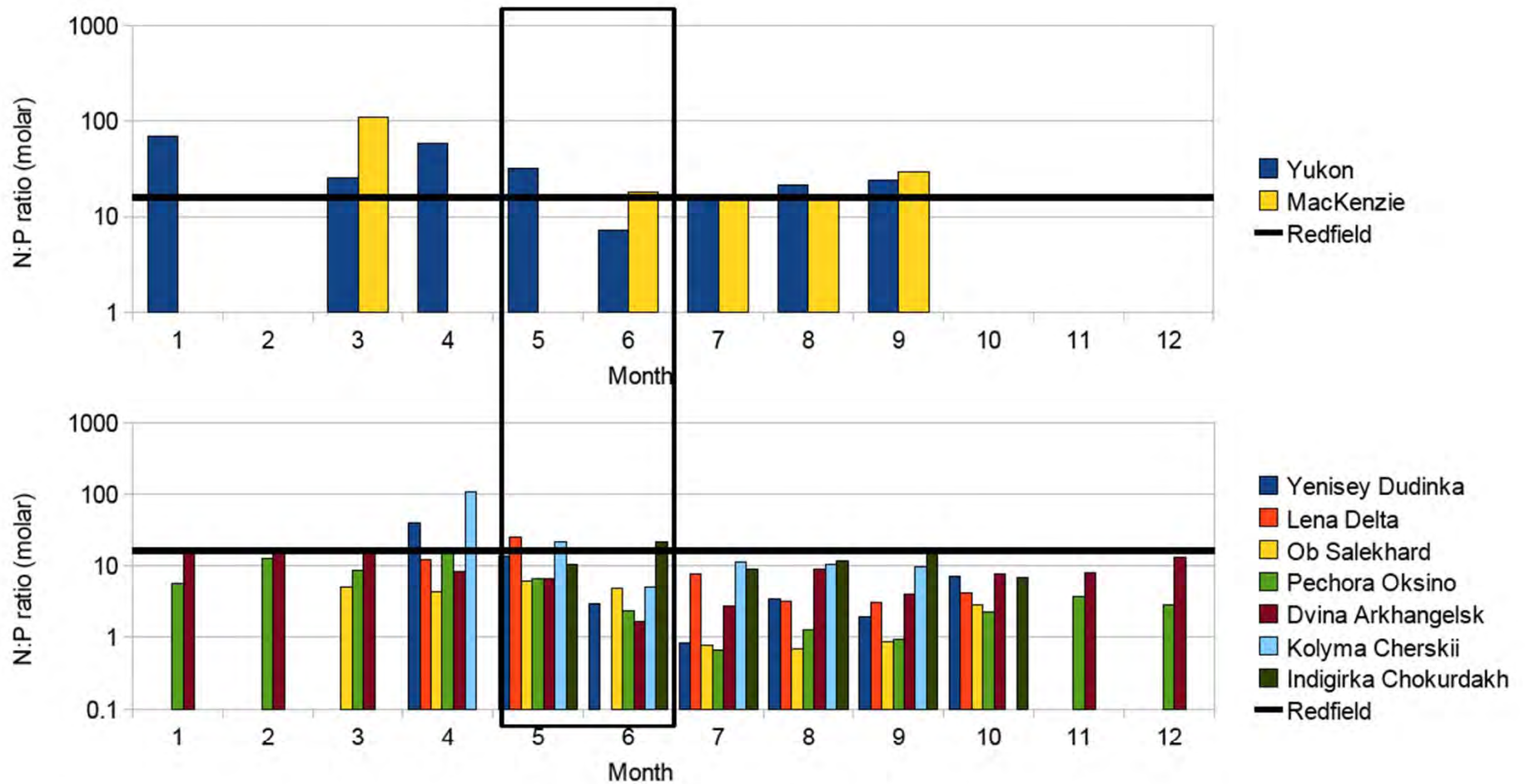
- 9 publications et 2 databases
- 9 most important pan-arctic rivers

concentration
×
discharge



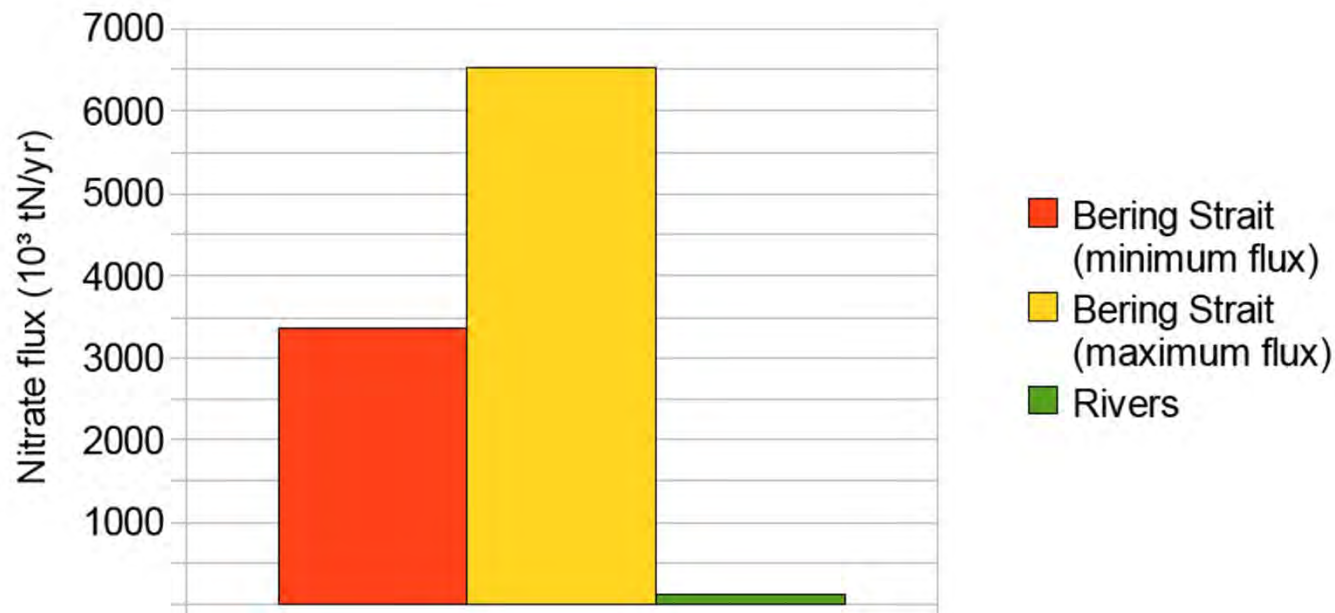
Lefouest et al.

What contribution for riverine nitrate on PP?



During the peak of river discharge (May-June), riverine nitrate is limiting relative to riverine phosphate in 70% of the cases

What contribution for riverine nitrate on PP?



Assuming :

A total Arctic Ocean PP of ~350 TgC/yr

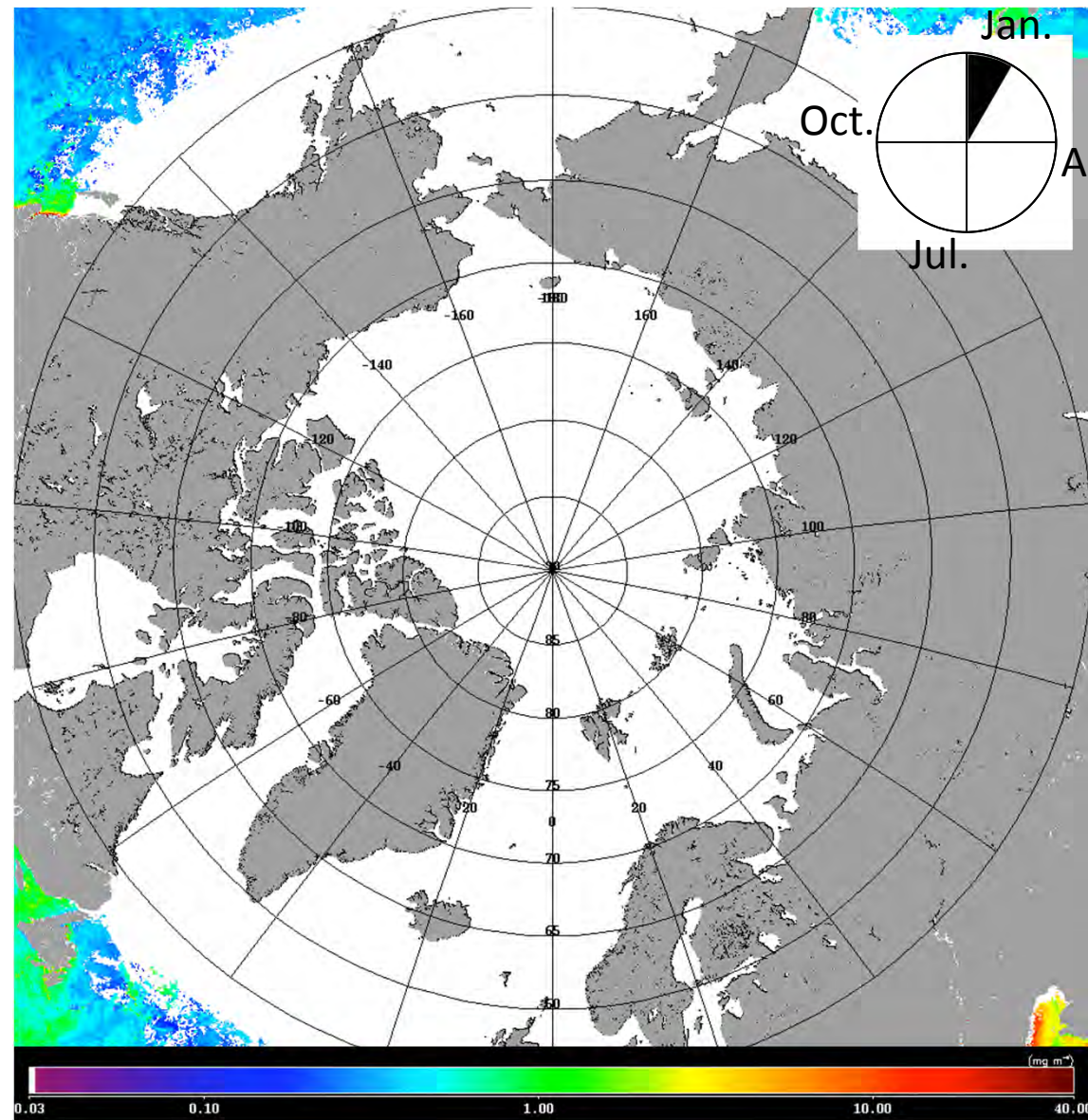
All nitrate are converted into biogenic carbon

then:

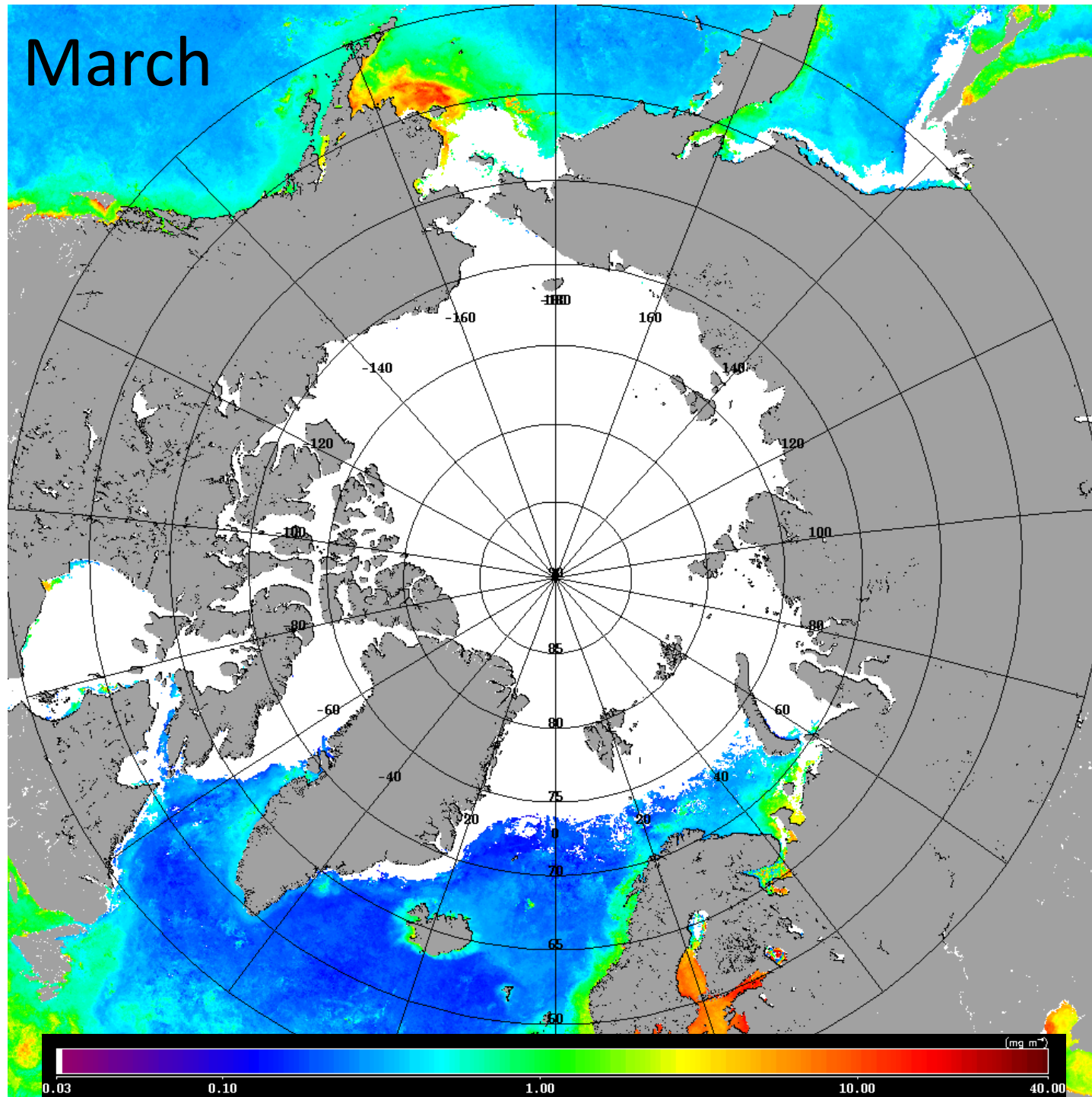
Bering Strait inputs would account for 4-8% of the PP

Riverine inputs would account for 0.14% of the PP

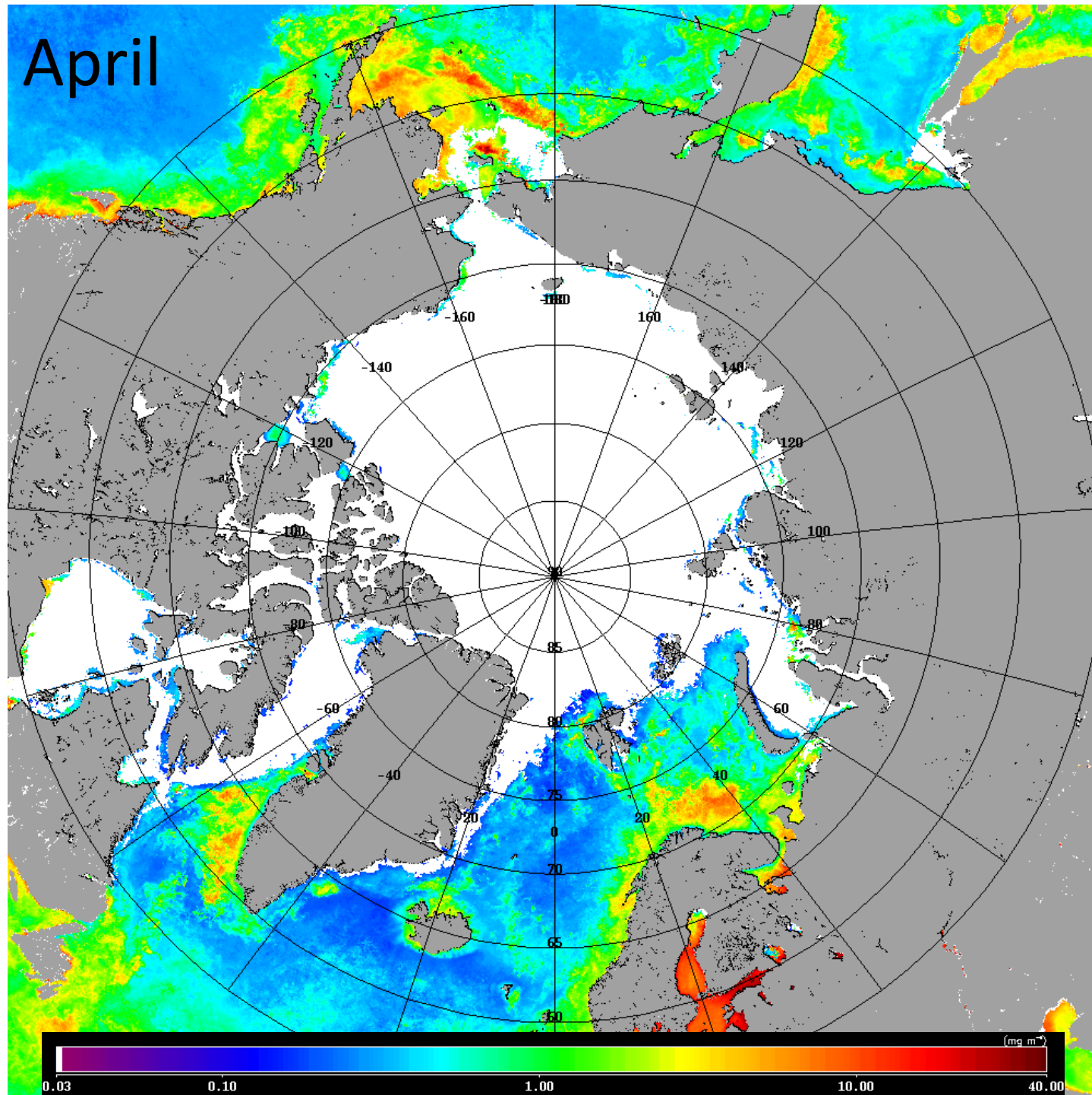
MODIS monthly climatology (2003 à 2011) of chlorophyll concentration

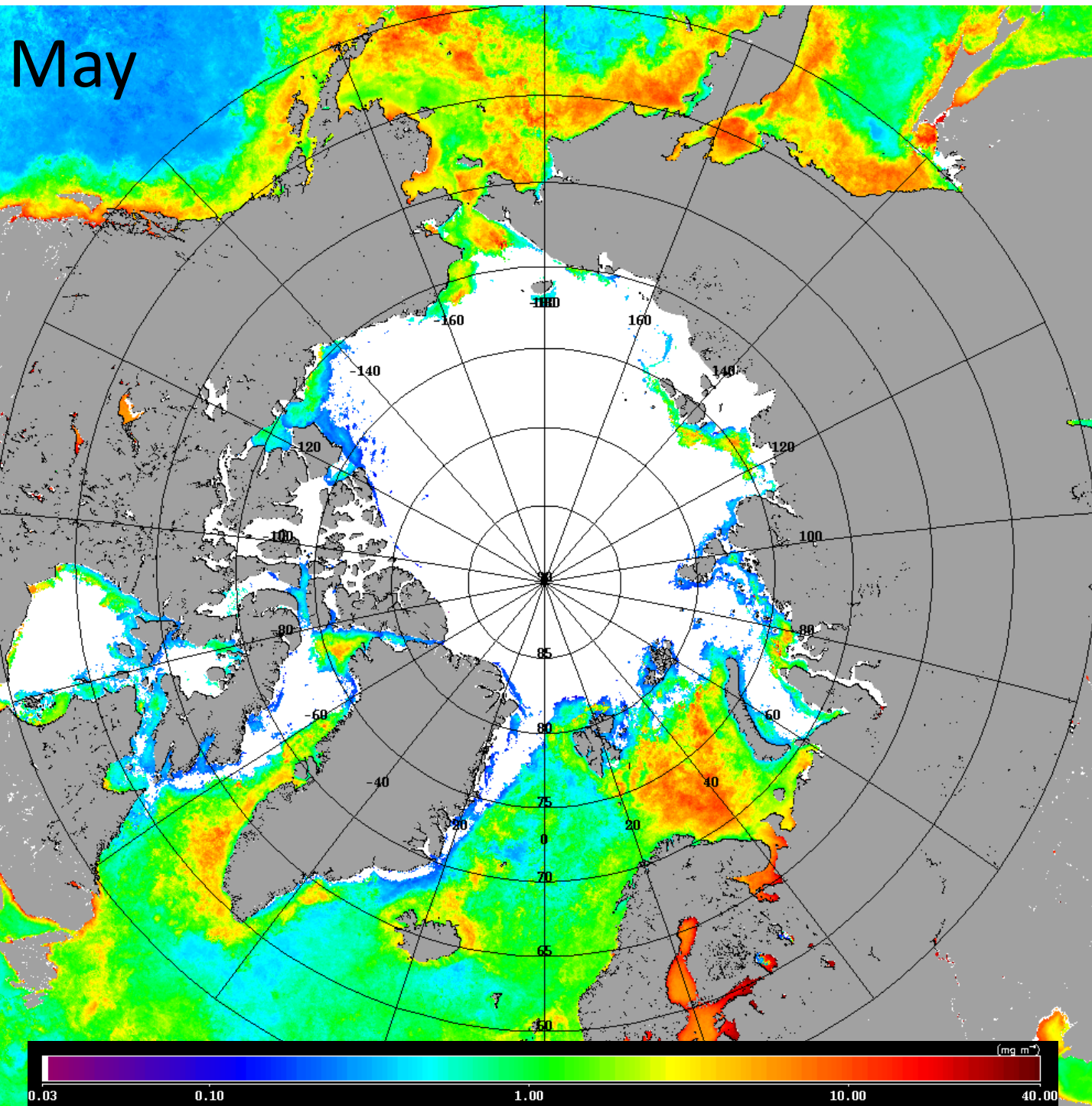


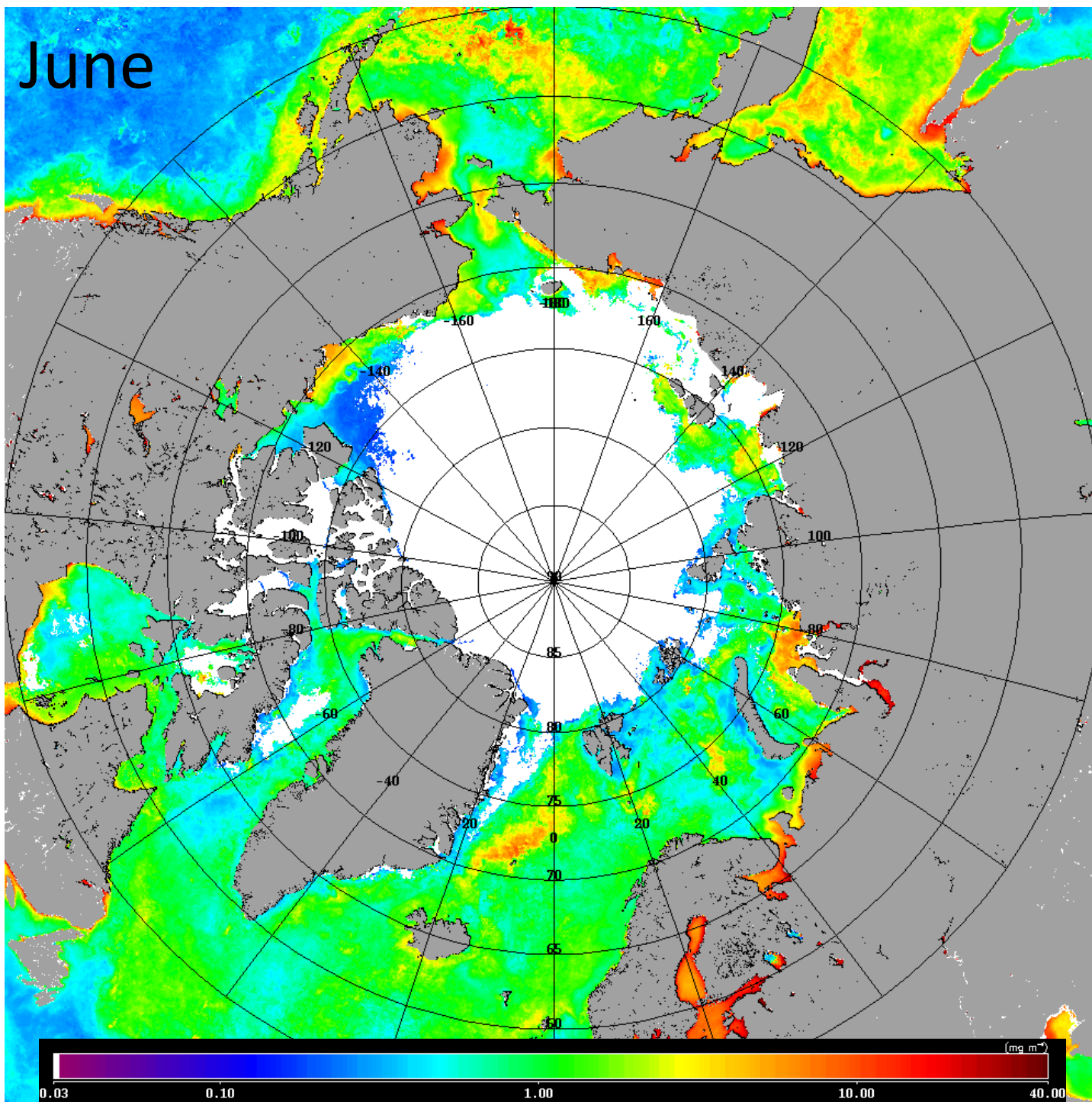
March

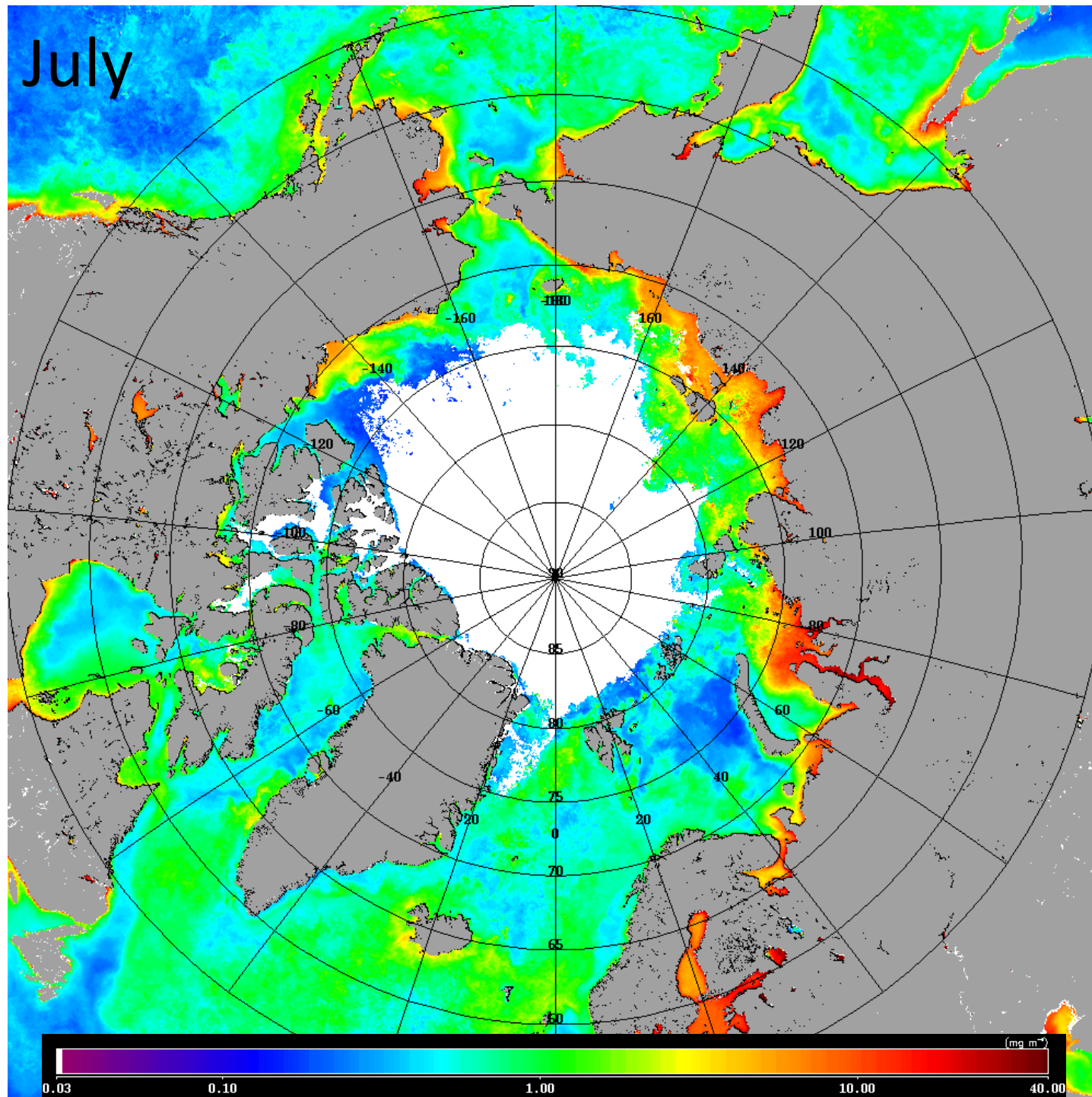


April

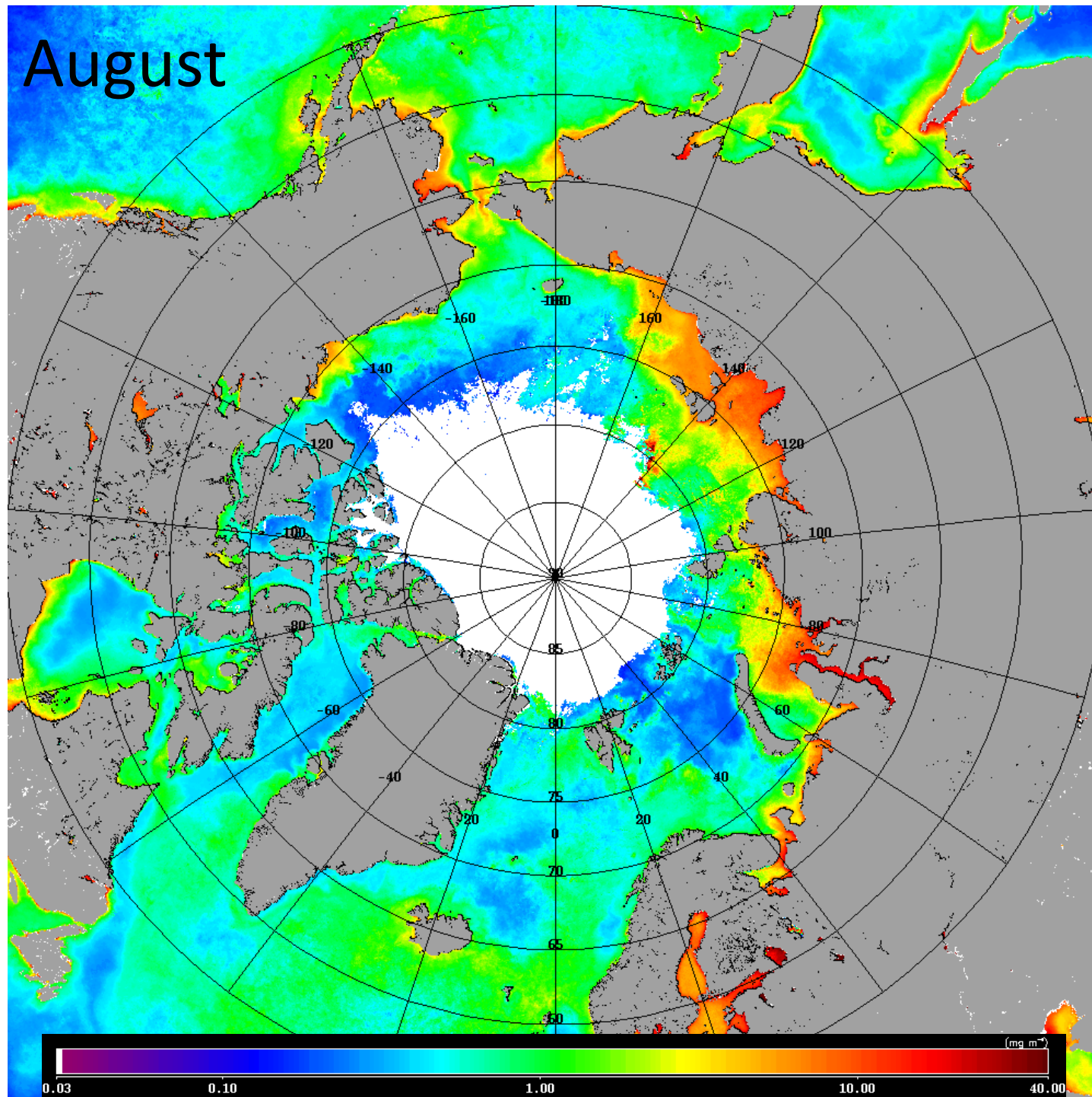


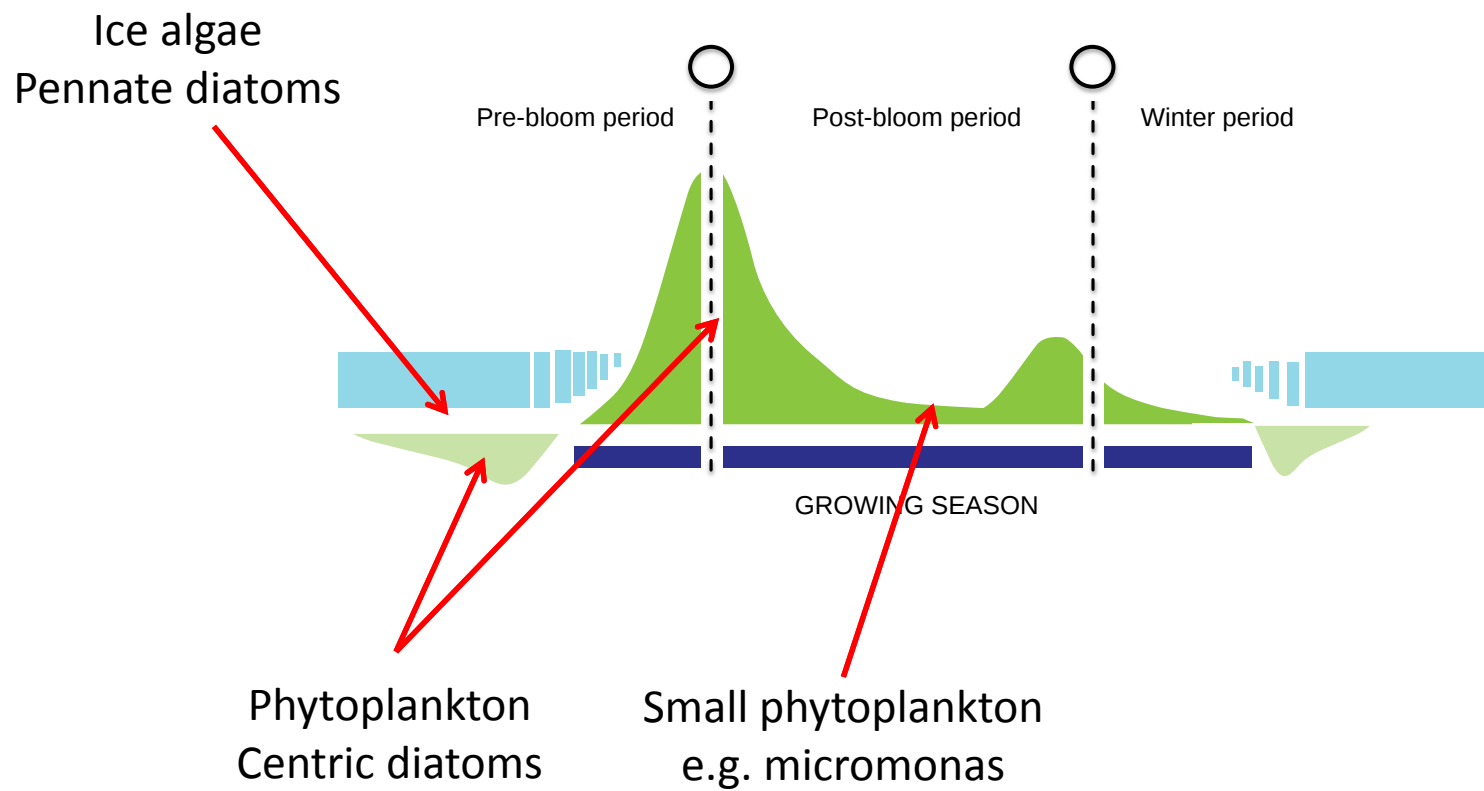




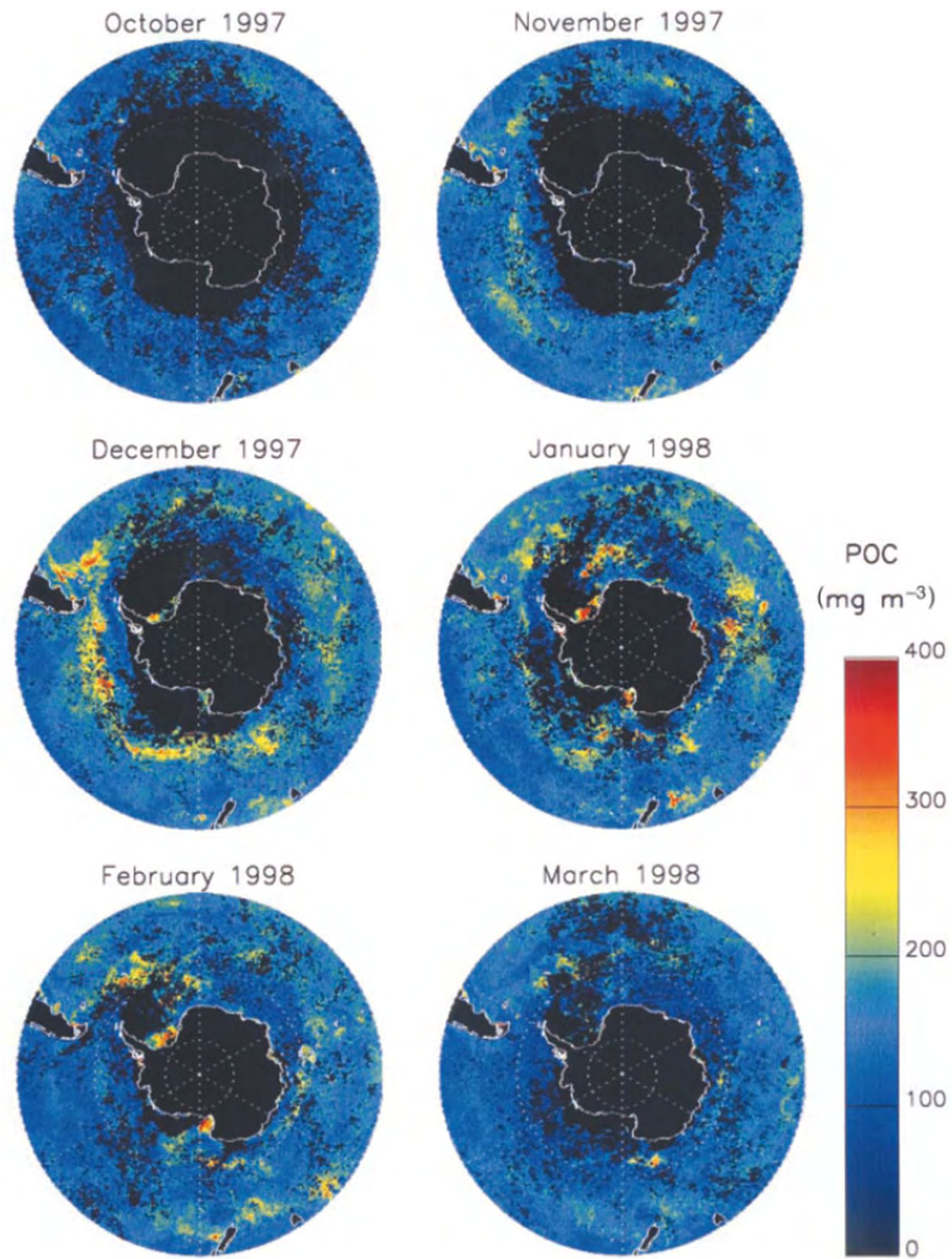


August



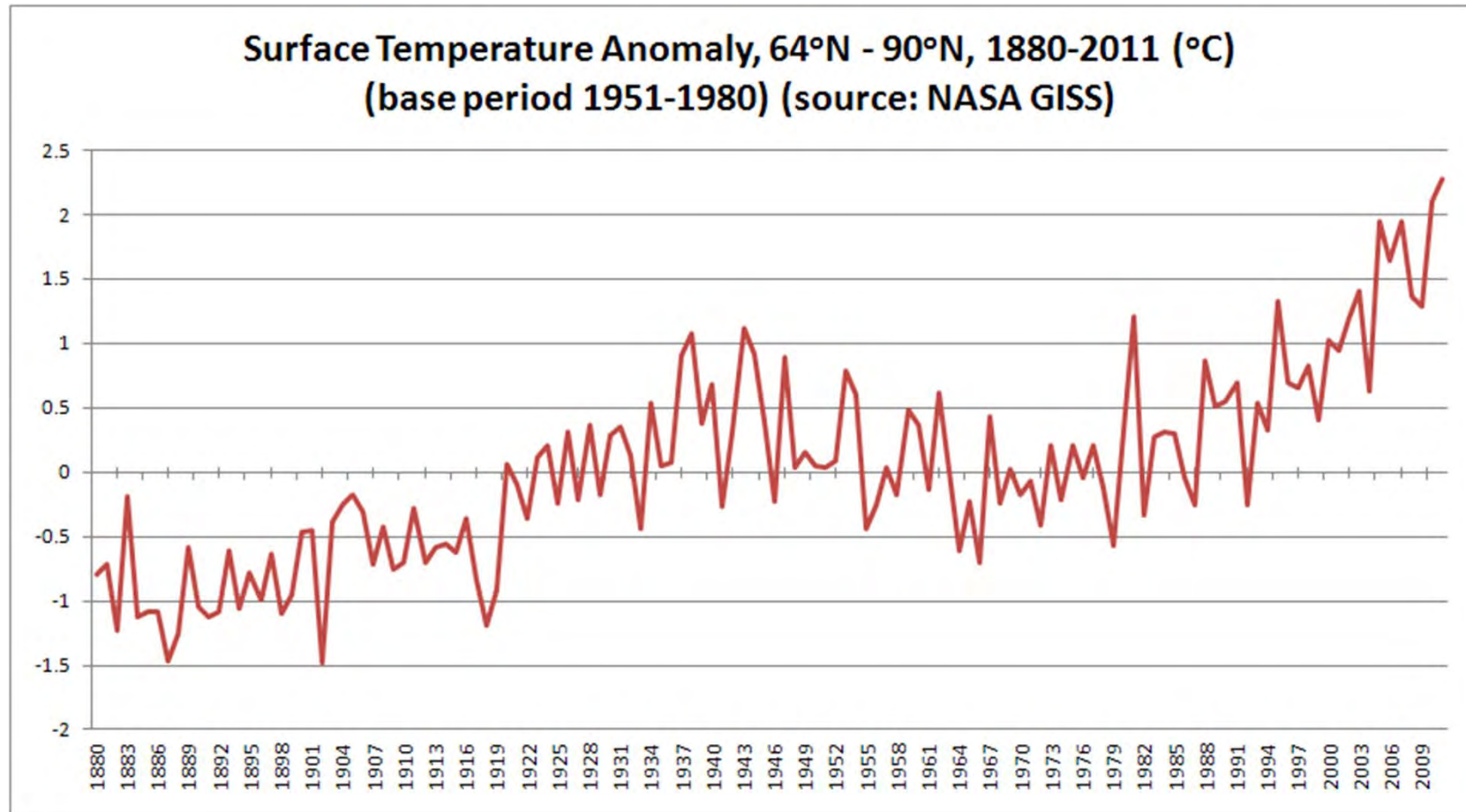


Courtesy of Mathieu Ardyna



Stramski et al. (1999)

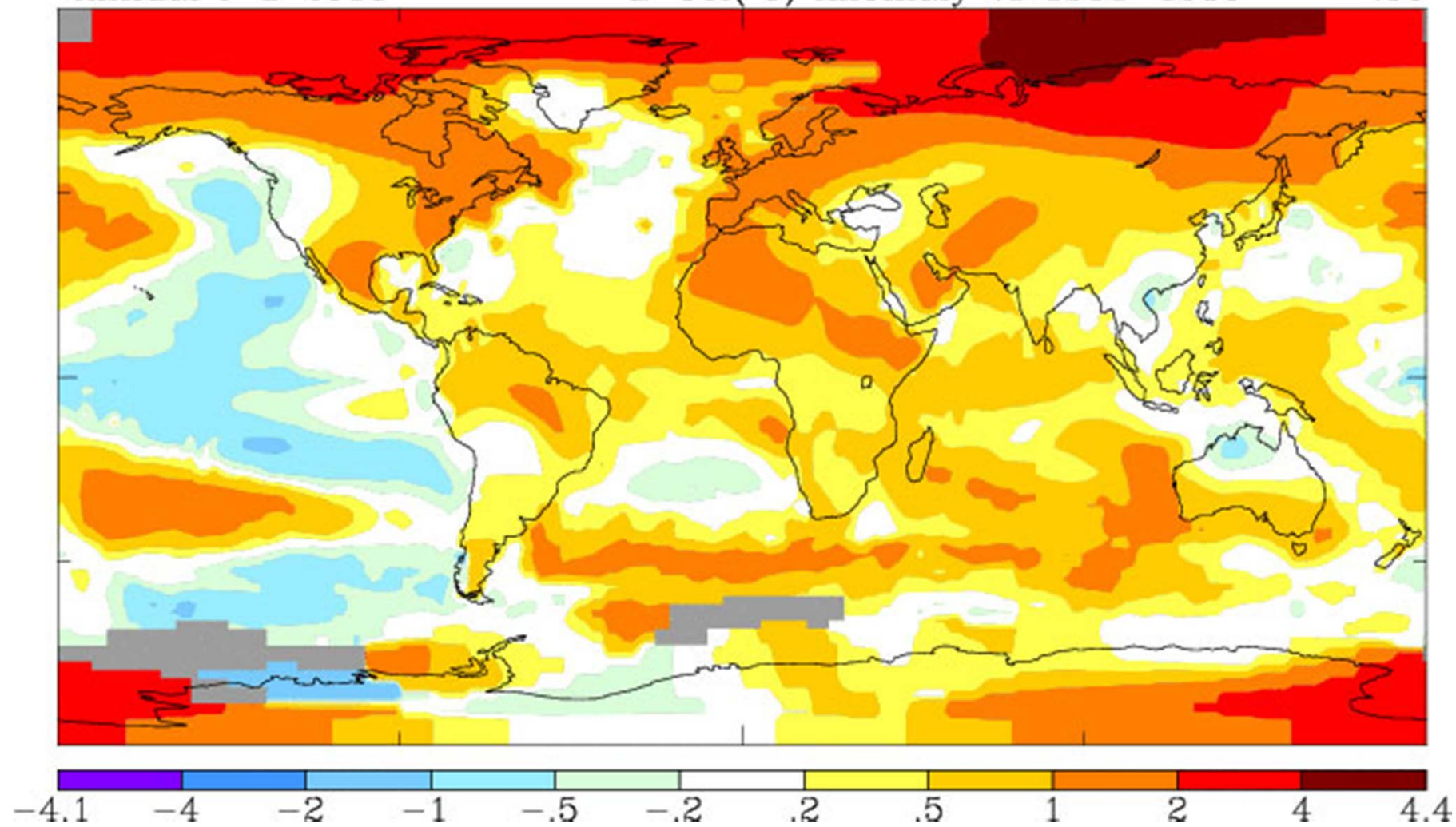
Ongoing changes



Annual J-D 2011

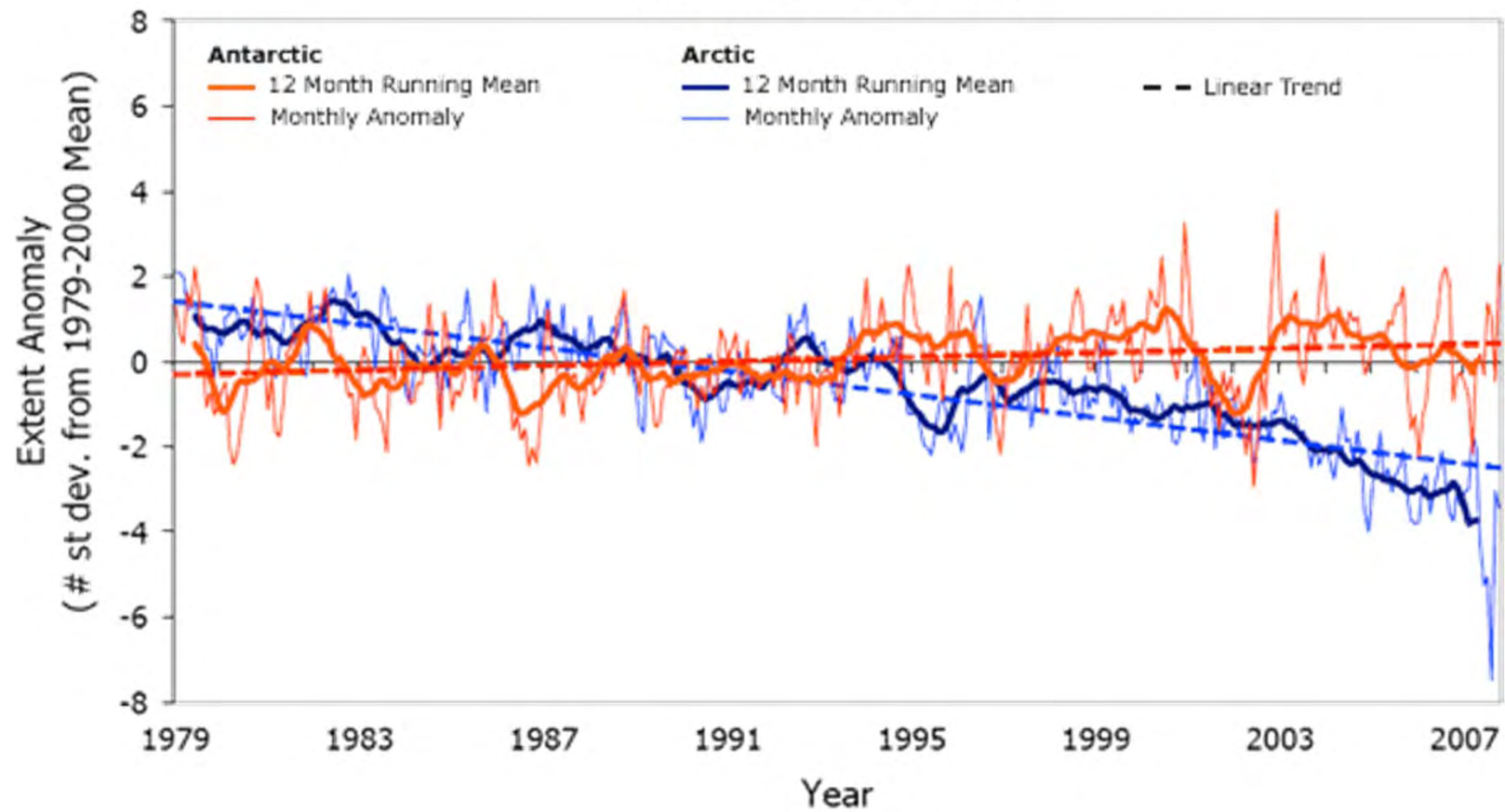
L-OTI(°C) Anomaly vs 1951-1980

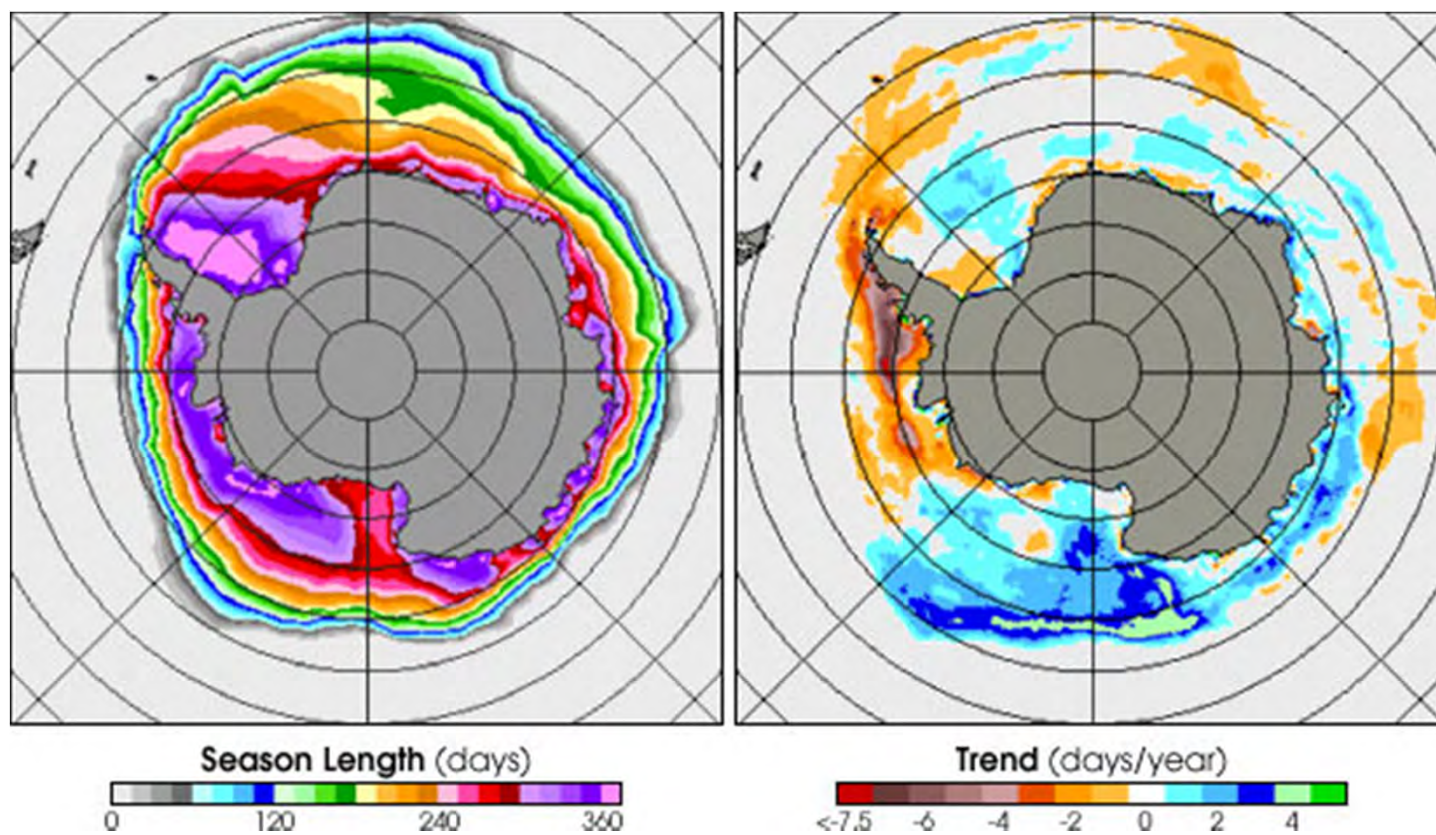
.50



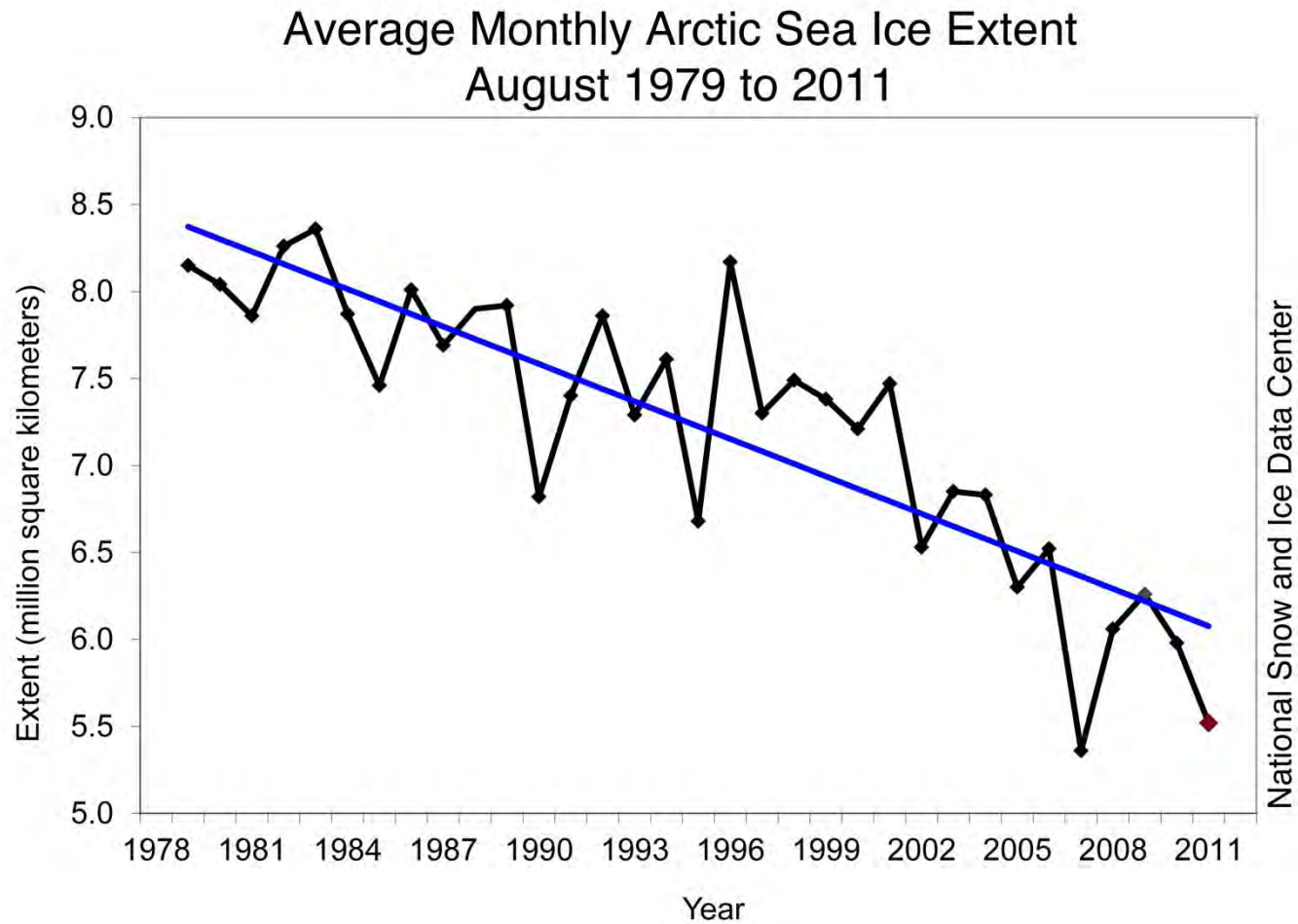
Ongoing changes

Arctic and Antarctic Standardized Anomalies and Trends
Jan 1979 - Dec 2007

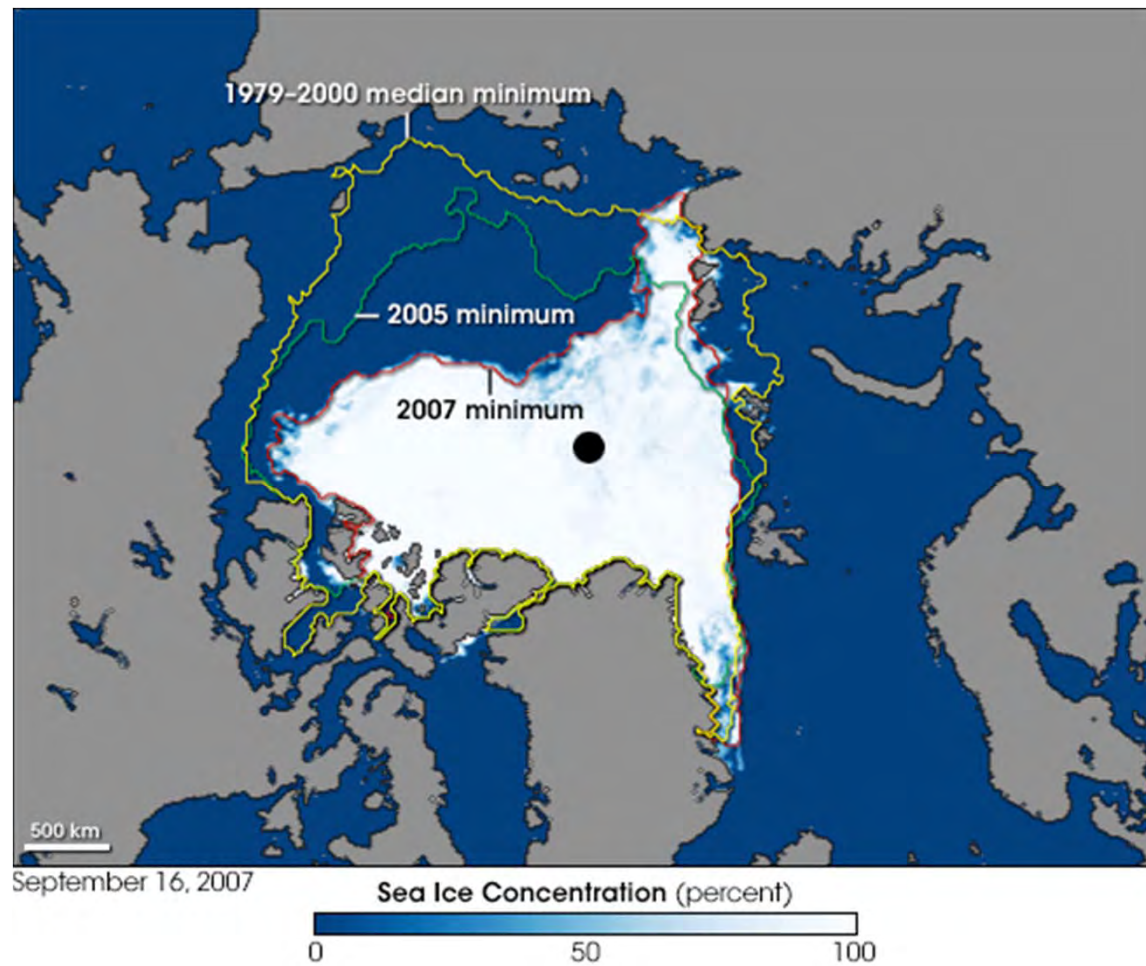




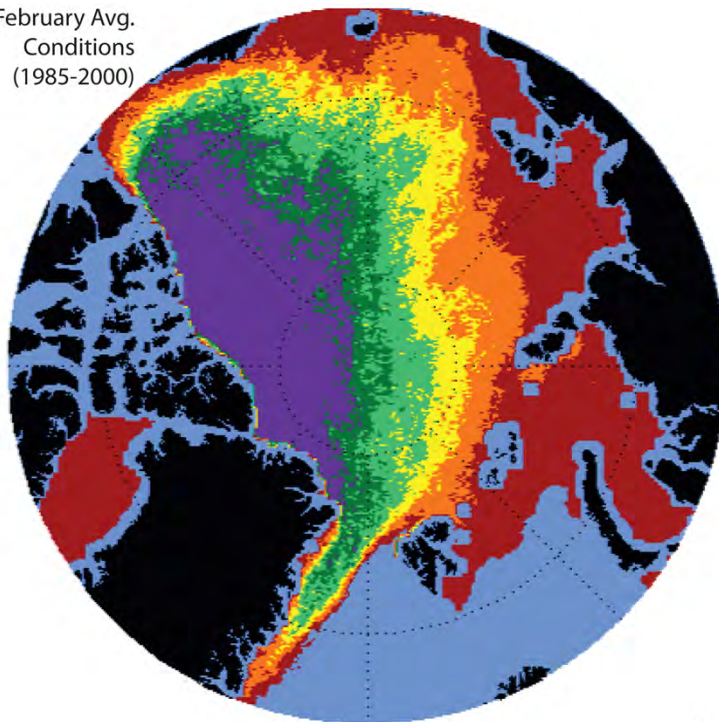
Ongoing changes



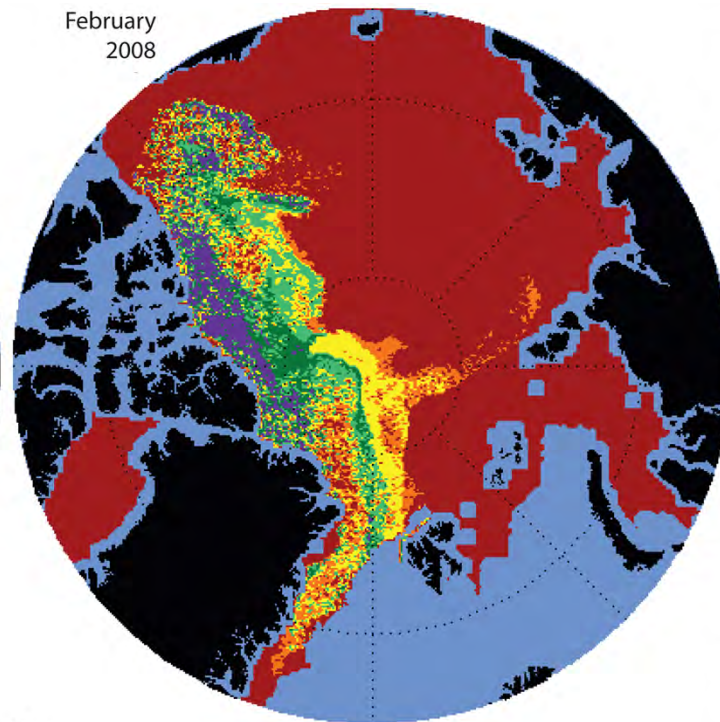
Ongoing changes



February Avg.
Conditions
(1985-2000)



February
2008

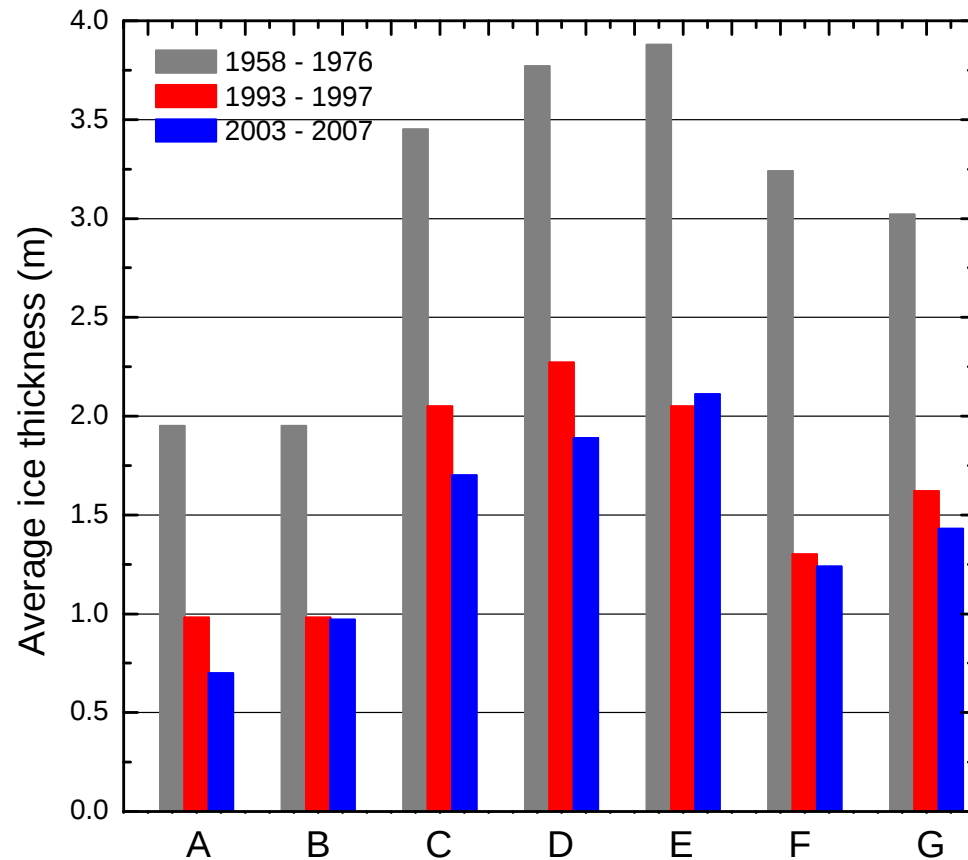


Ice Age



Sea ice thickness: then and now

*Changes in summer thickness
Comparing 1958-1976 to the 1990's and 2000's*

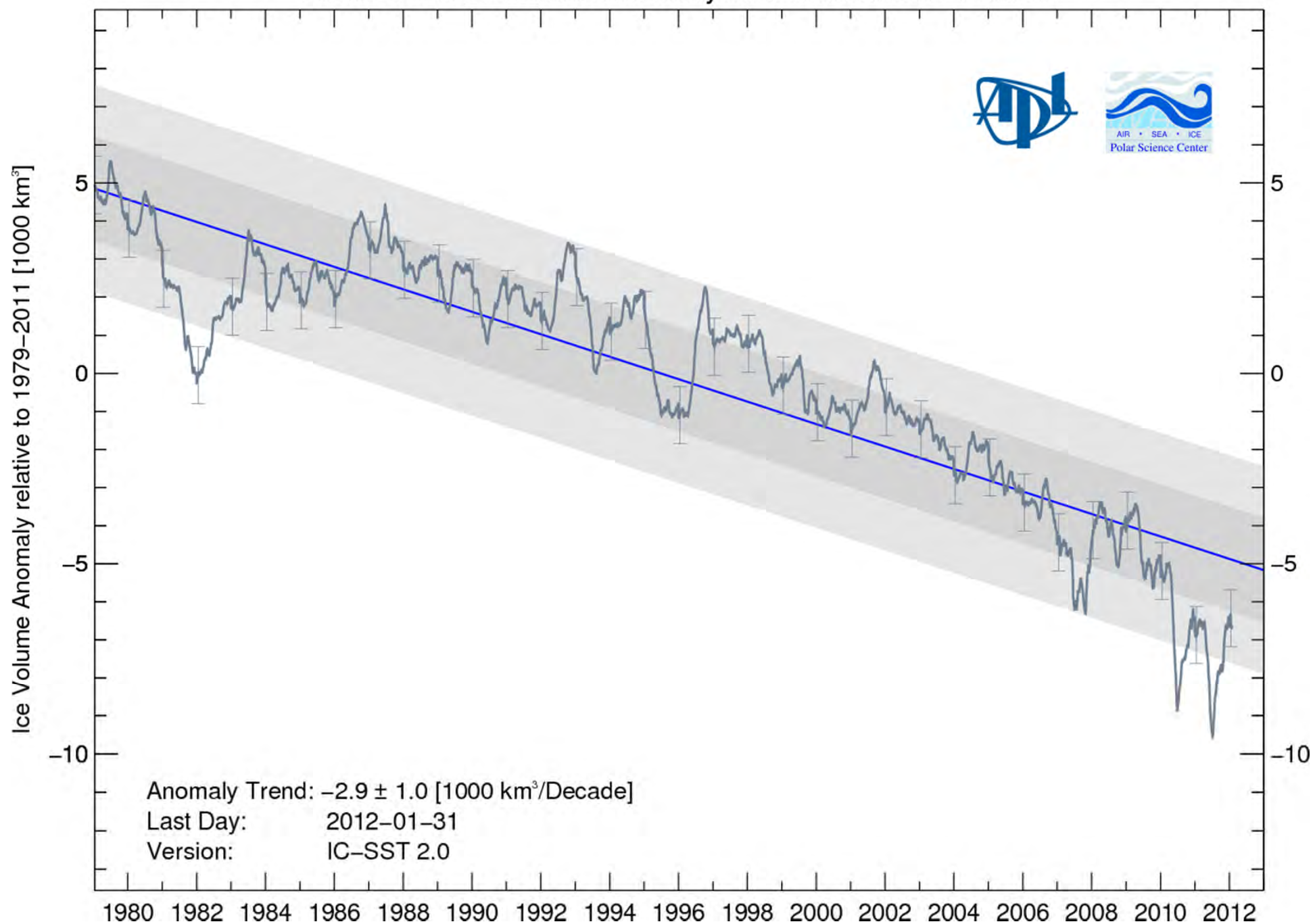


- Analysis of submarine and satellite data
- Rothrock et al. show thinning everywhere!
- Average decrease was 40%
- From 3 m to under 2 m

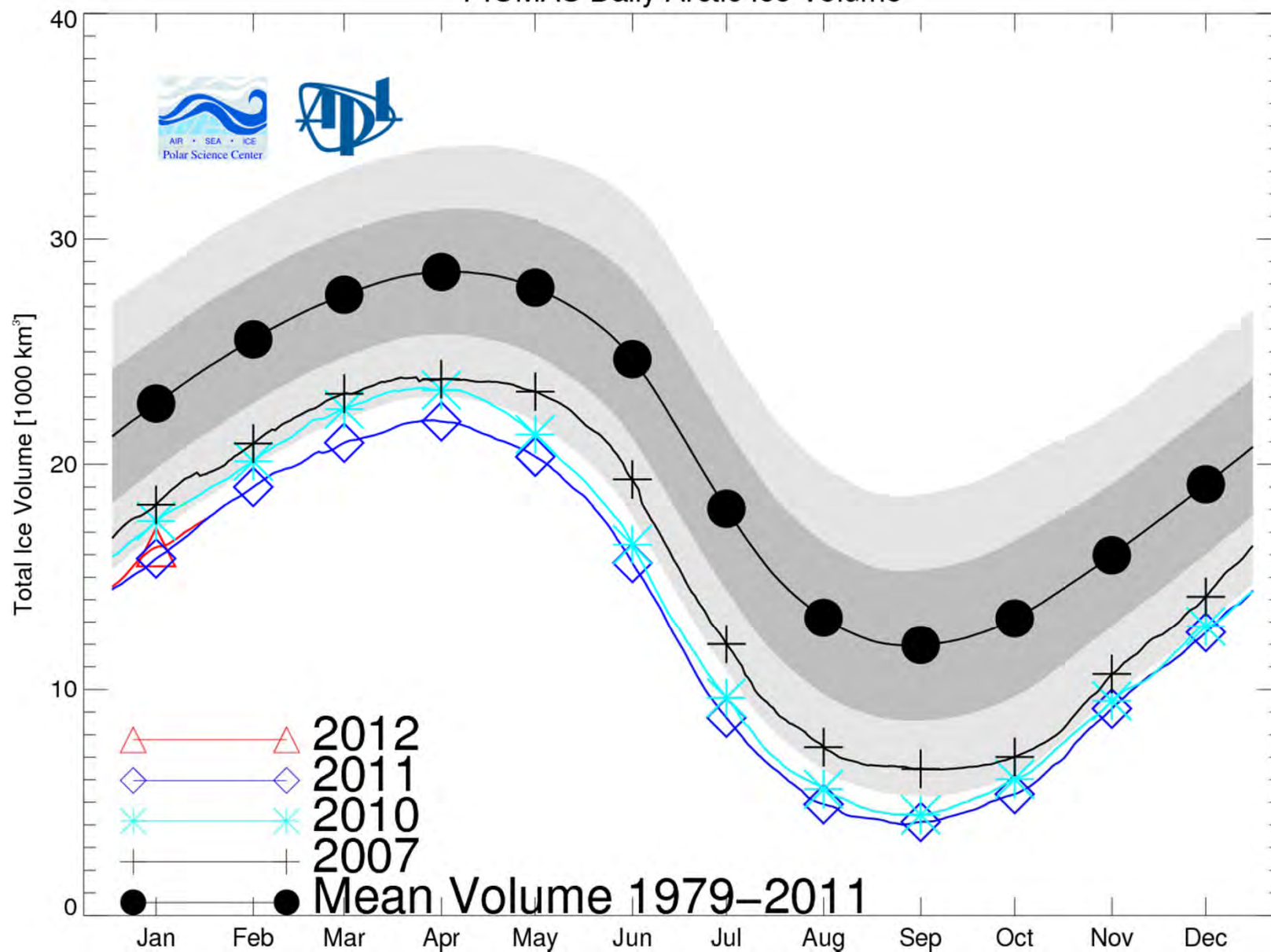
Sea ice is thinning ... everywhere

Courtesy of Don Perovich

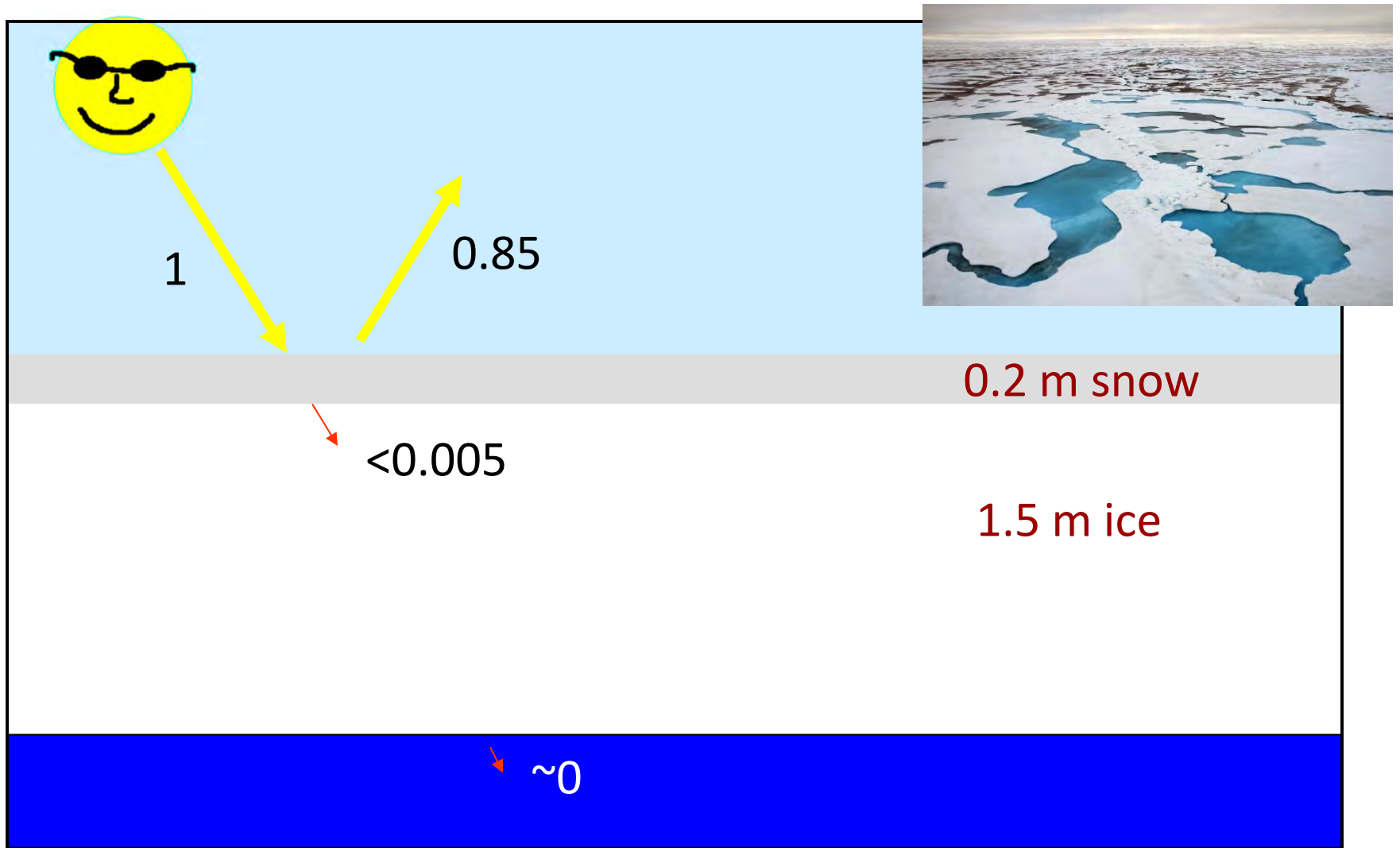
Arctic Sea Ice Volume Anomaly and Trend from PIOMAS



PIOMAS Daily Arctic Ice Volume



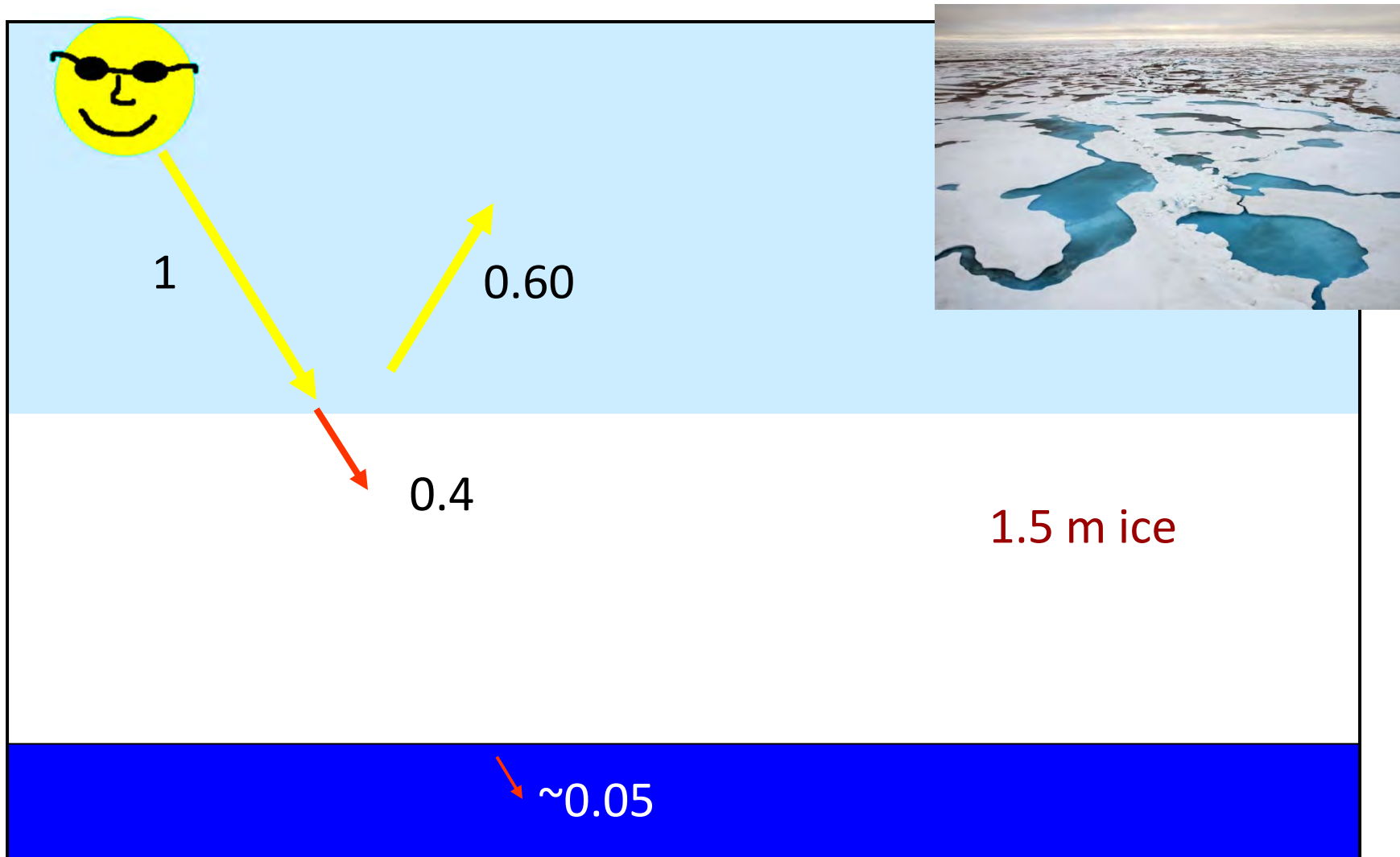
Today: Assume a simple thin slab



Snow covered, thin, first year ice

Courtesy of Don Perovich

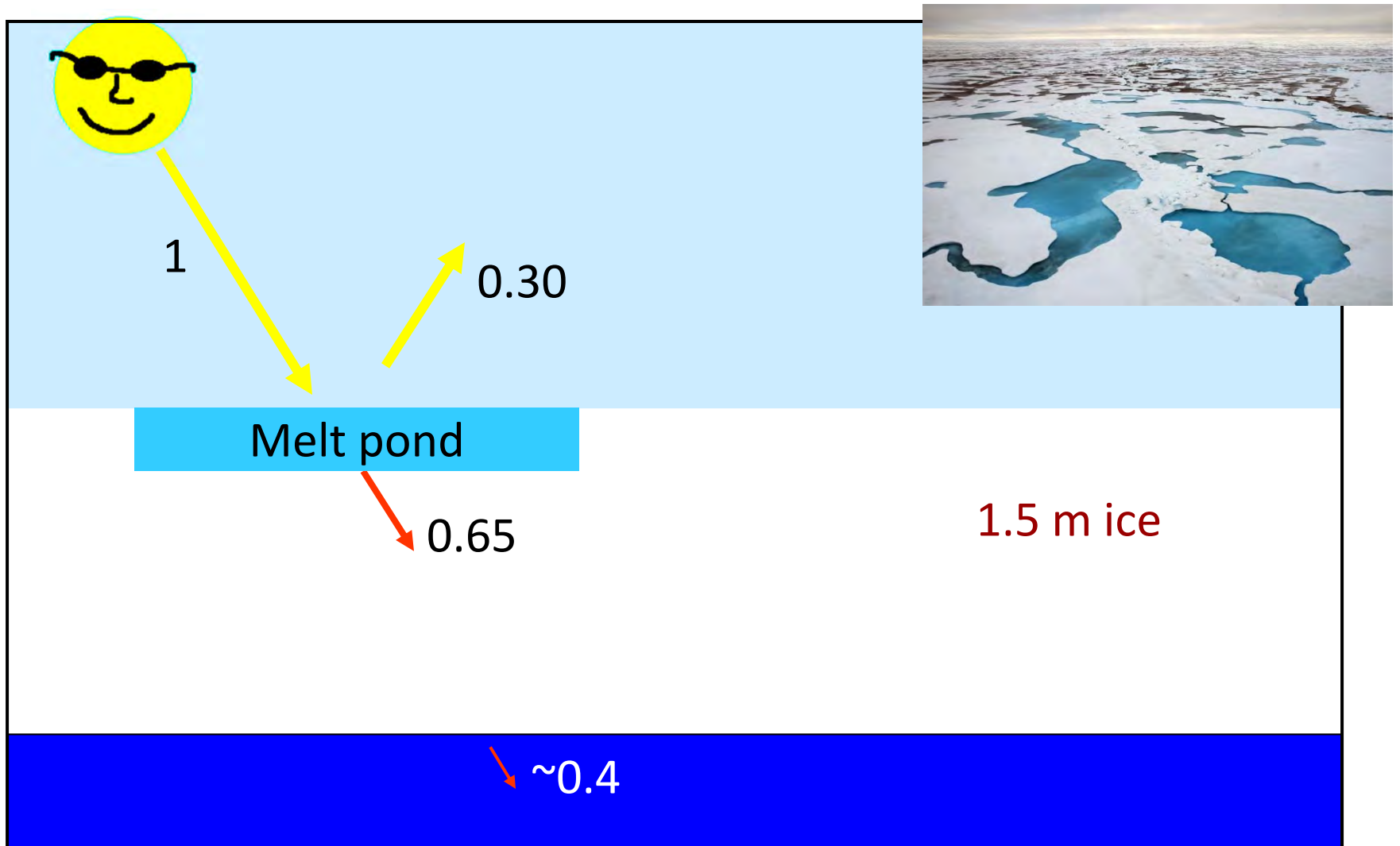
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Let the snow melt

Courtesy of Don Perovich

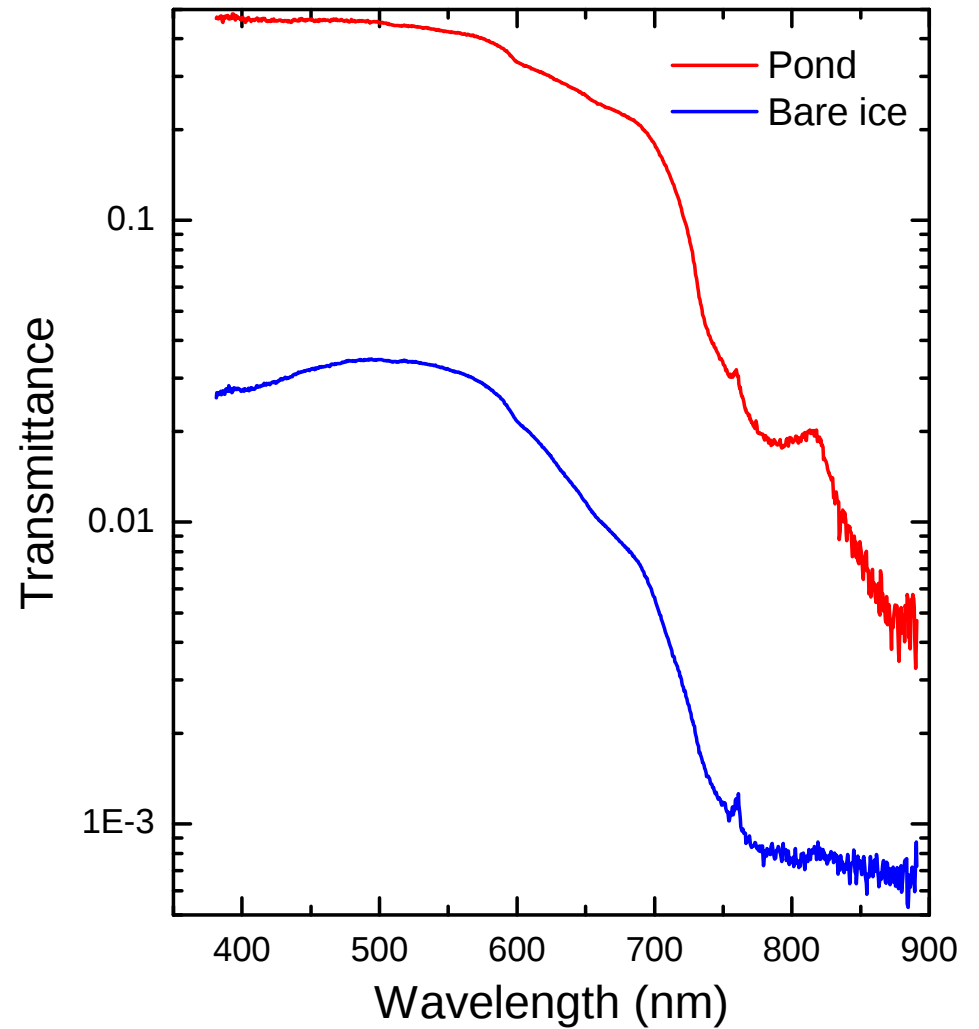
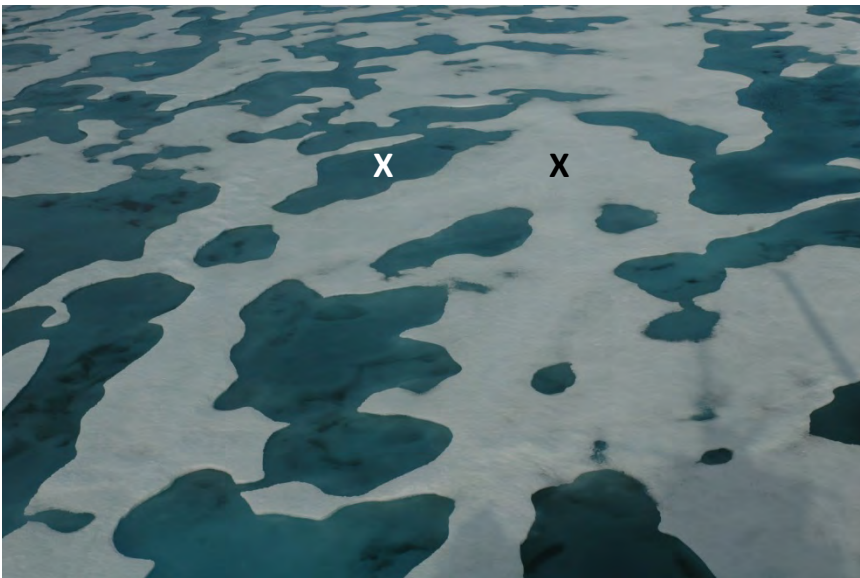
Today: Assume a simple thin slab



Let the snow melt and add a pond

Courtesy of Don Perovich

Spectral transmittance

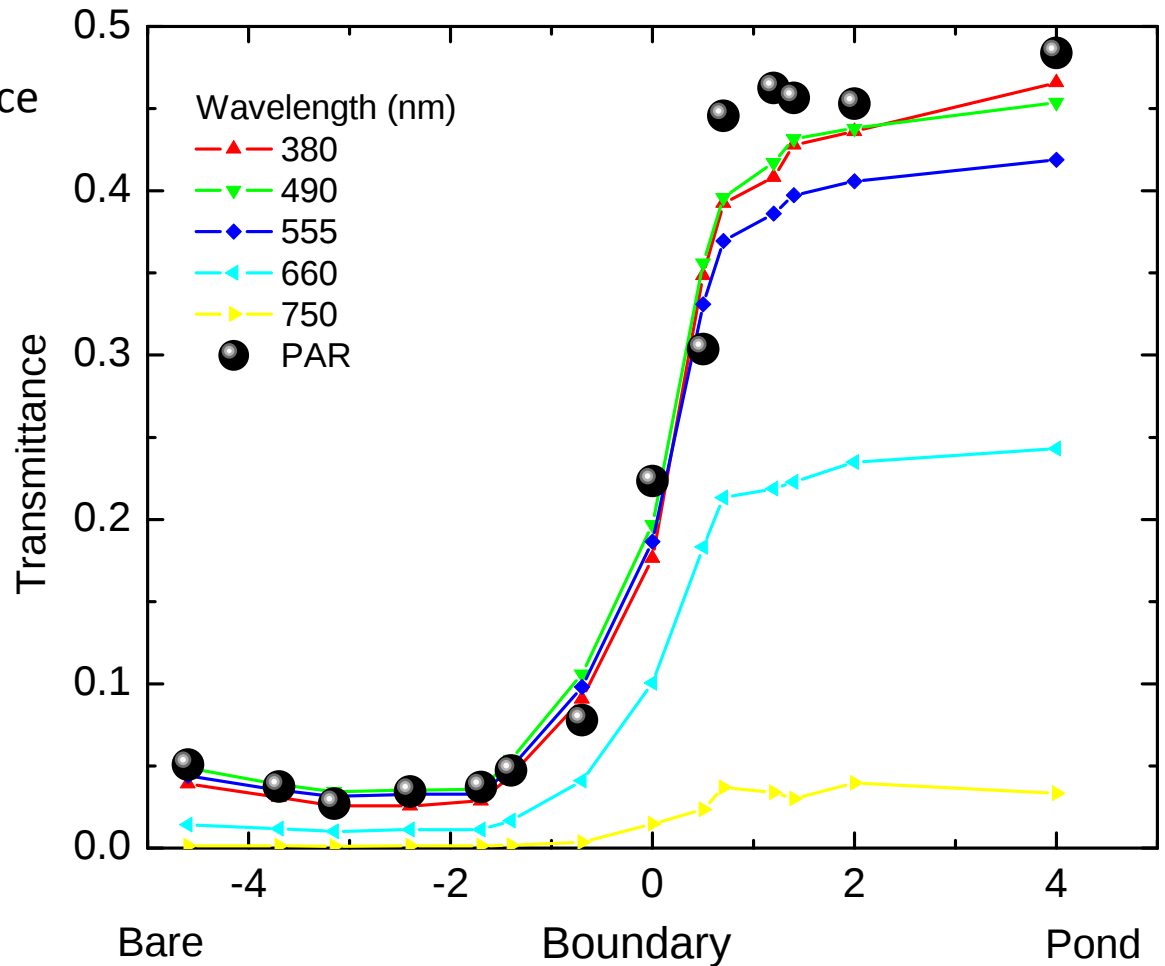


Pond transmittance order of magnitude higher than bare ice

Courtesy of Don Perovich

Spatial variability in light field

- Transmittance transition
- Distance ~ 3 to 4 X thickness
- Short wavelengths = longer distance



The sea ice light field has tremendous spatial variability

Courtesy of Don Perovich

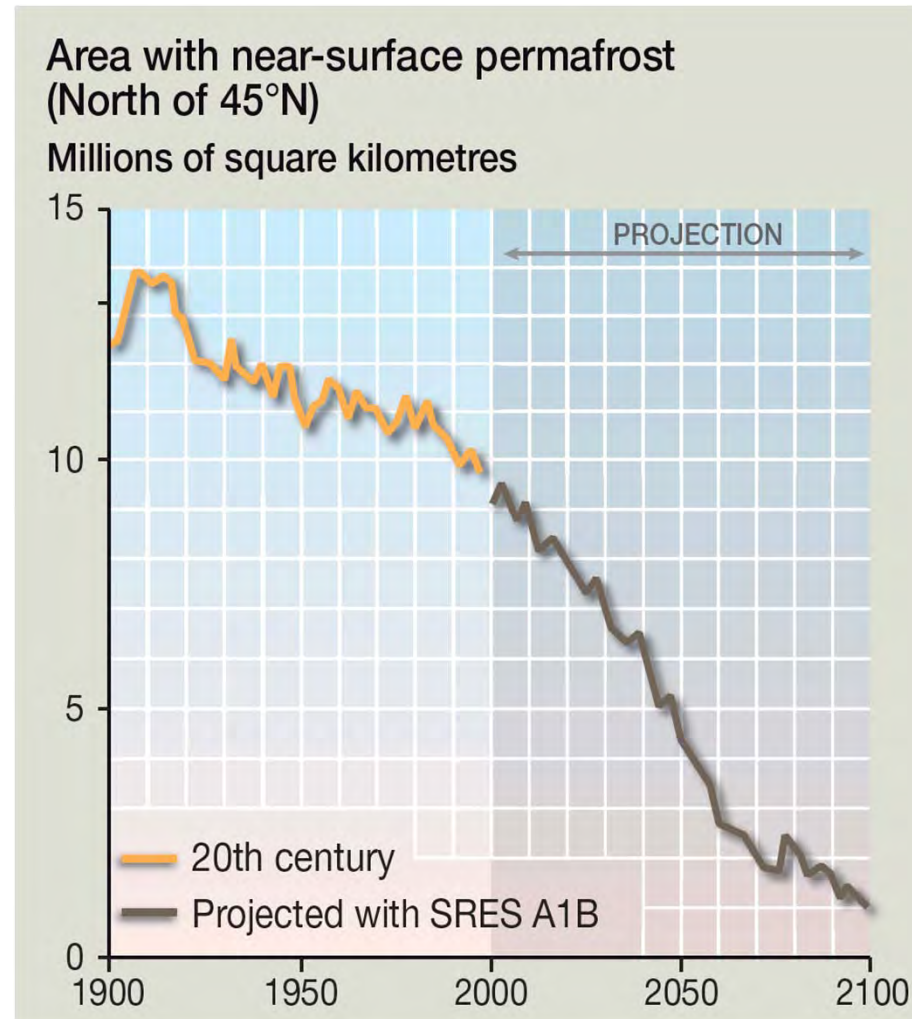
Ongoing changes

+

- Permafrost thawing

|

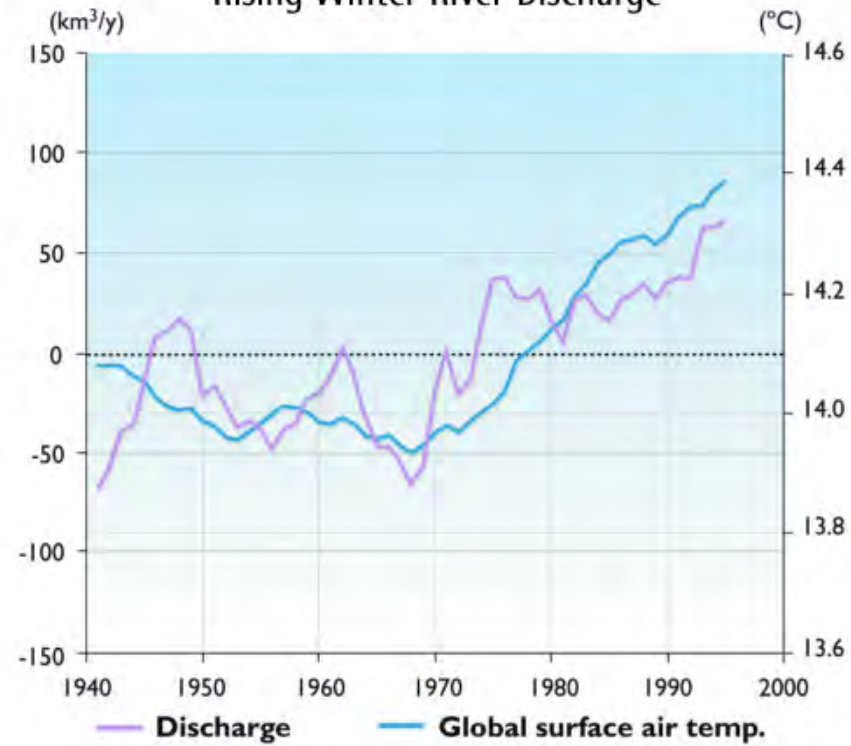
Ongoing changes



Ongoing changes



Rising Winter River Discharge



©2004, ACIA

Ongoing changes

+

- Permafrost thawing
- Increase in river runoff (+7% from 1936 to 1999)
- UV increase (+15% since 1979)

In brief

- Temperature has risen twice faster in the Arctic than in other regions; air temperature is expected to increase by up to 7°C during the next century.

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- Permafrost, which represents 25% of the continental surface of the northern hemisphere, has been observed to have undergone a temperature increase since the 1960s and, in many places, to gradually thaw. The permafrost contains up to 50% of all soil organic carbon.
- From 1936 to 1999, an increase of 7% was observed for river discharge to the Arctic Ocean. It may impact on circulation, including deep-water formation.

In brief

- Temperature has risen twice faster in the Arctic than in others regions; air temperature is expected to increase by up to 7°C during the next century.
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- From 1936 to 1999, an increase of 7% was observed for river discharge to the Arctic Ocean. It may impact on circulation, including deep-water formation.
- The summer ice cover over the Arctic Ocean decreased by more than 30% since 1979; it is predicted that it will disappear almost completely by the end of the century.

In brief

- Temperature has risen twice faster in the Arctic than in others regions; air temperature is expected to increase by up to 7°C during the next century.
- Permafrost, which represents 25% of the continental surface of the northern hemisphere, has been observed to have undergone a temperature increase since the 1960s and, in many places, a gradual thaw. The permafrost contain up to 50% of all soil organic carbon.
- From 1936 to 1999, an increase of 7% was observed for river discharge to the Arctic Ocean. It may impact on circulation, including deep-water formation.
- The summer ice cover over the Arctic Ocean decreased by 20% over the last 26 years; it is predicted that it will disappear almost completely by the end of the century.
- The amount of atmospheric ozone above the Arctic during the spring, has decreased by 10 to 15 % since 1979.

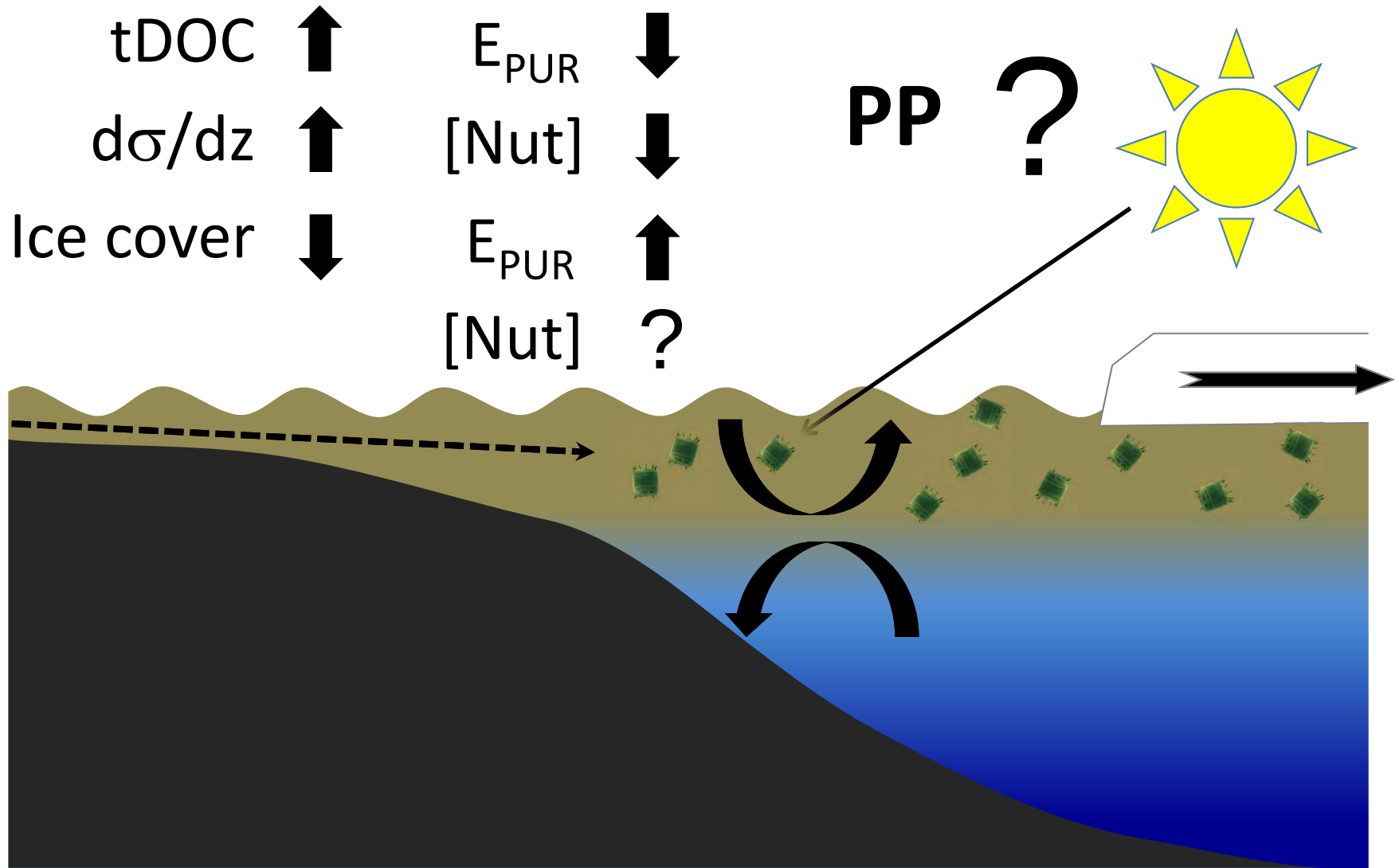
One may expect that...

1. An increasing fraction of the organic carbon sequestered into the permafrost will be transported toward the Arctic Ocean together with inorganic nutrients
2. The Arctic Ocean surface layer will increasingly be exposed to light, including UV
3. The organic matter of terrestrial origin will be oxidized to CO₂ both through photo-oxidation, and bacterial activity amplified by light
4. Photosynthesis will be increasingly stimulated by light and inorganic nutrients, and will lead to more carbon sequestration

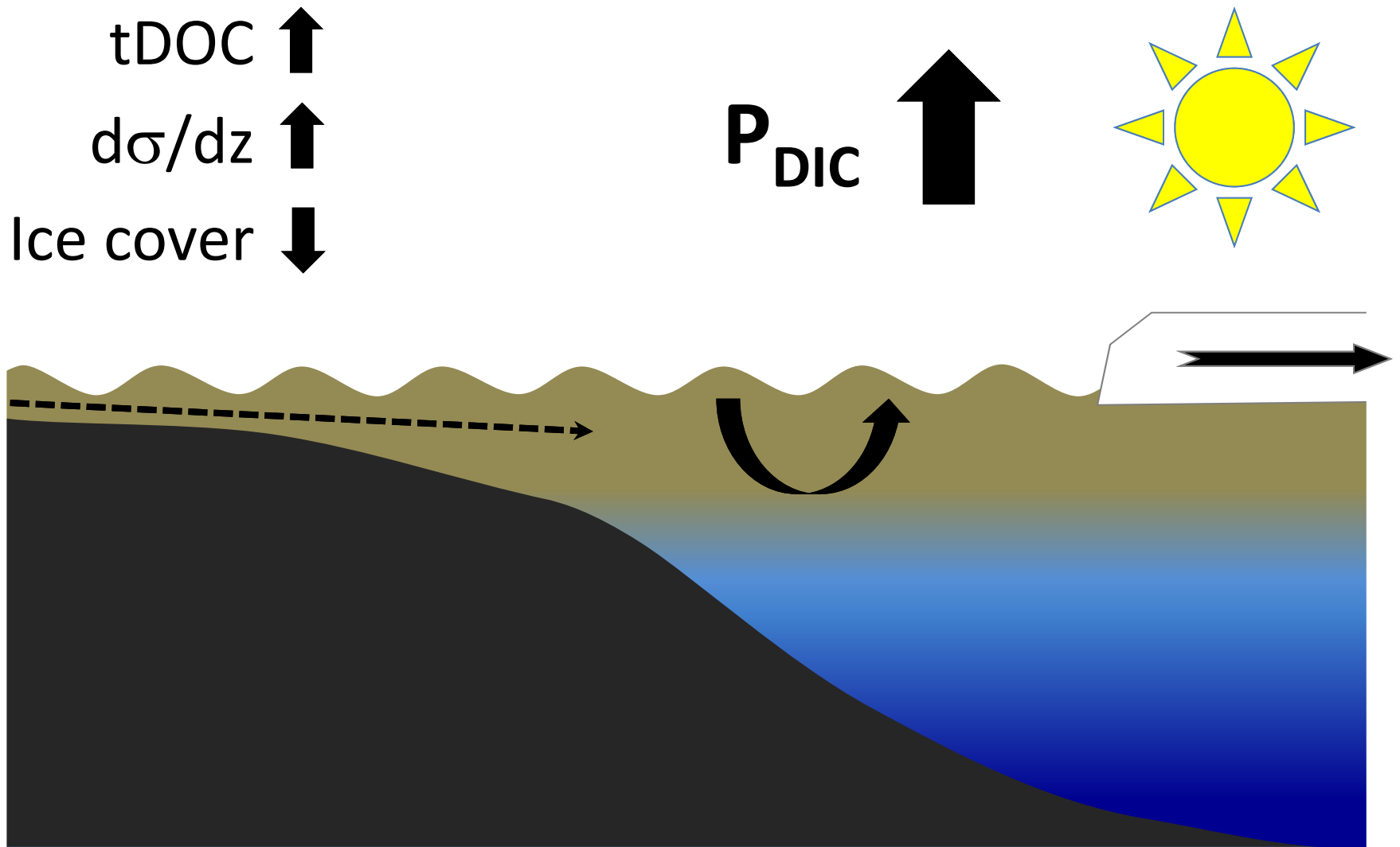
And wonder whether...

The Arctic Ocean will become a new net source of CO₂ originating from organic carbon that was sequestered in the permafrost (analogous to the combustion of fossil fuel), or a stronger biological sink of CO₂ leading to more sequestration of carbon in the sediments

What's the possible impact on PP?



And on P_{DIC} ?

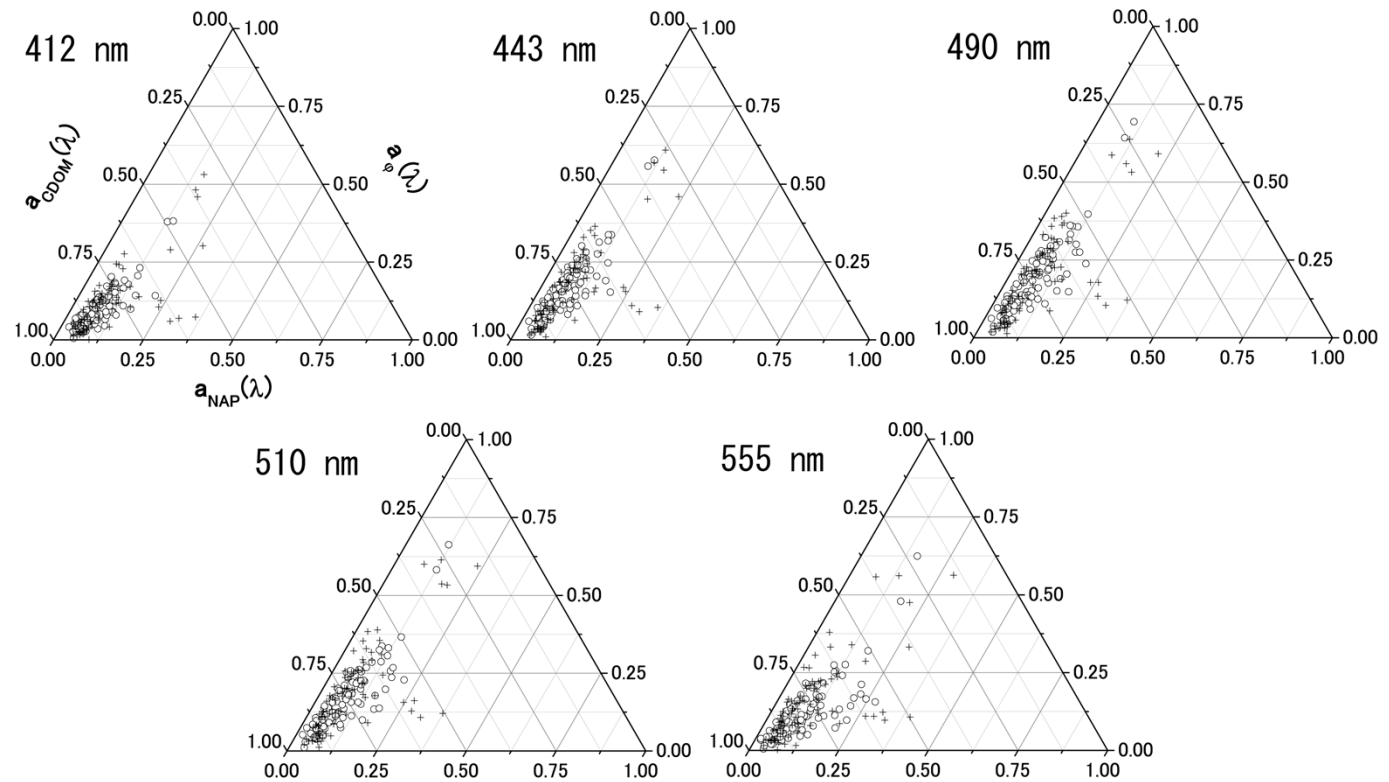


1. Ocean colour remote sensing in polar seas

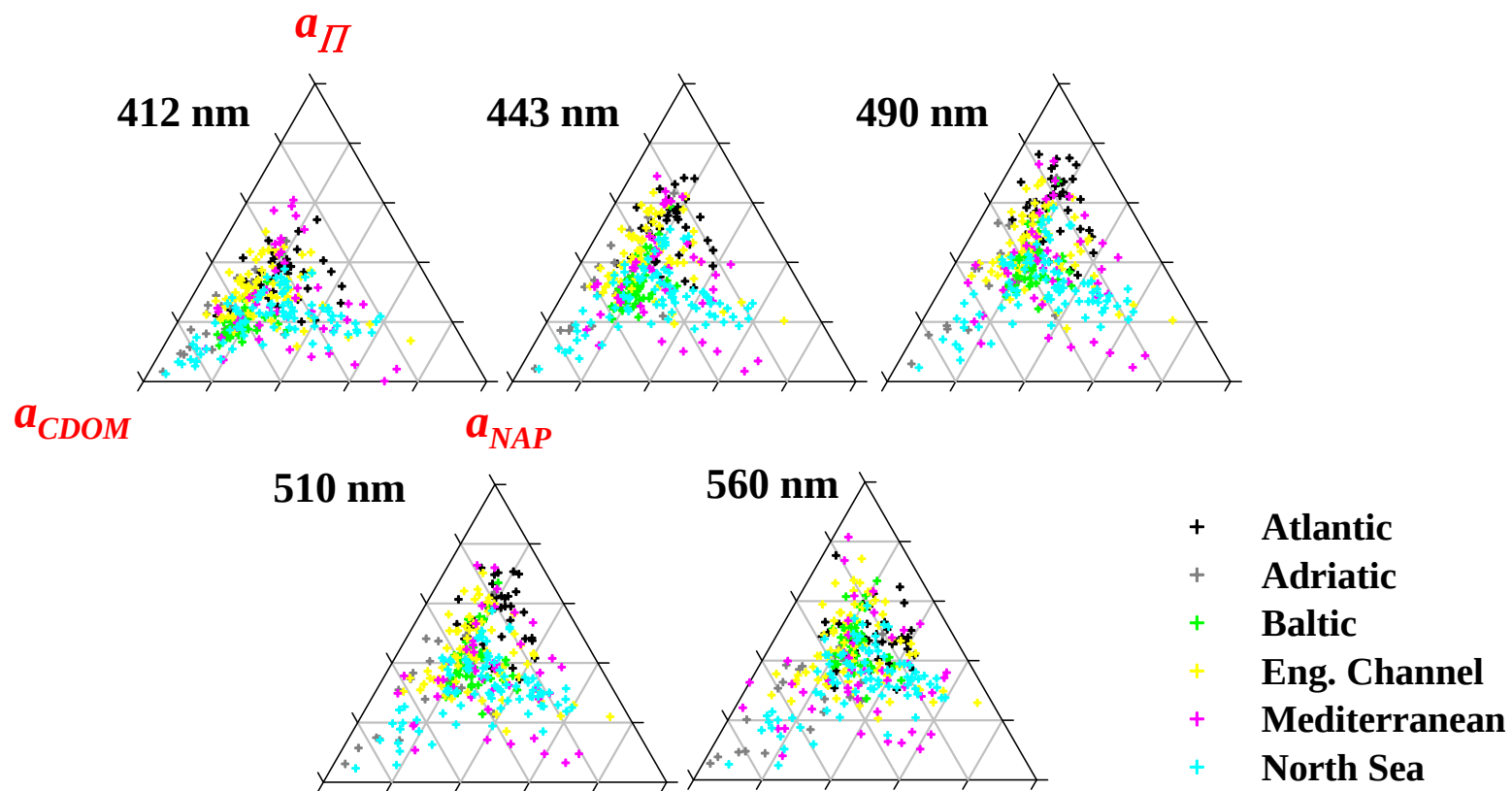
1.2. Seawater optical properties

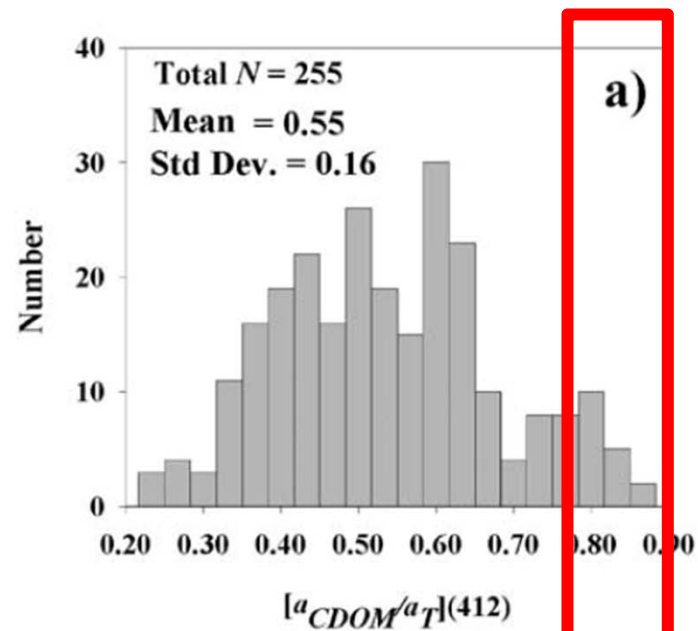
What's special?

Arctic waters are optically dominated by CDOM

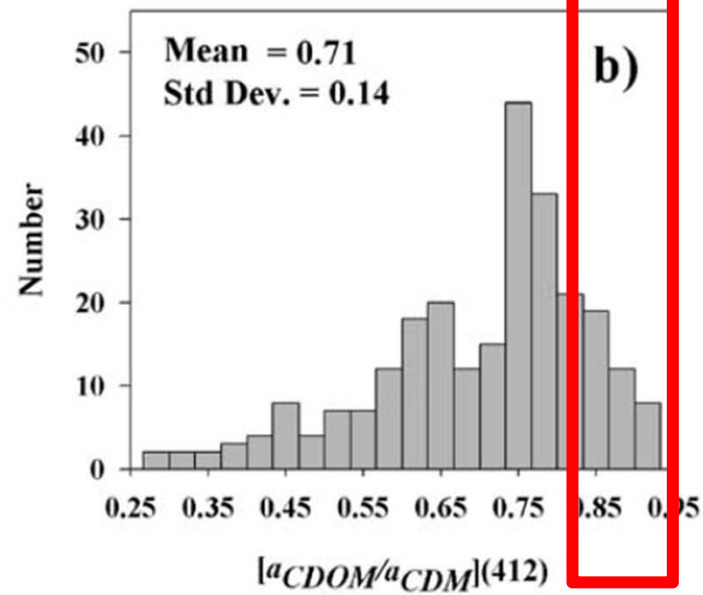


Matsuoka et al. (2007)

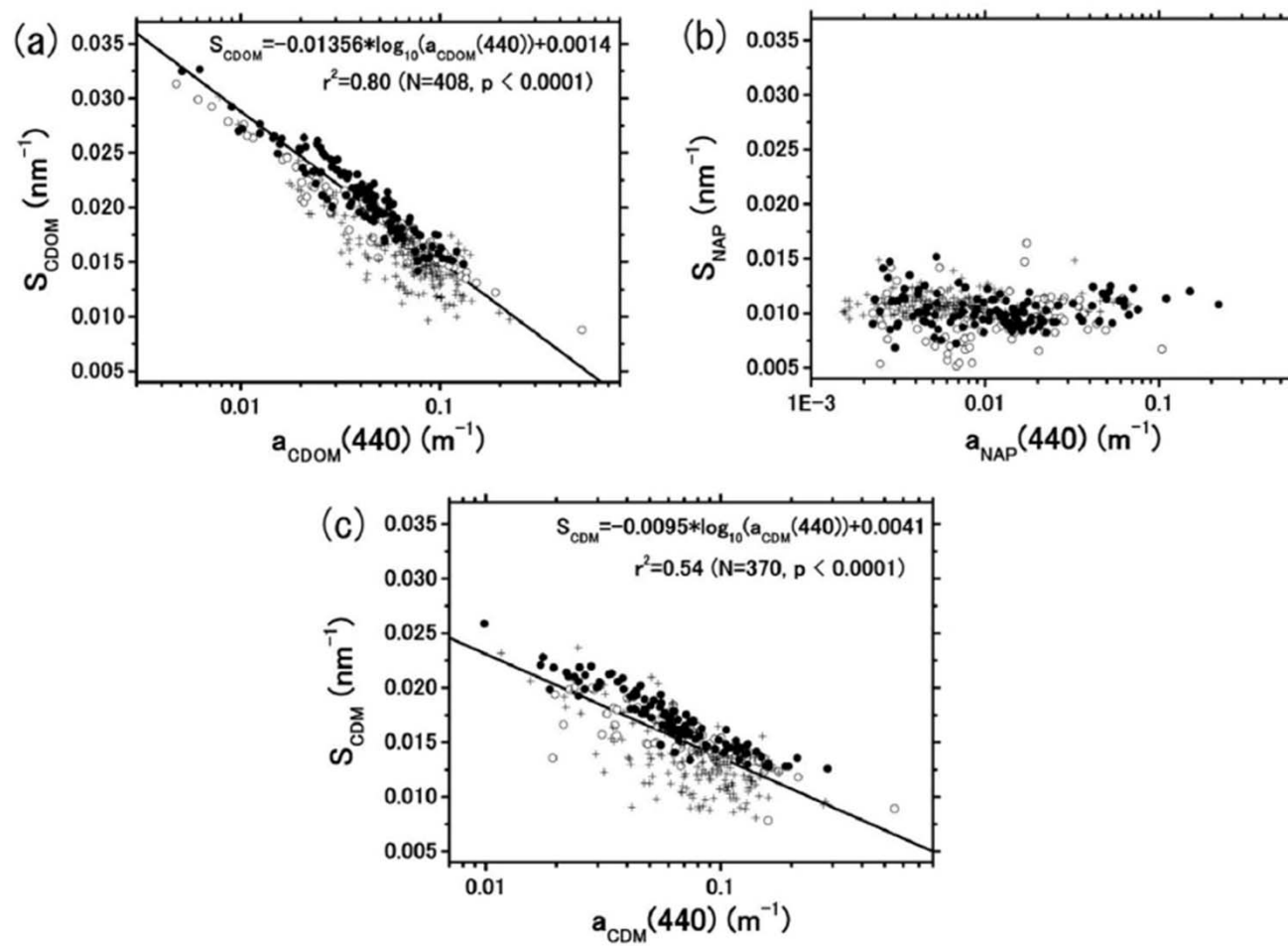




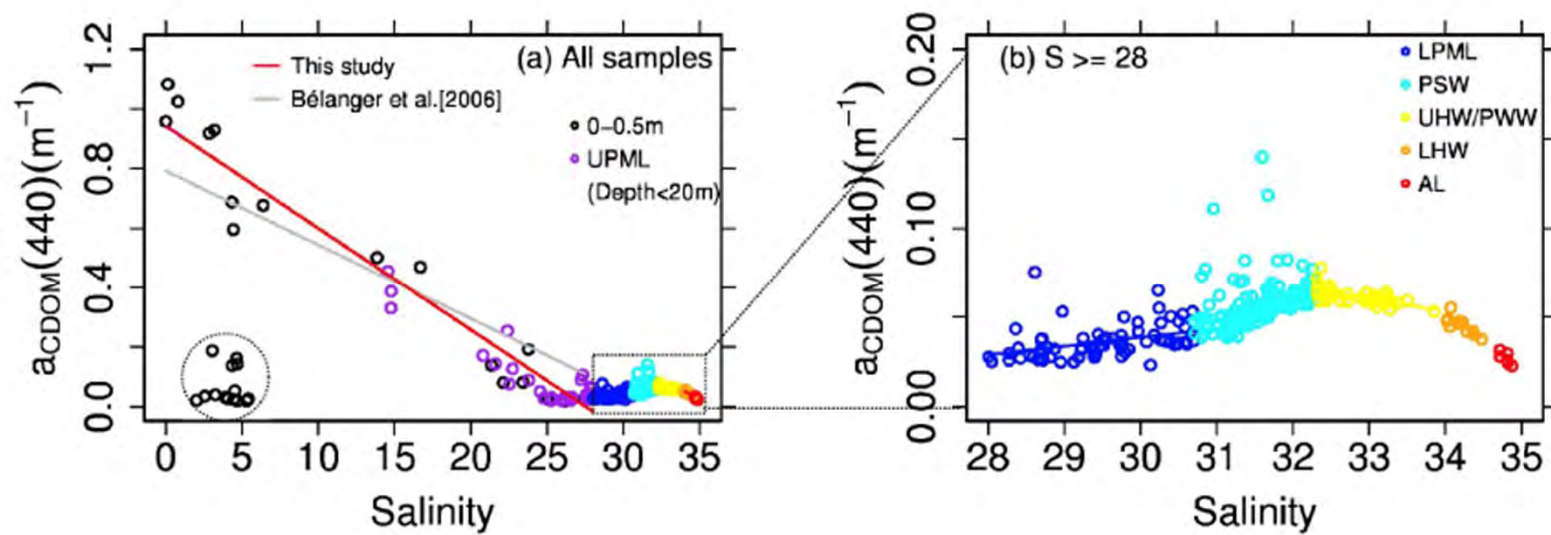
Beaufort Sea



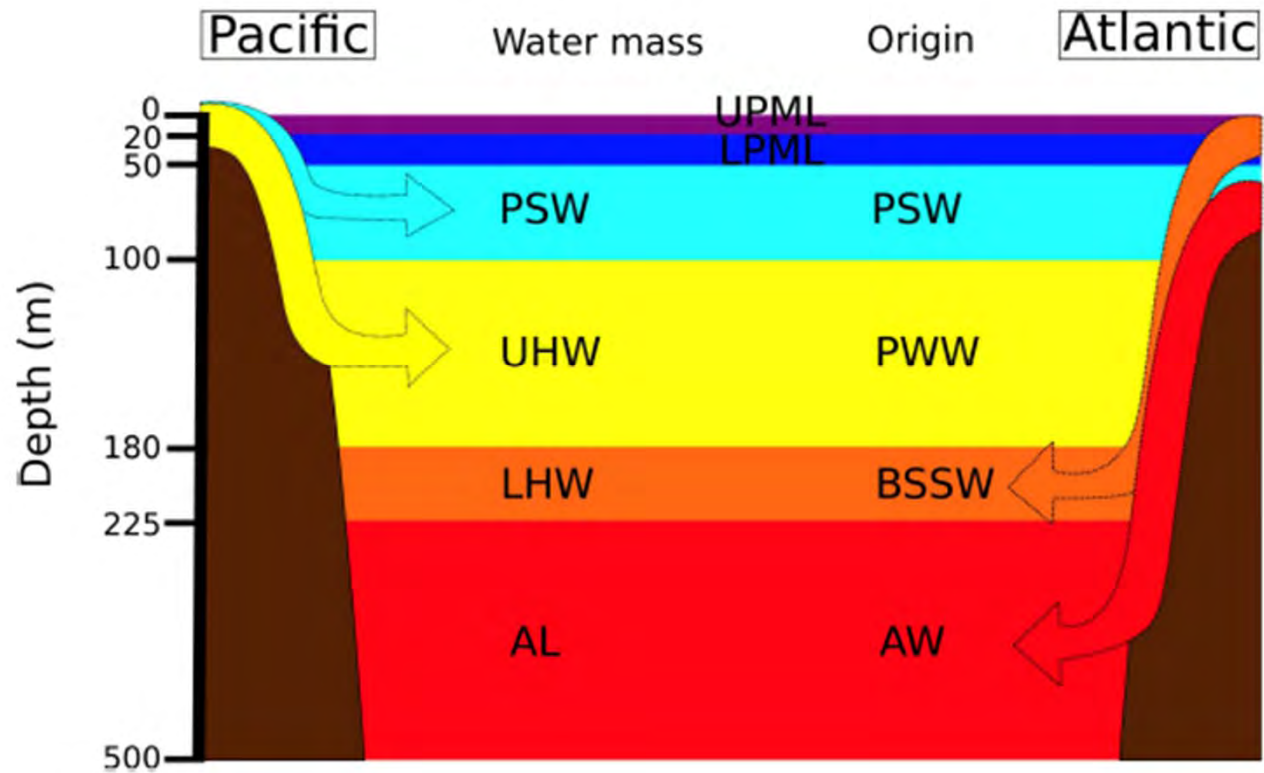
Bélangier et al. (2008)



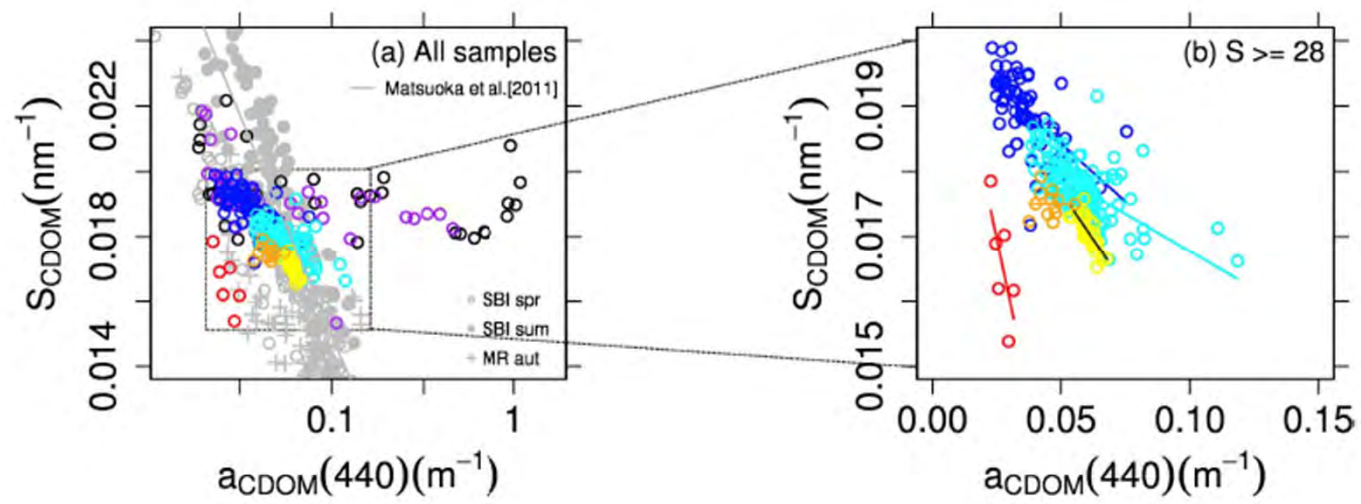
Matsuoka et al. (2011)



Matsuoka et al. (2012)

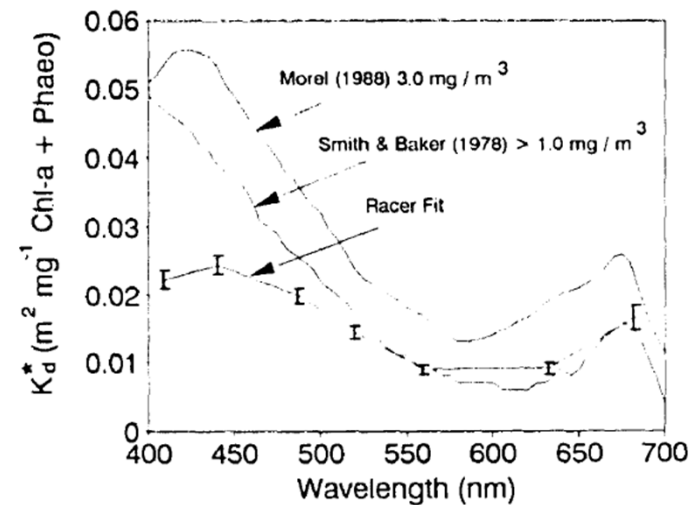


Matsuoka et al. (2012)

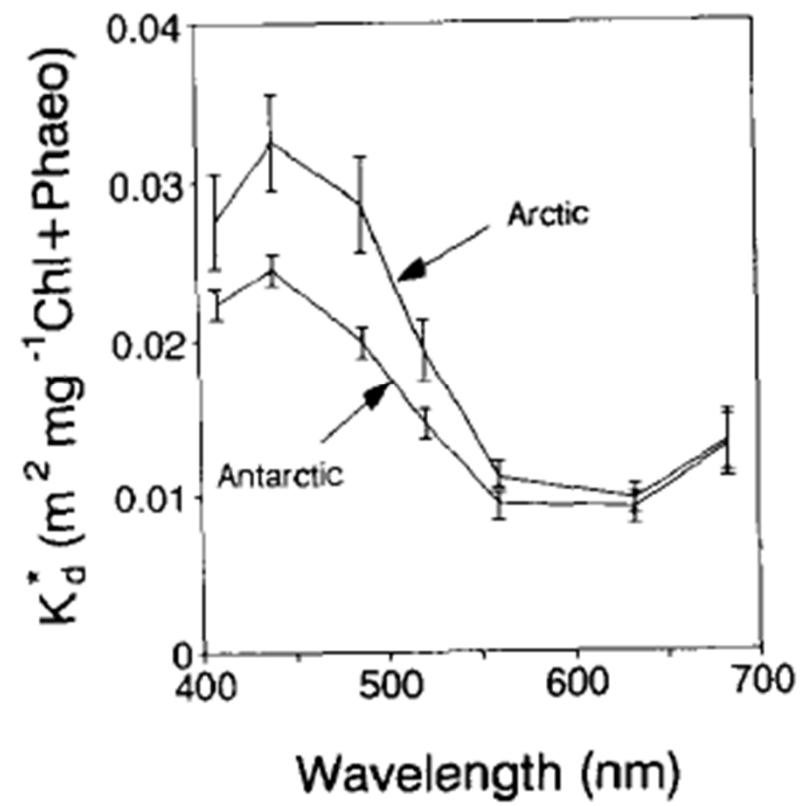


Matsuoka et al. (2012)

Chl-specific absorption is significantly smaller

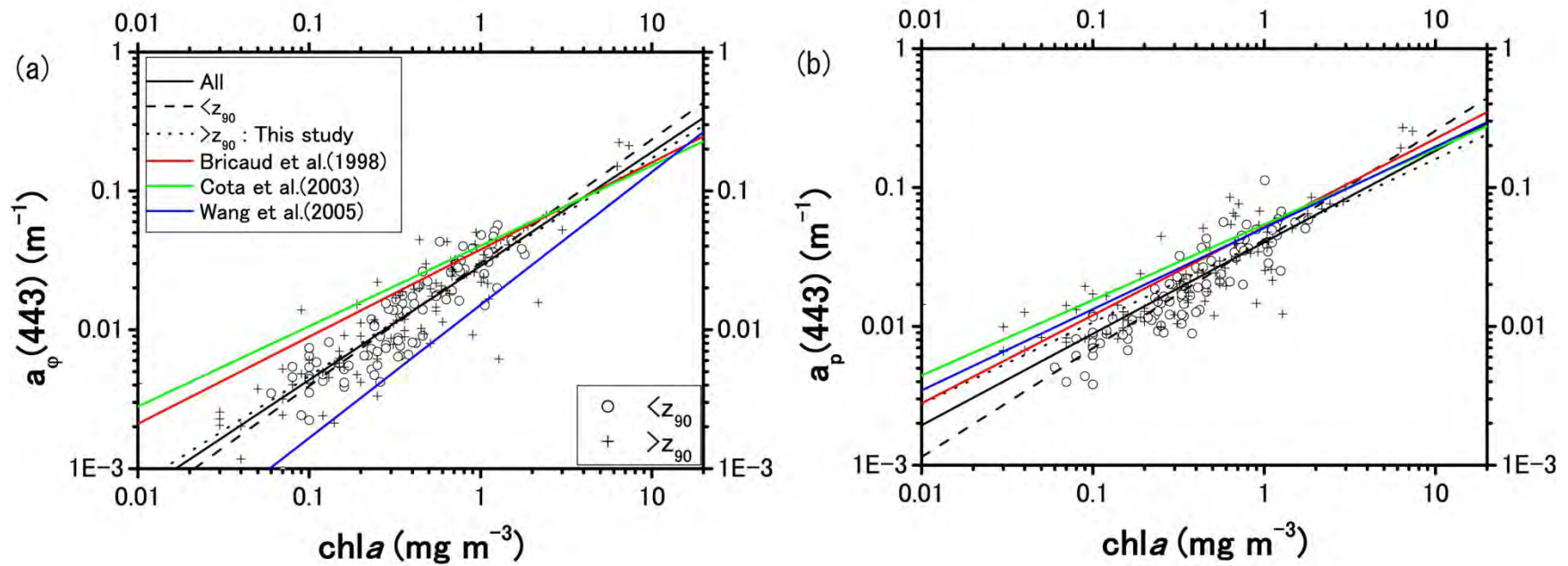


Mitchel and Holm-Hensen (1991)

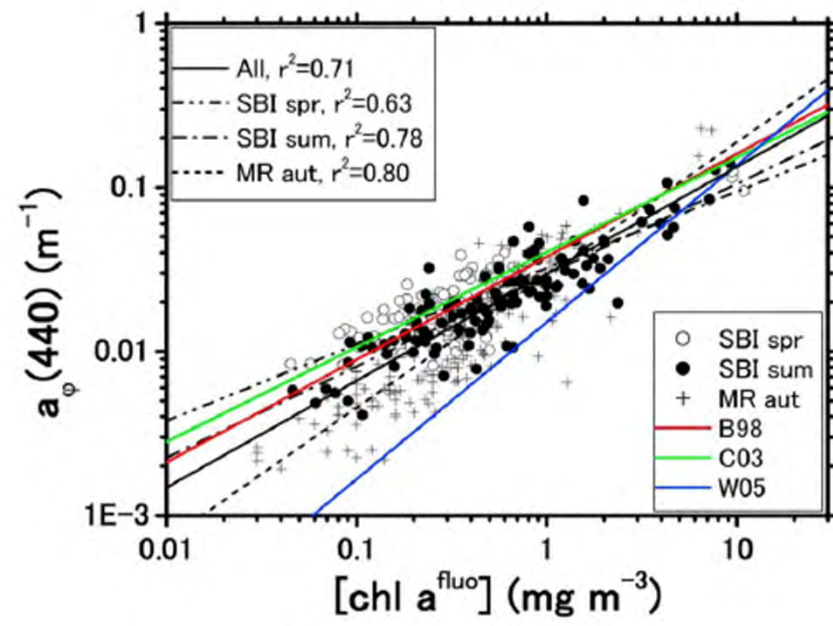


Mitchell (1990)

$\bar{a}_\phi^*$ is generally lower in the Arctic Ocean because of the light regime



Matsuoka et al. (2007)

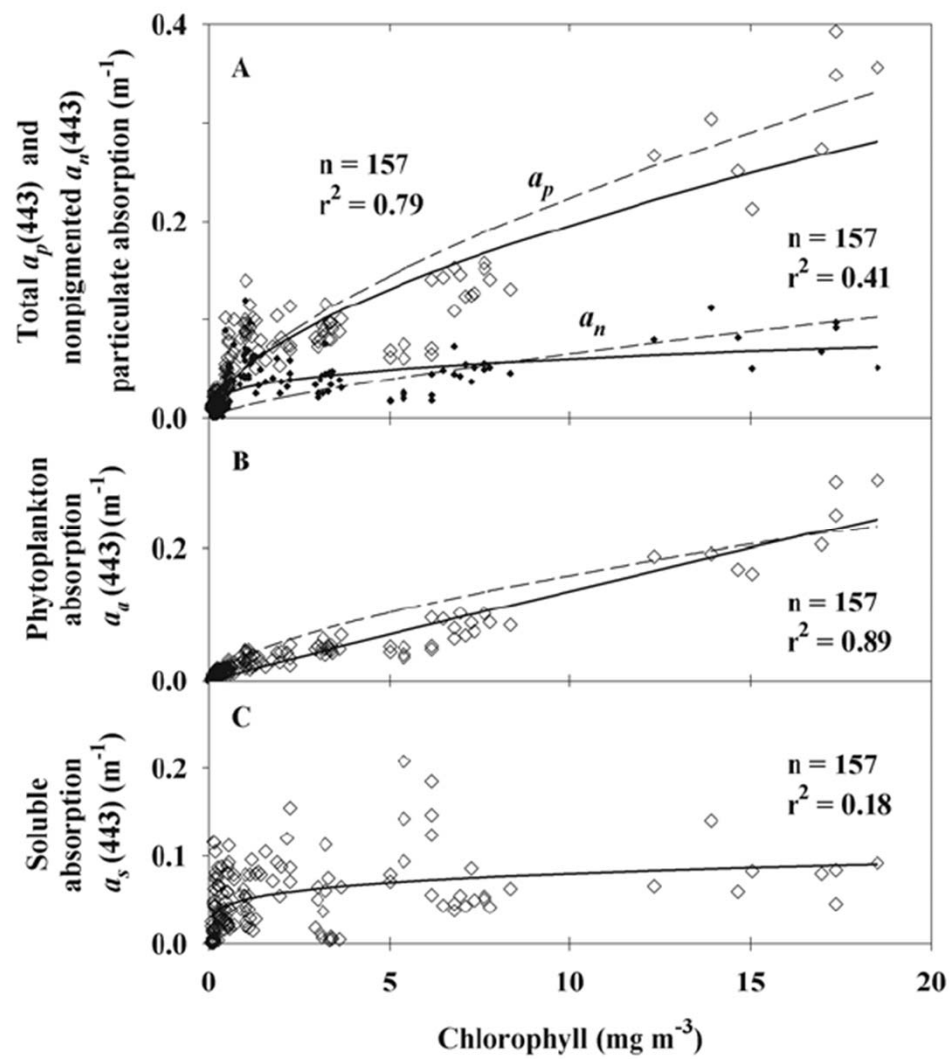


Matsuoka et al. (2011)

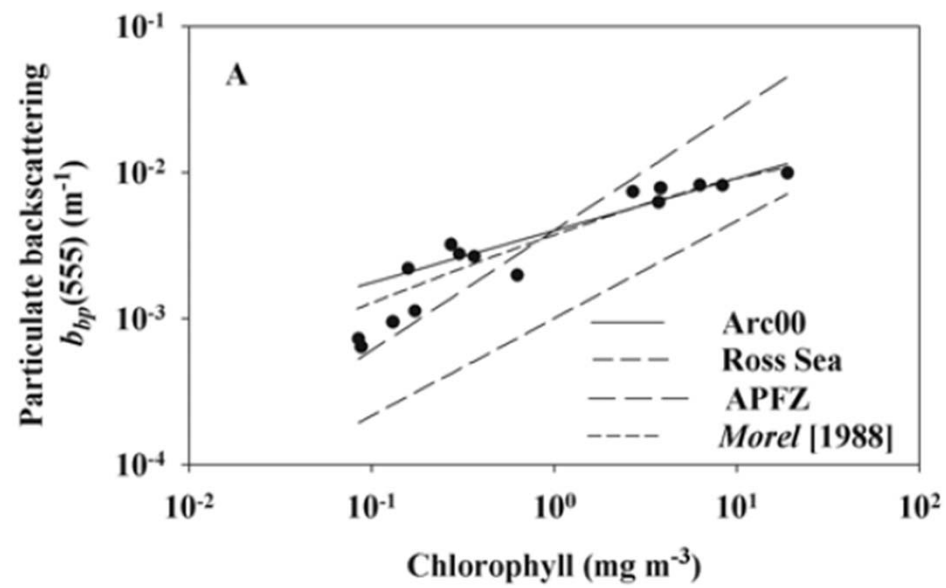
Table 1. Coefficients for the Nonlinear Regression Expressed as $a_{\phi}(\lambda) = \alpha(\lambda)[a_{\phi}(440)]^{\beta(\lambda)}$, Where λ is the Wavelength^a

λ (nm)	$\alpha(\lambda)$	$\beta(\lambda)$	r^2
400	0.8865	1.022	0.917
405	0.9035	1.011	0.945
410	0.9318	1.008	0.961
415	0.9422	1.002	0.974
420	0.9475	0.999	0.980
425	0.9575	0.998	0.986
430	0.9614	0.992	0.989
435	0.9975	0.997	0.997
440	1.0000	1.000	1.000
445	0.8995	0.990	0.993
450	0.8550	0.989	0.993
455	0.8322	0.991	0.991
460	0.8104	0.988	0.991
465	0.8011	0.991	0.990
470	0.7434	0.979	0.989

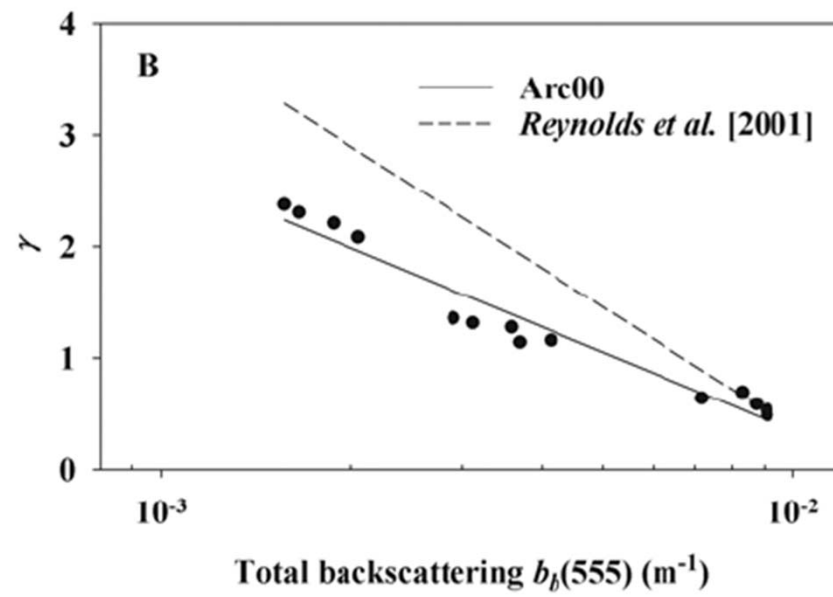
Matsuoka et al. (2011)



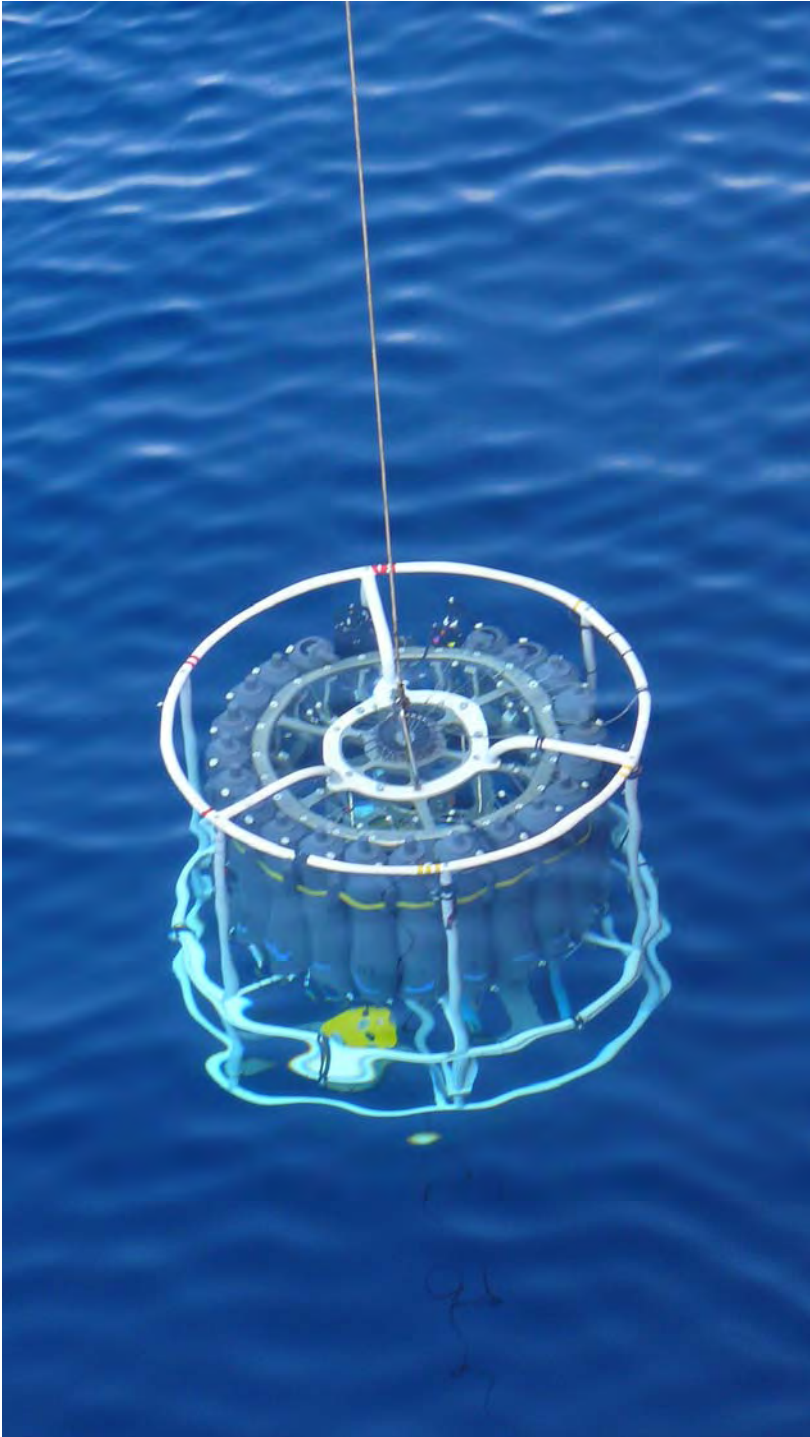
Wang and Cota (2005)



Wang and Cota (2005)



Wang and Cota (2005)



Optical properties, in brief

1. CDOM is relatively high
2. Chl-specific phytoplankton absorption is low
3. Turbid waters at the coast and in river plumes
4. But most of the time: blue clear water!

1. Ocean colour remote sensing in polar seas

1. 1.3. Retrieval of ocean properties from ocean colour

1.3.1. Atmospheric corrections

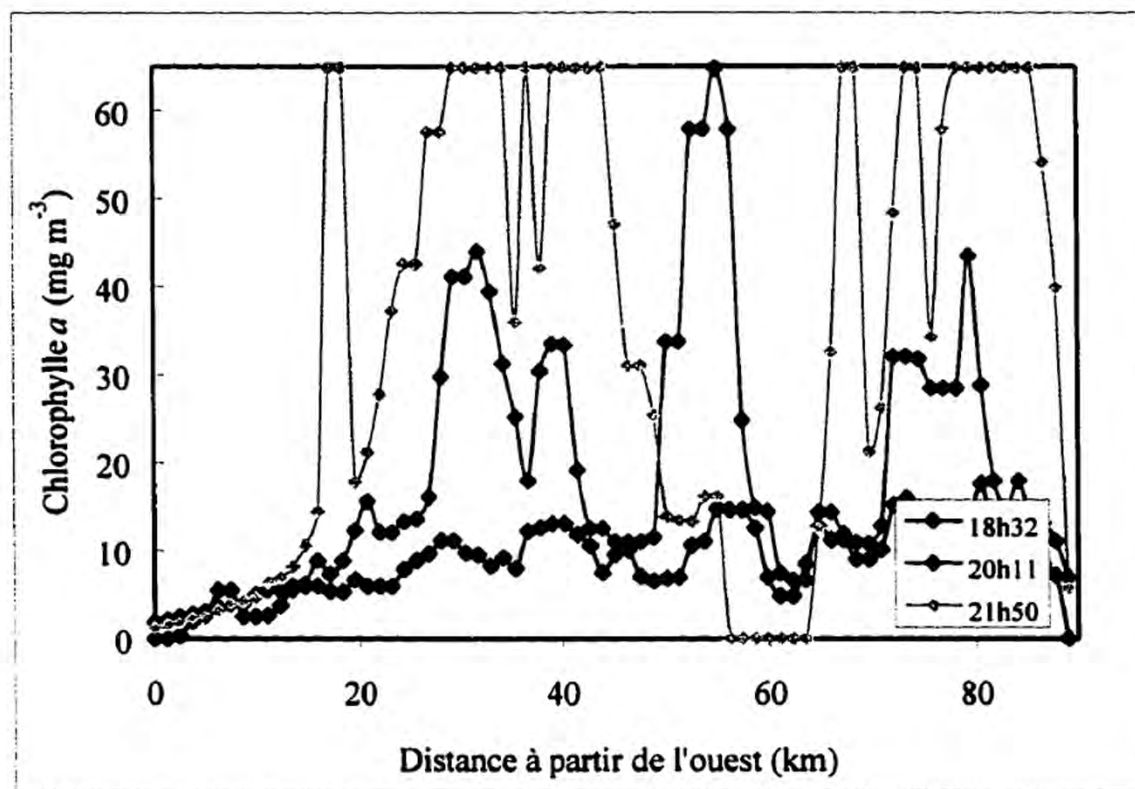
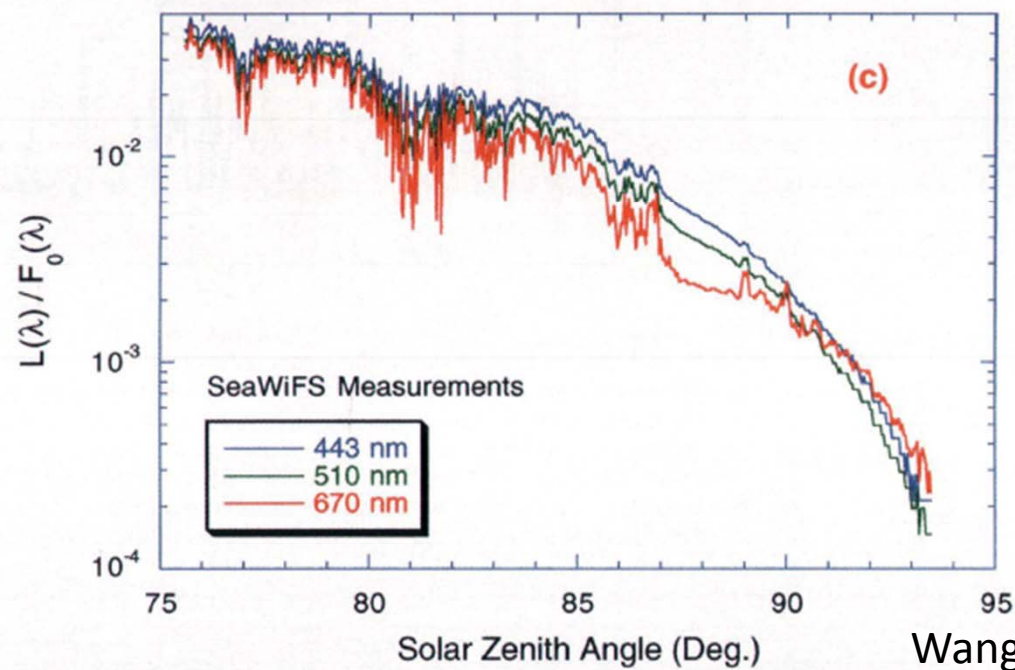
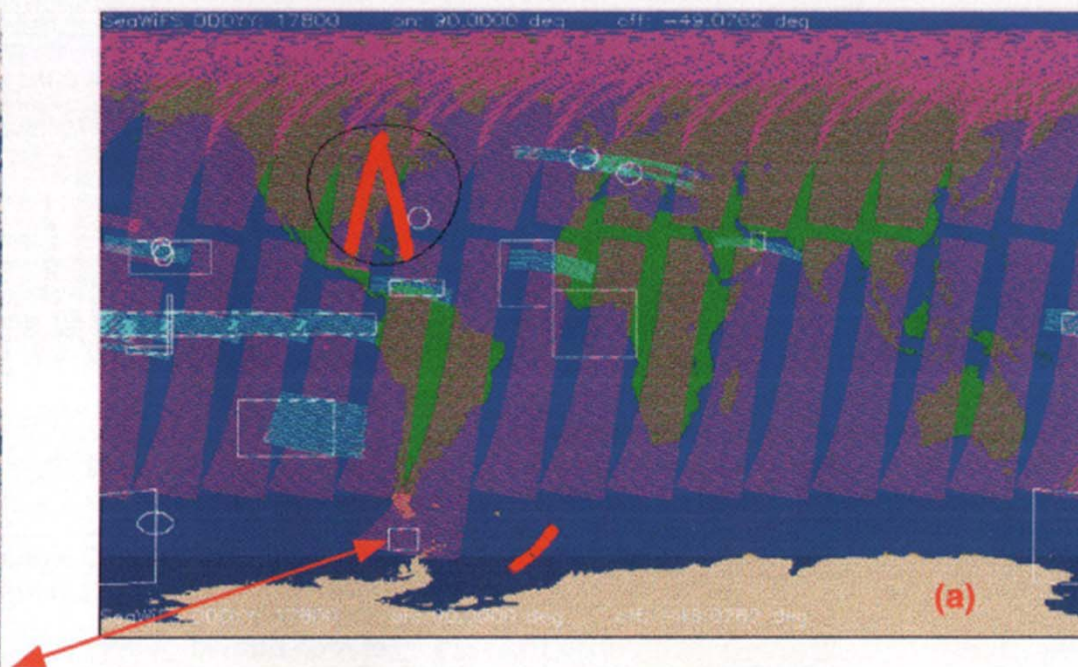
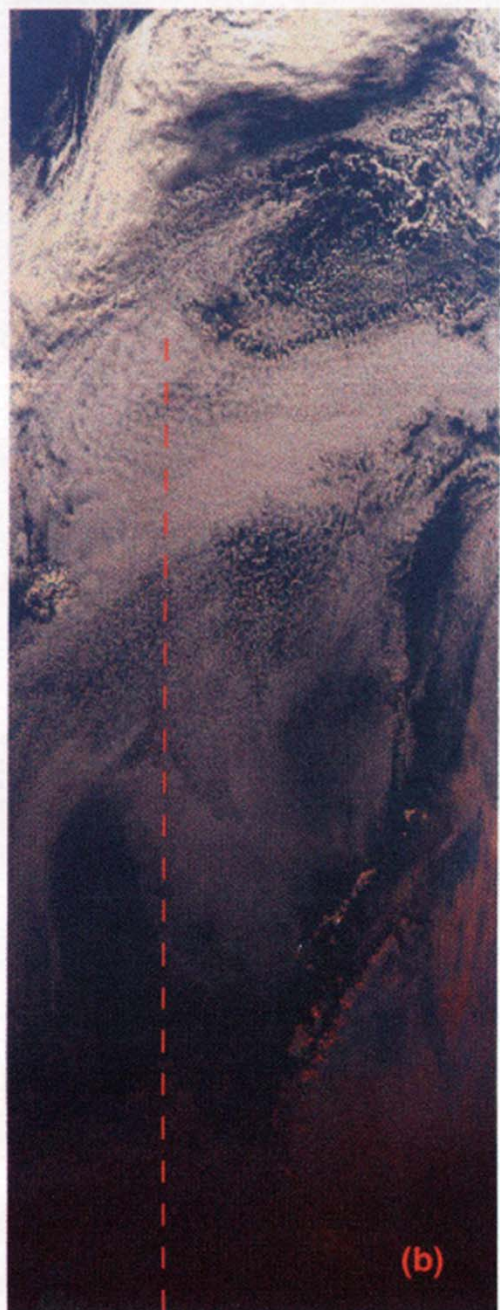


Figure 5.8. – Concentration de chlorophylle *a* pour trois heures d'acquisition le 19 juin 1999

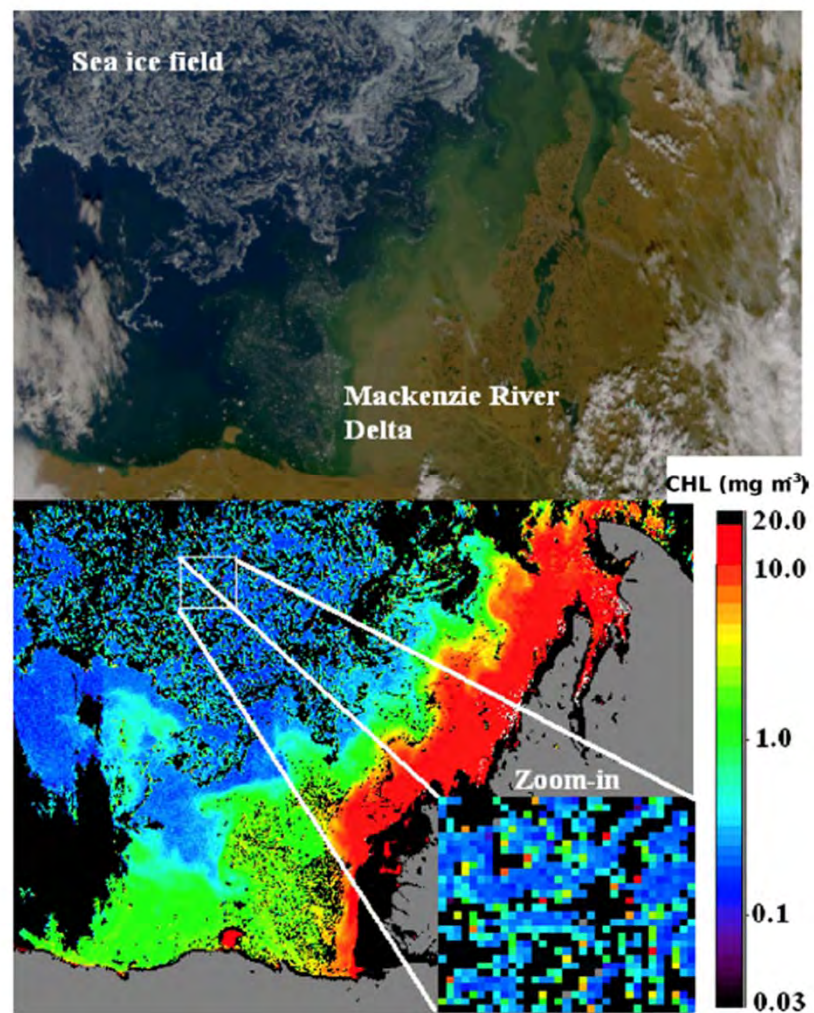
Courtesy of Simon Bélanger



1. Ocean colour remote sensing in polar seas

1.3. Retrieval of ocean properties from ocean colour

1.3.2. Contamination of the signal by sea ice

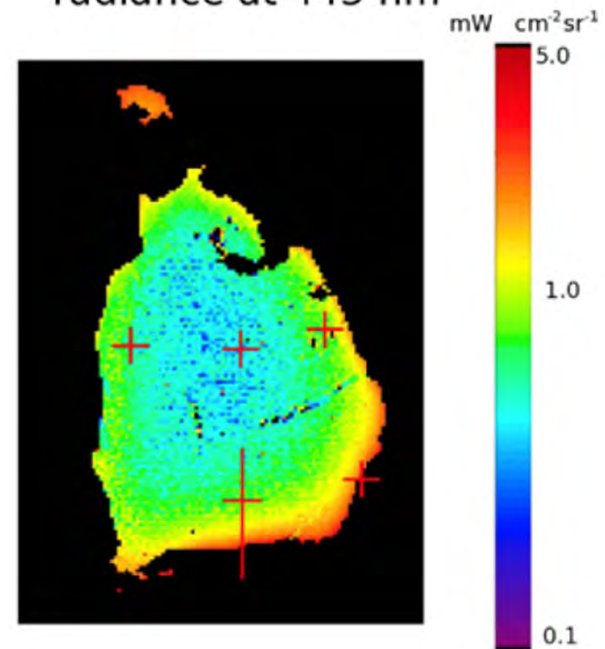


Bélanger et al. (2007)

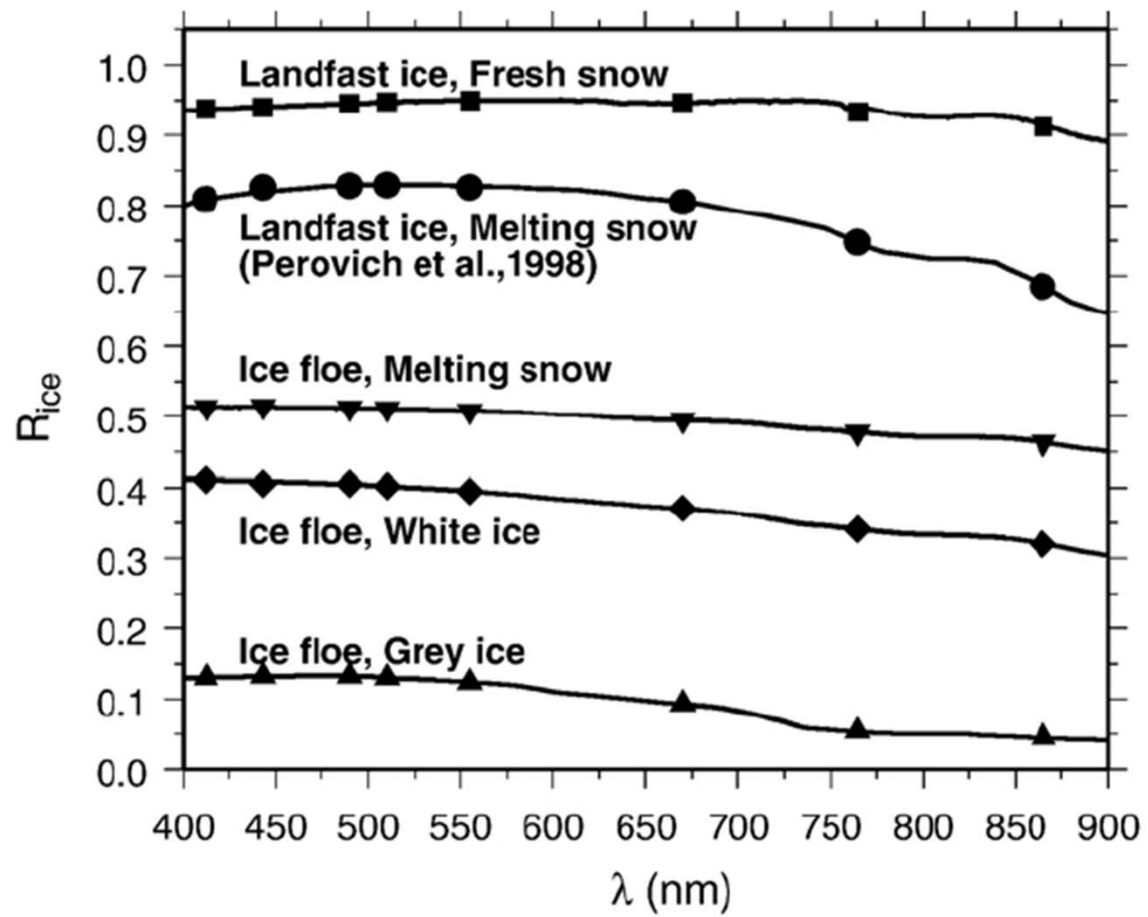
SeaWiFS quasi-true color



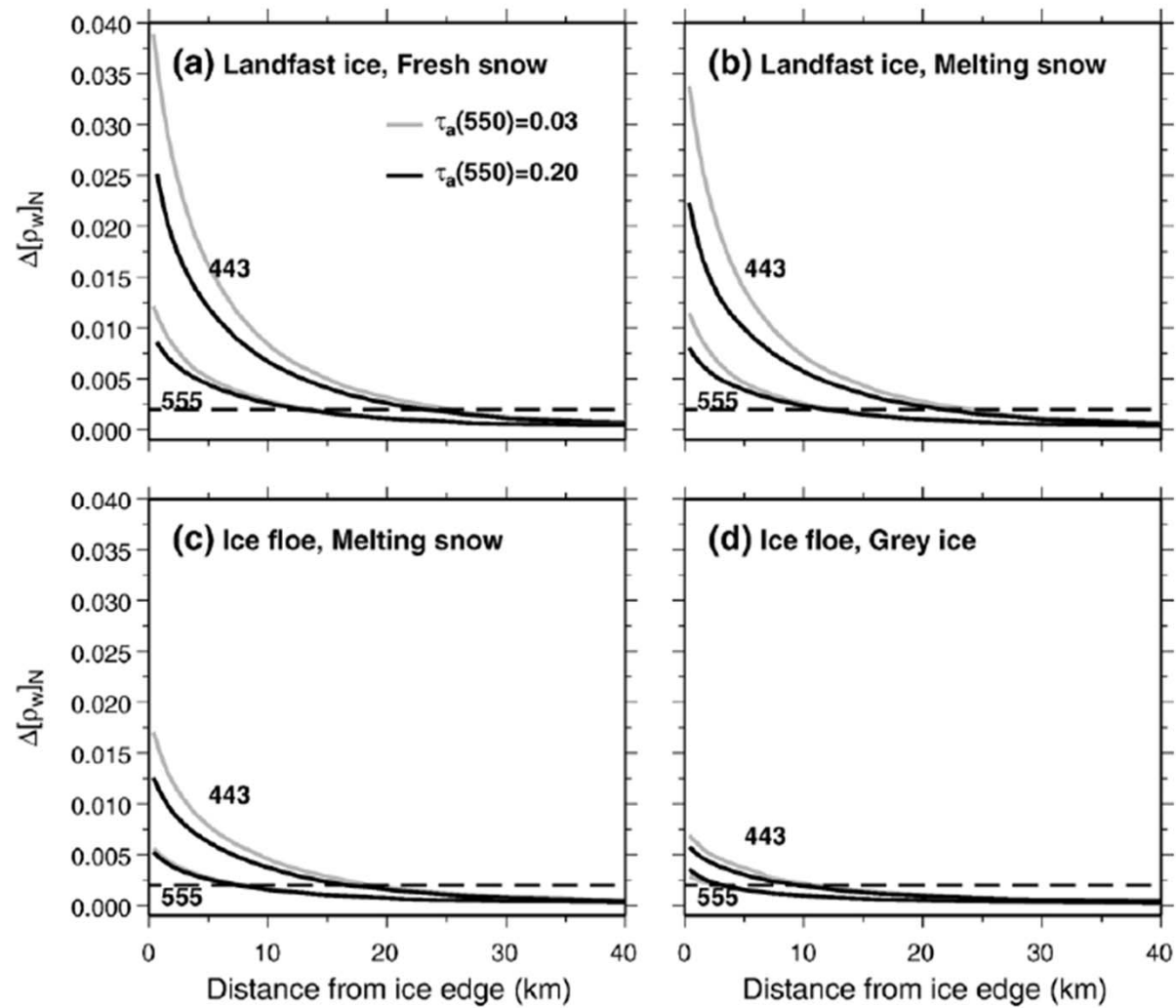
Normalized water-leaving
radiance at 443 nm



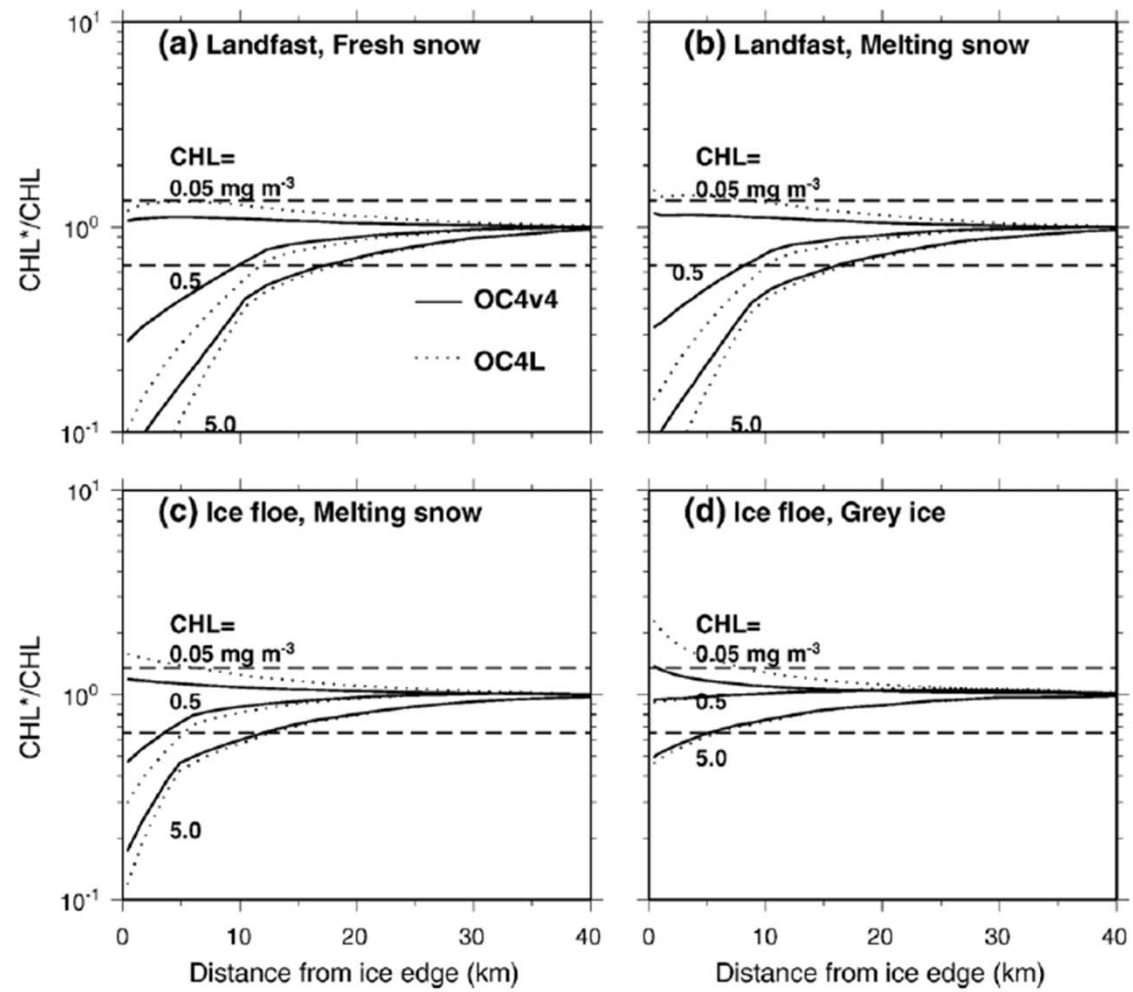
Bélanger et al. (2007)



Bélanger et al. (2007)



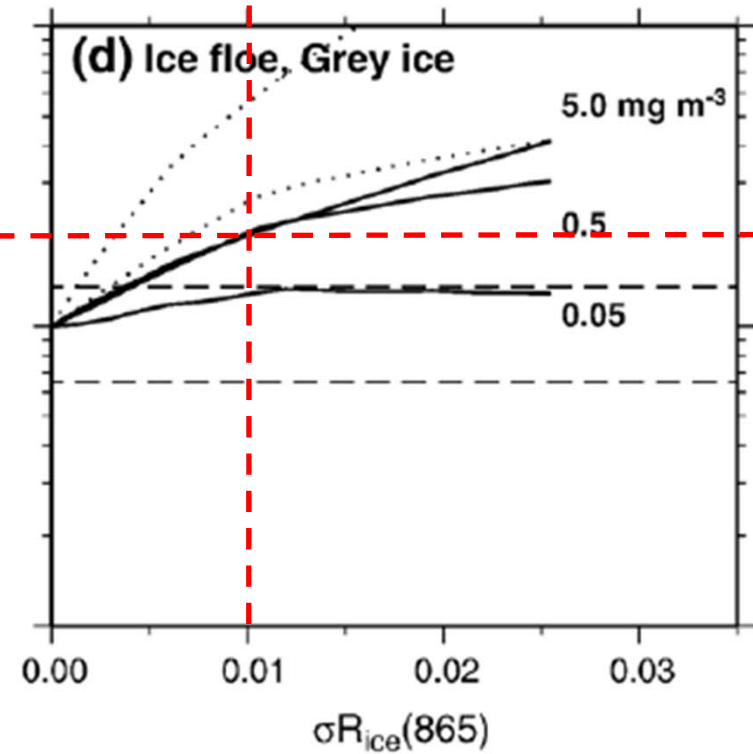
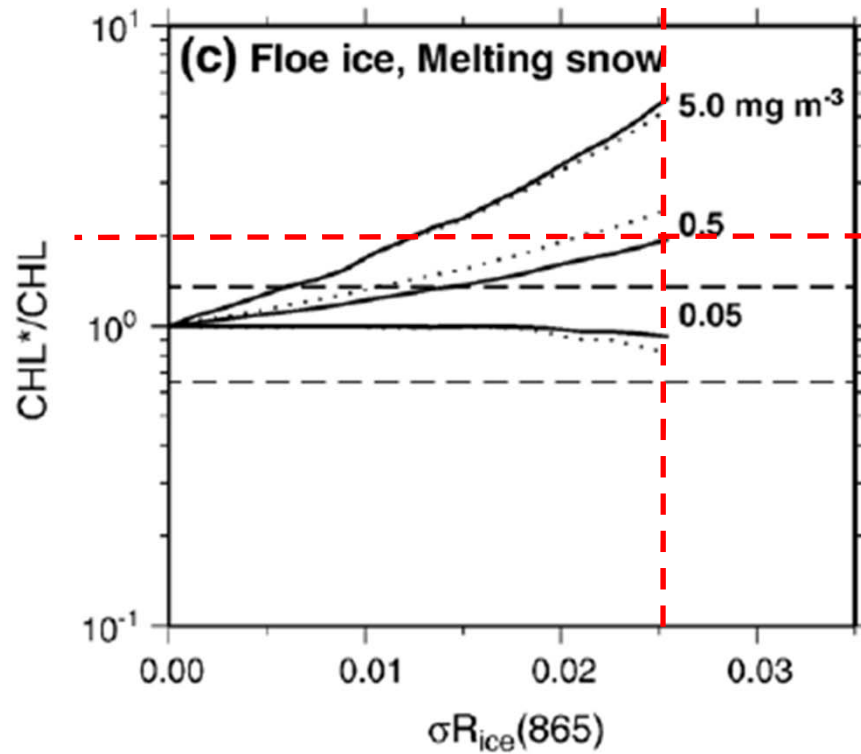
Bélanger et al. (2007)



Bélanger et al. (2007)

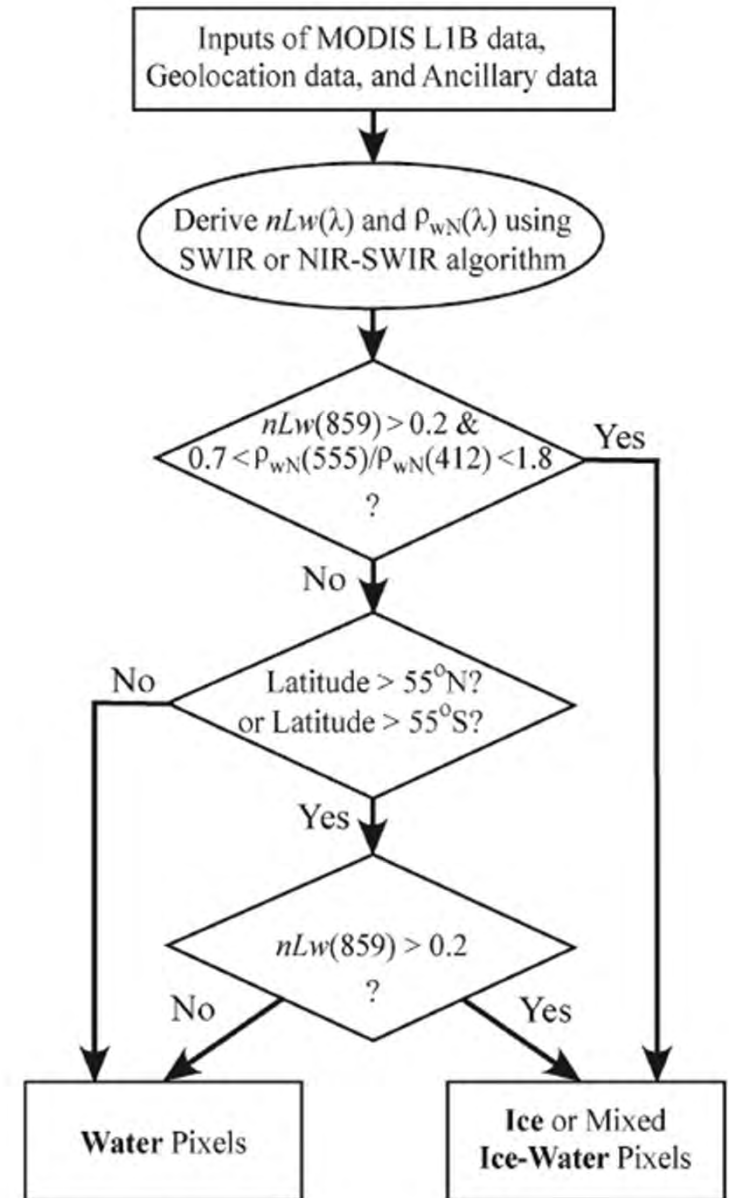
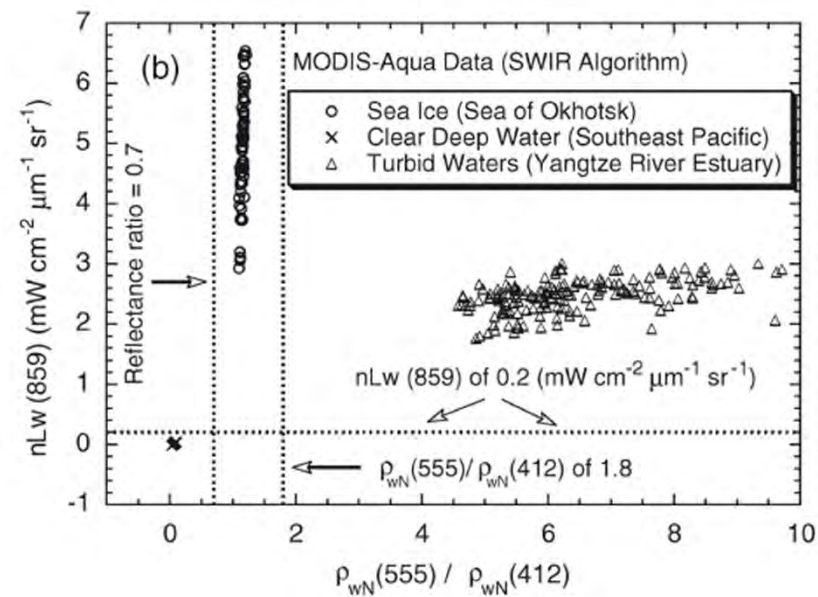
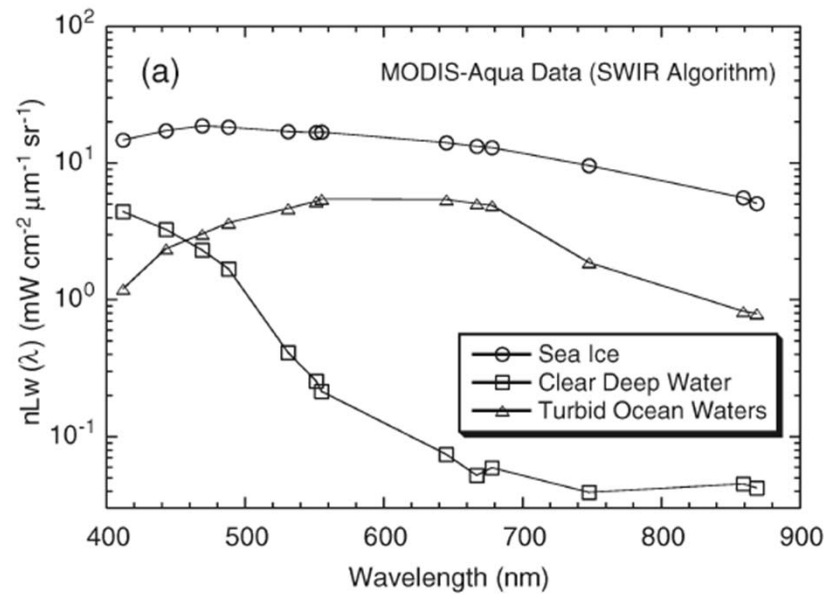
% ice cover = 7.5%

% ice cover = 10%

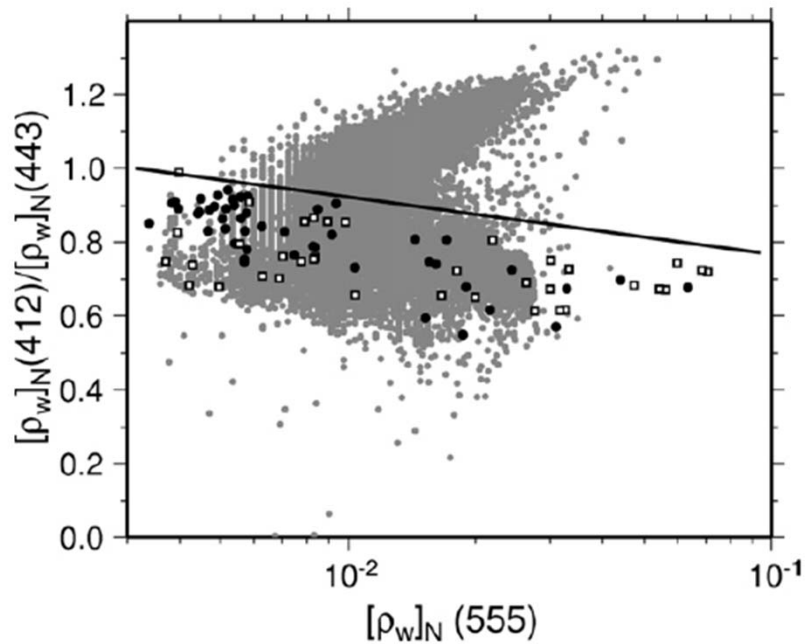


Bélanger et al. (2007)

How to flag ice pixels?

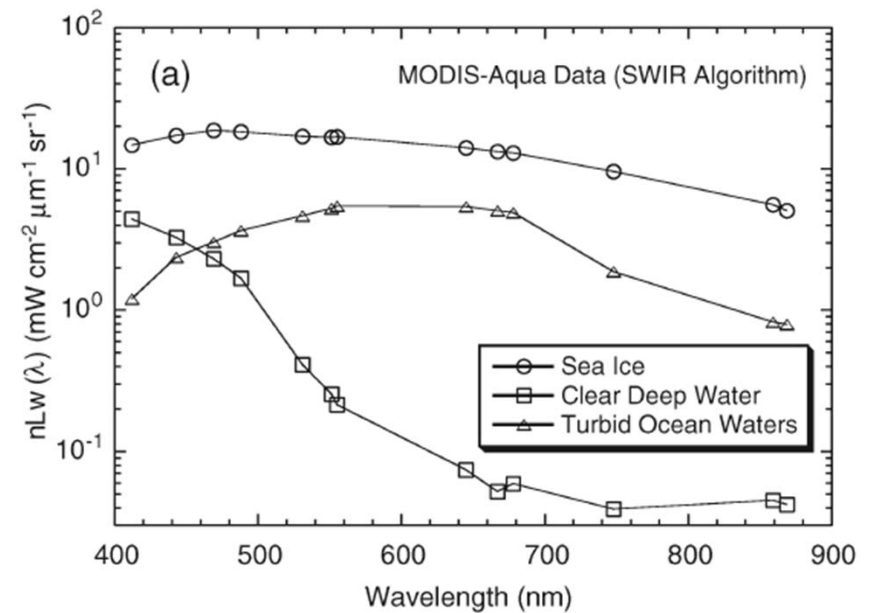


Wand and Shi (2009)

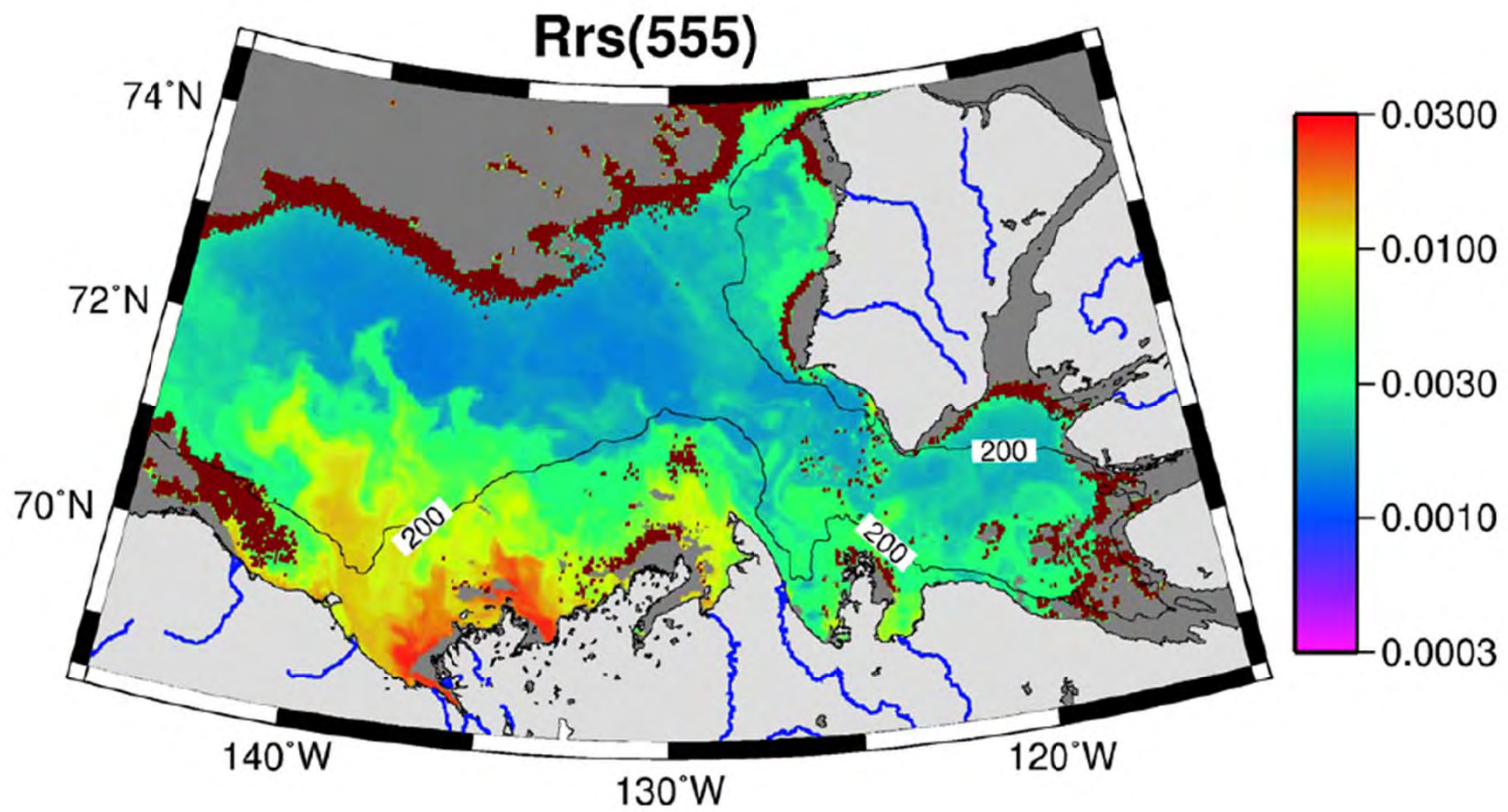


Bélanger et al. (2007)

How to flag pixels contaminated by the adjacency effect?



Wand and Shi (2009)



Bélanger et al. (2007)

Outline

1. Ocean colour remote sensing in polar seas

- 1.1. Ocean and sea ice in Arctic and Antarctic: relevant features
- 1.2. Seawater optical properties
- 1.3. Retrieval of ocean properties from ocean colour
 - 1.3.1. Atmospheric corrections
 - 1.3.2. Contamination of the signal by sea ice
 - 1.3.3. Retrieval of IOPs and AOPs, and biogeochemically relevant variables
- 1.4. Availability of data as favoured by polar orbits and limited by elevated cloudiness

2. Primary production estimates from OC in polar seas

- 2.1. PP model and validation: an example
- 2.2. Results from PP models

1. Ocean colour remote sensing in polar seas

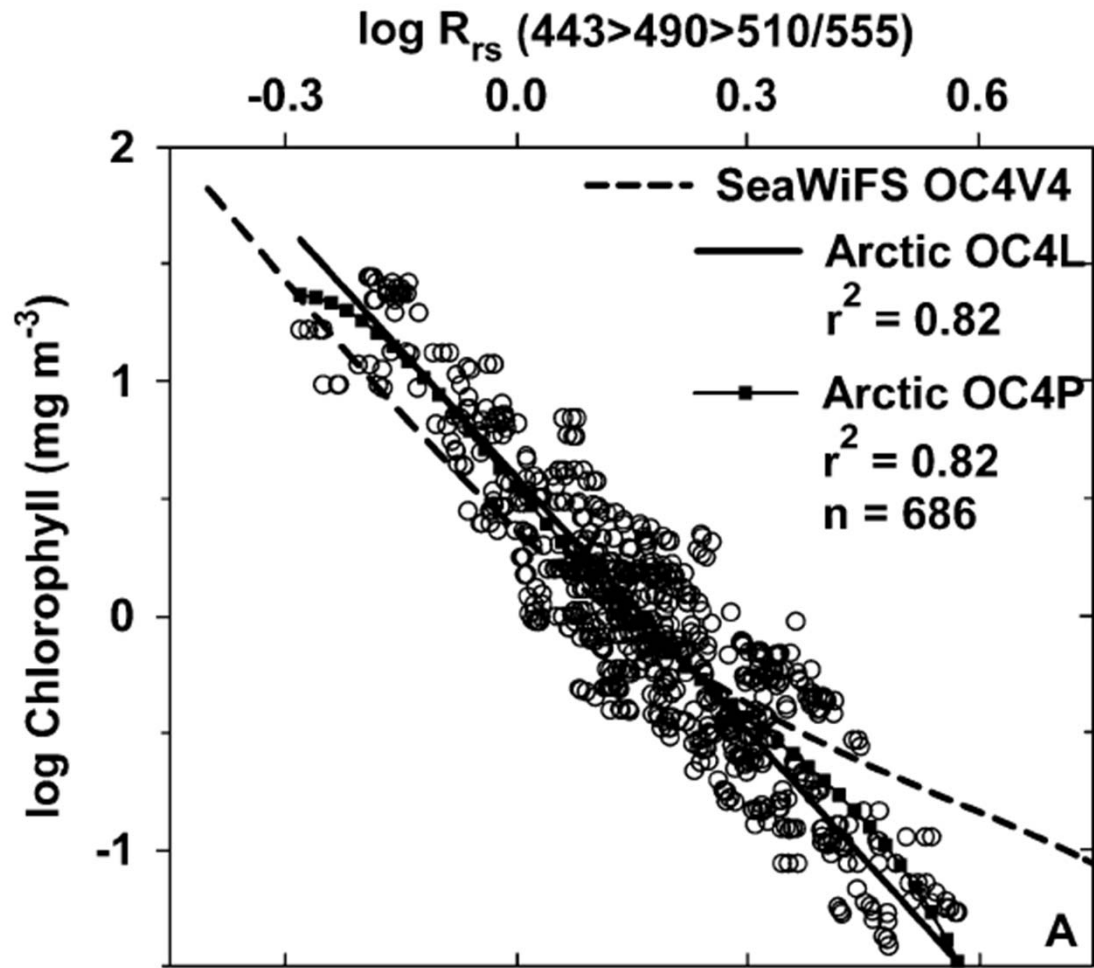
1.3. Retrieval of ocean properties from ocean colour

1.3.3. Retrieval of IOPs and AOPs, and biogeochemically relevant variables

Cota et al. 2004

Data from:

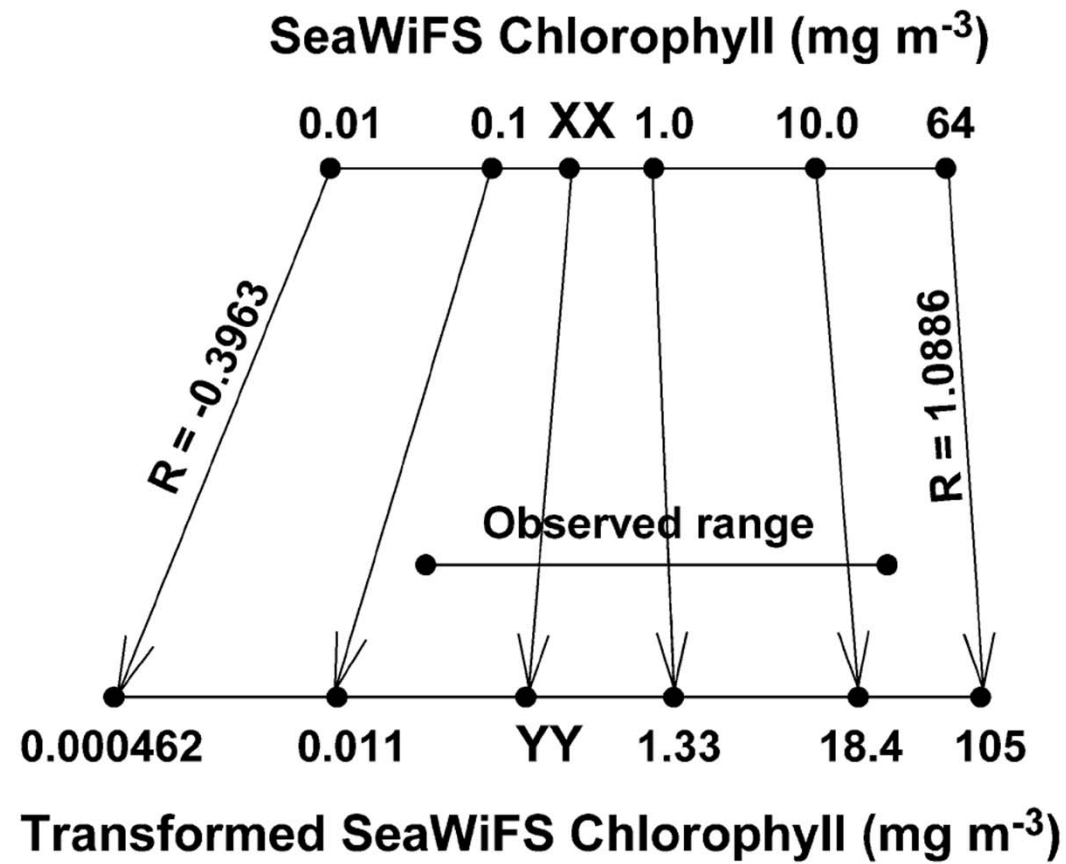
- Labrador Sea
- Chukchi Sea
- Beaufort Sea



Arctic OC4L algorithm

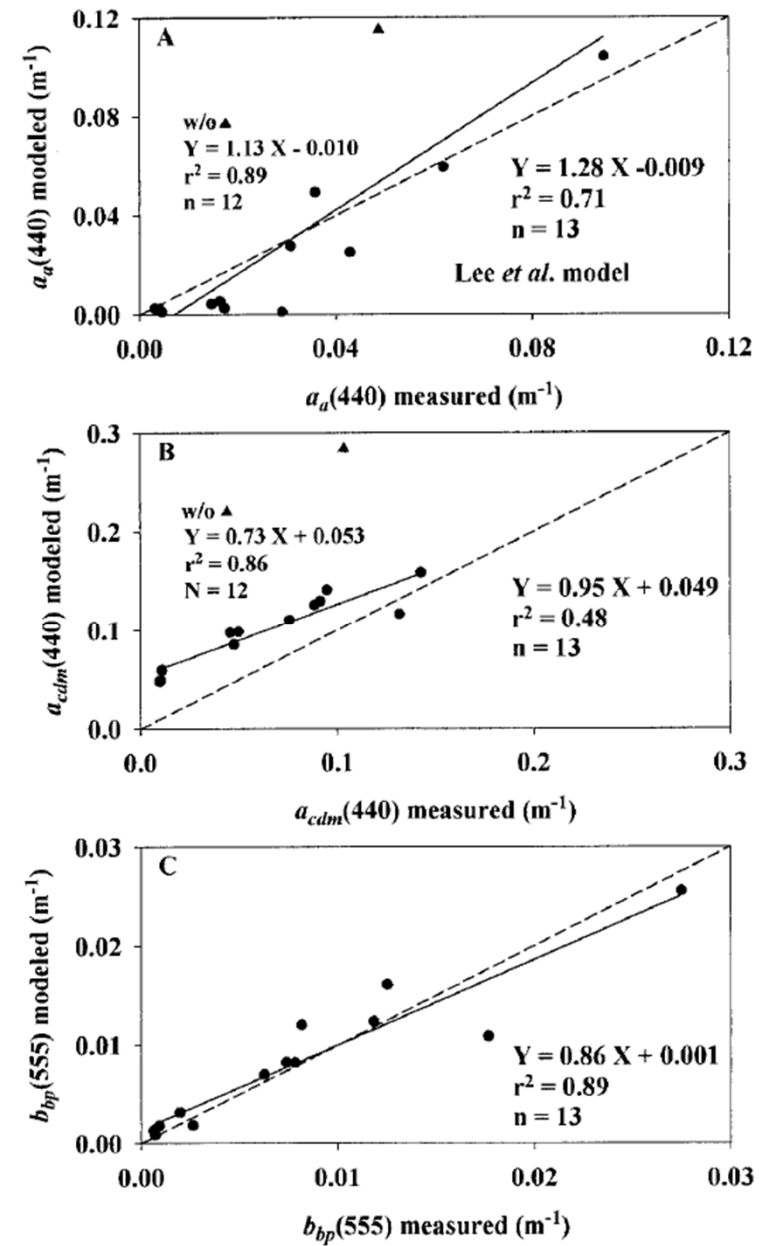
$$\text{Chl (Arc)} = 10^{(a_1 + b_1 R)}$$

$$R = \log(R_{rs443>490>510}/R_{rs555})$$

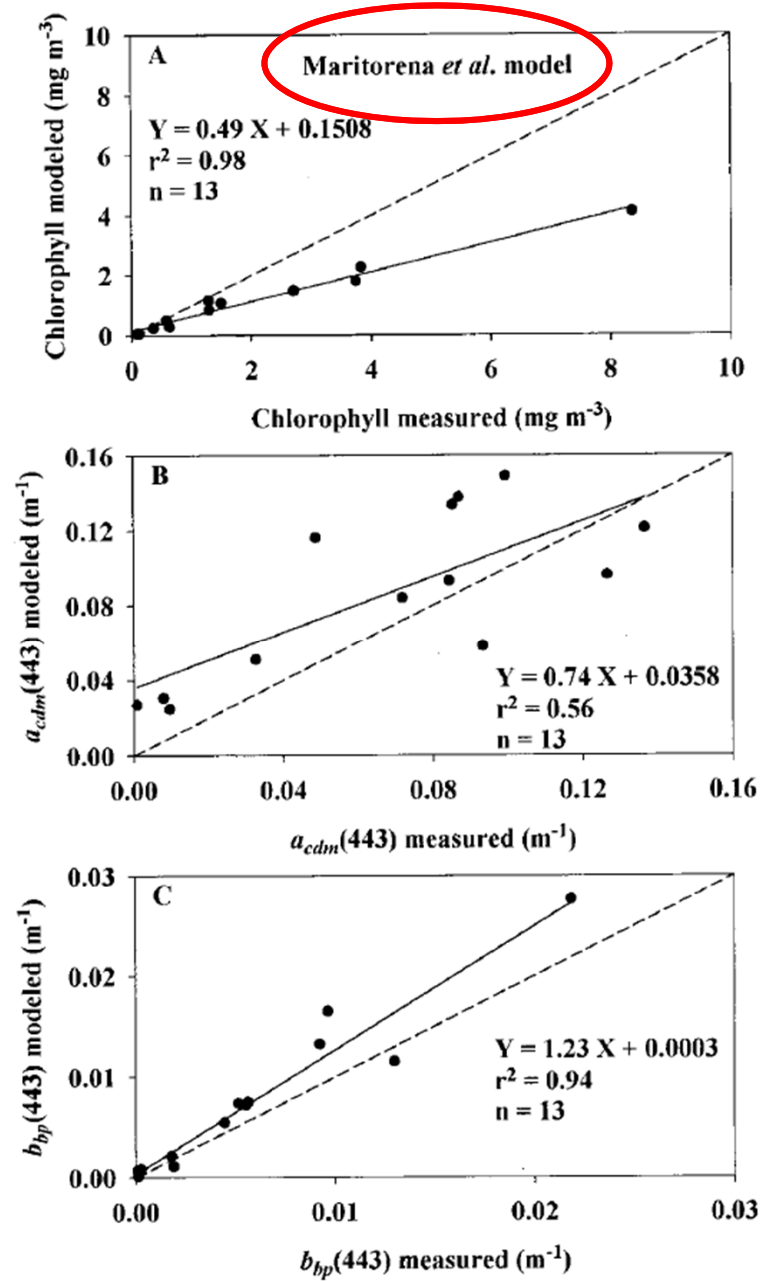


Wang et al. (2003)

Tuned Lee et al. (2001)

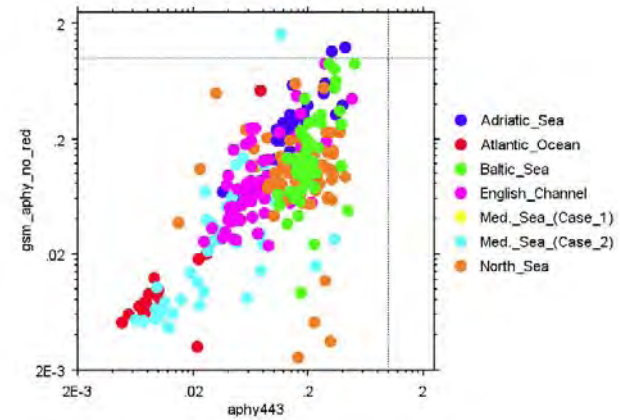
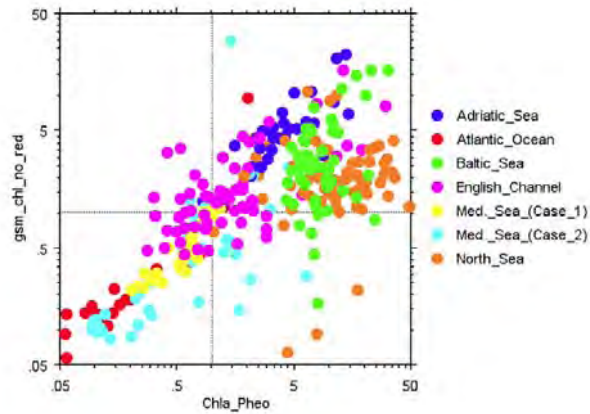


Wang et al. (2003)

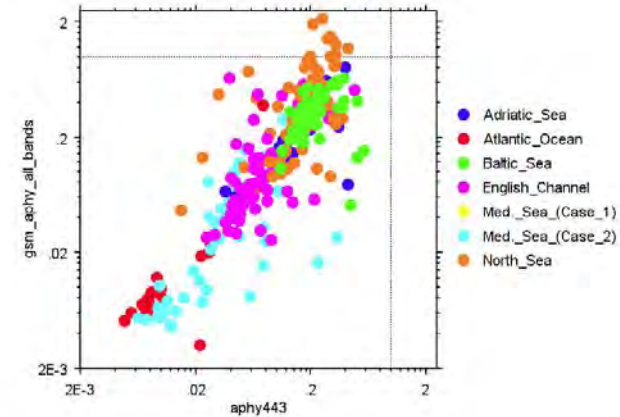
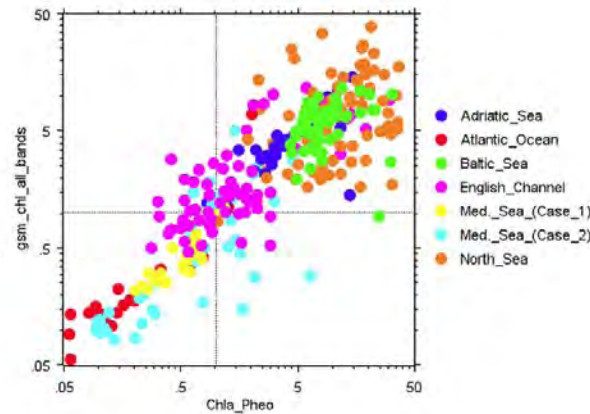


GSM validation using Coastlooc data from Babin (2003)

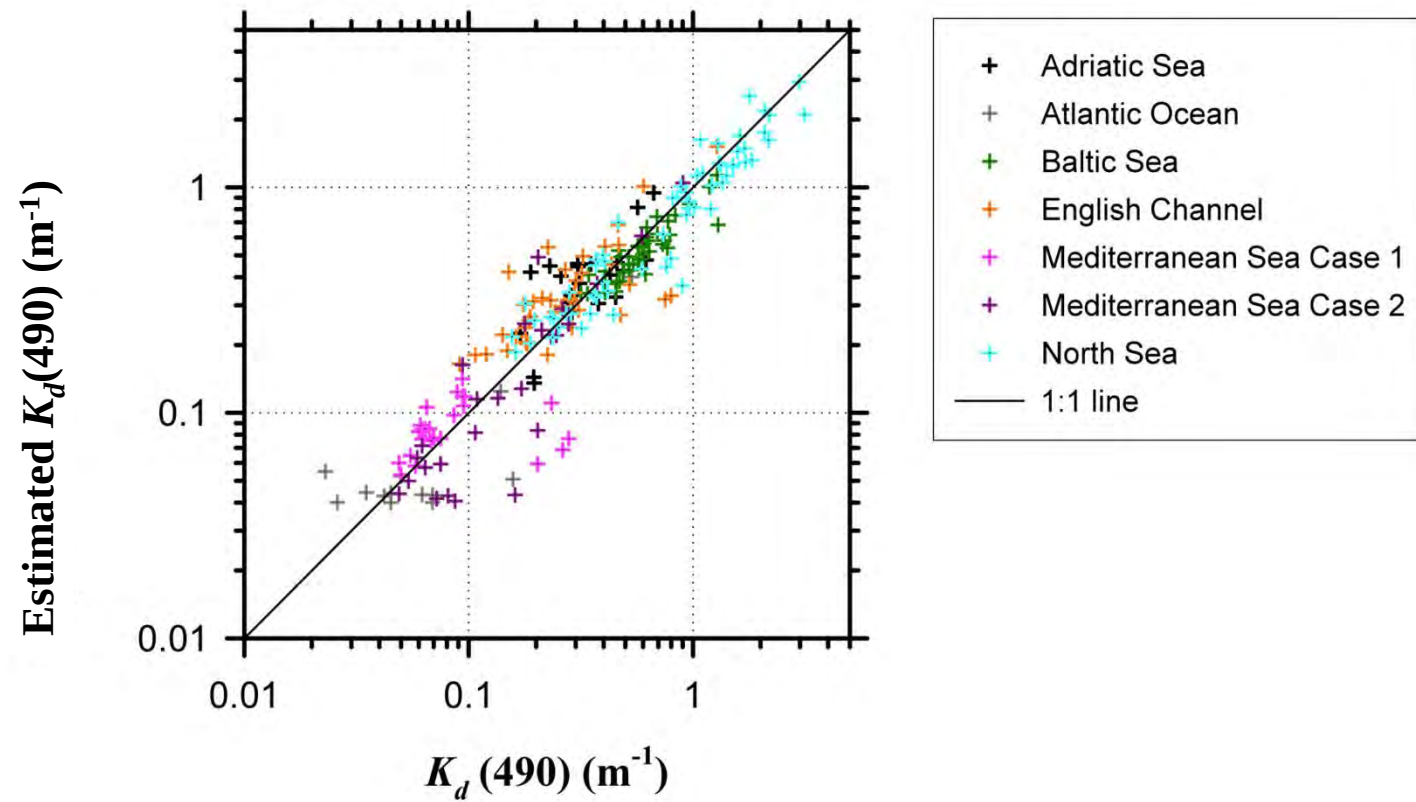
Blue & green channels



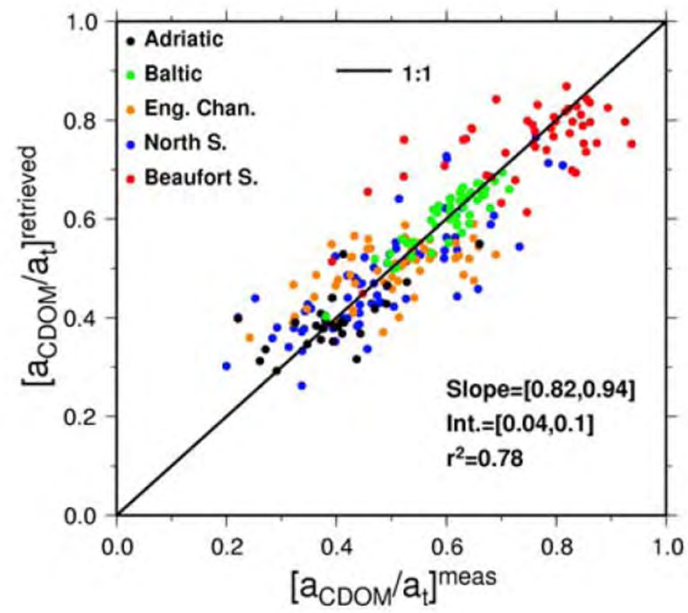
Blue, green & red channels



QAA validation in Case 2 waters

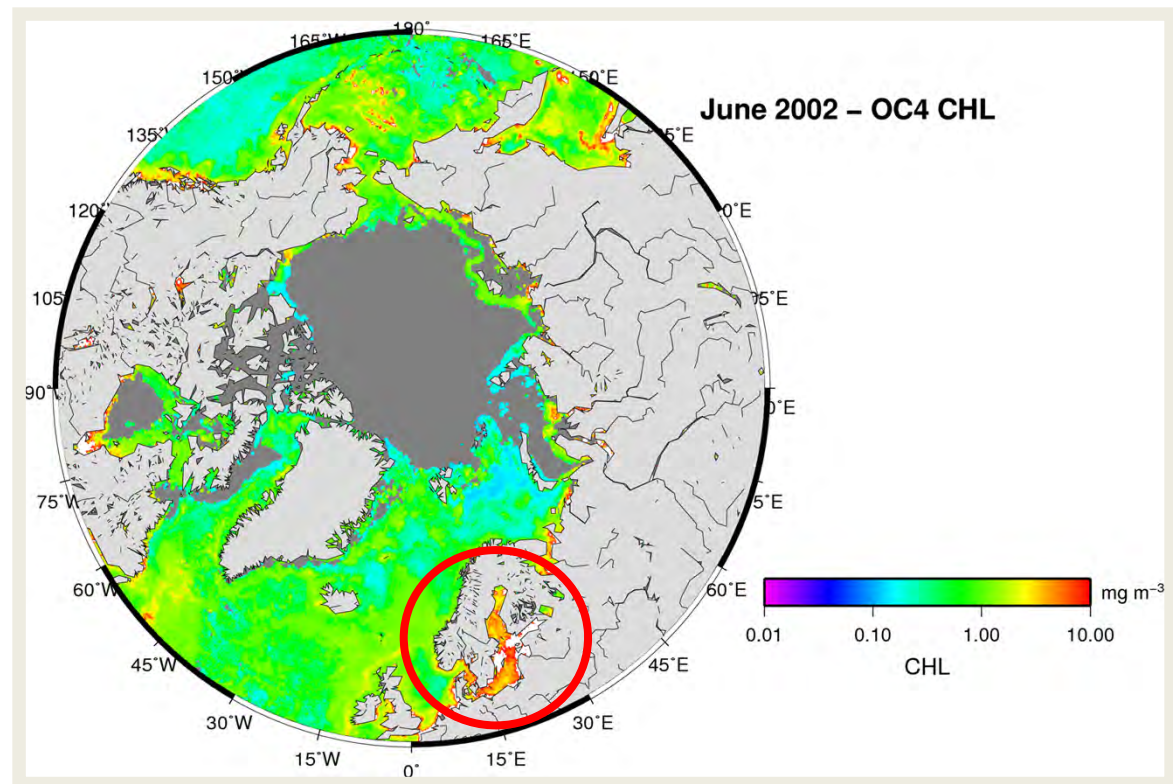


Doron et al. (2007)

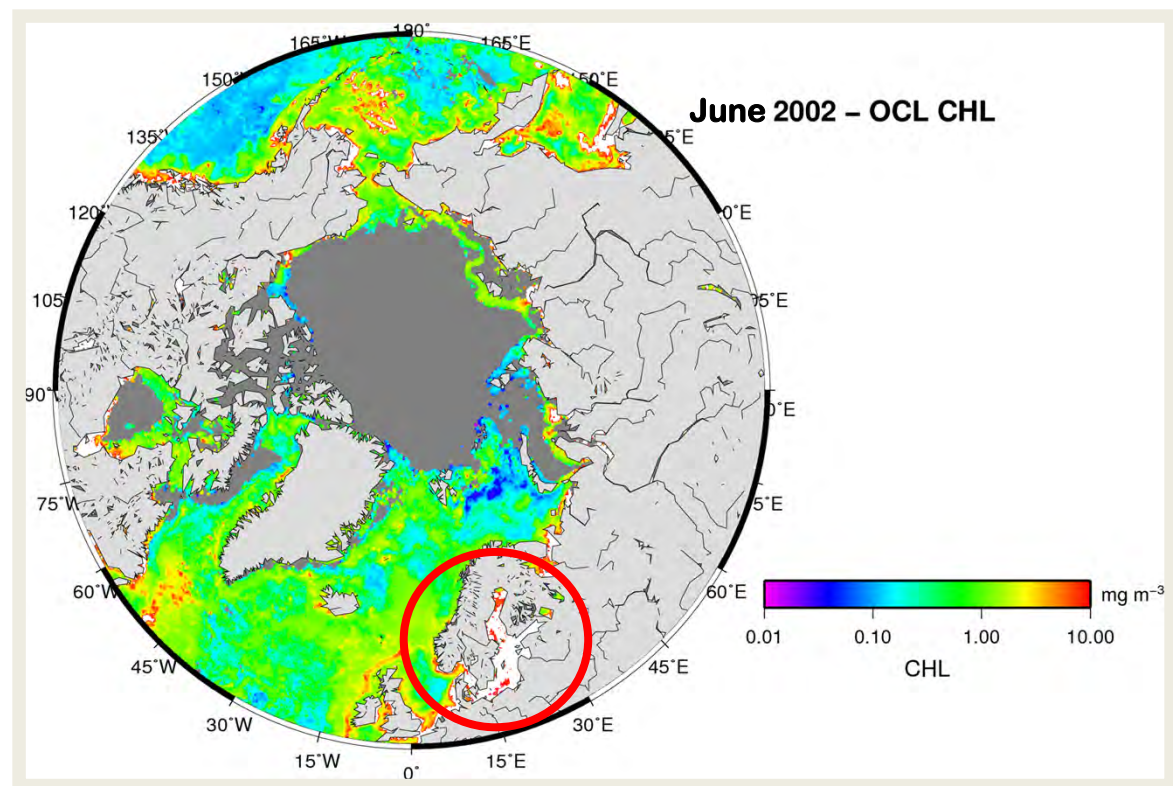


Bélanger et al. (2008)

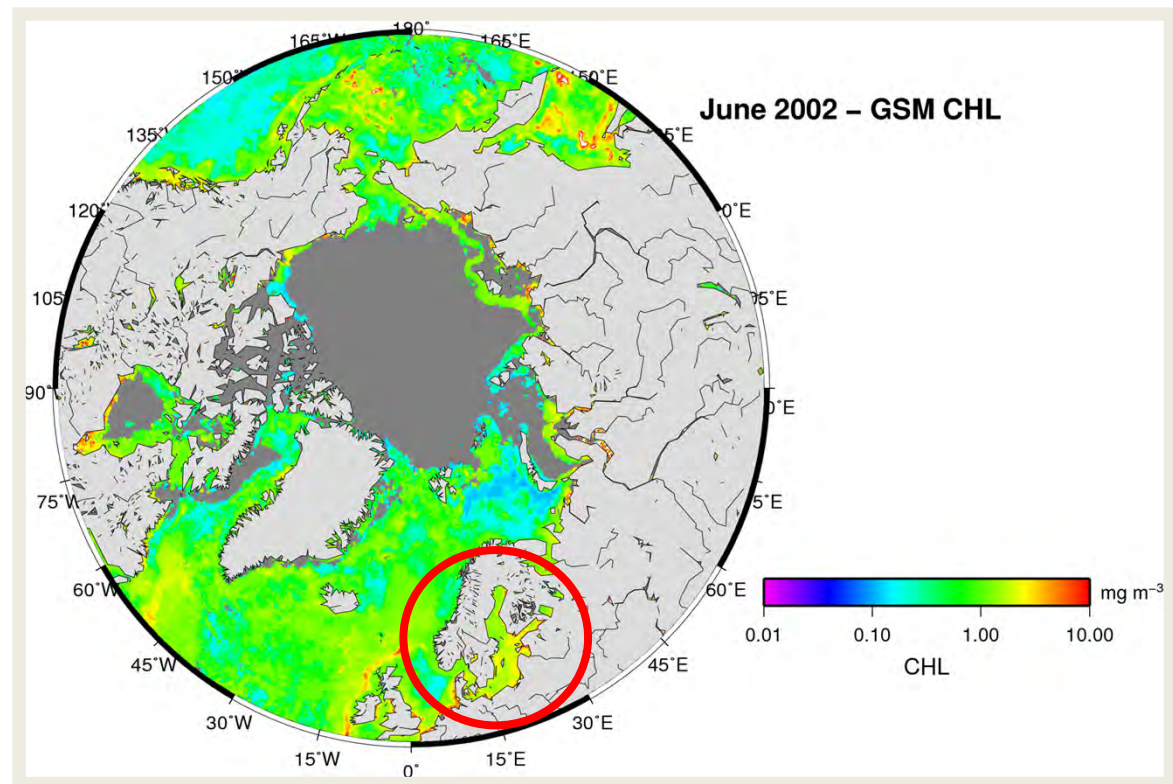
Standard ocean color chl products are not reliable in the Arctic Ocean



Standard ocean color chl products are not reliable in the Arctic Ocean



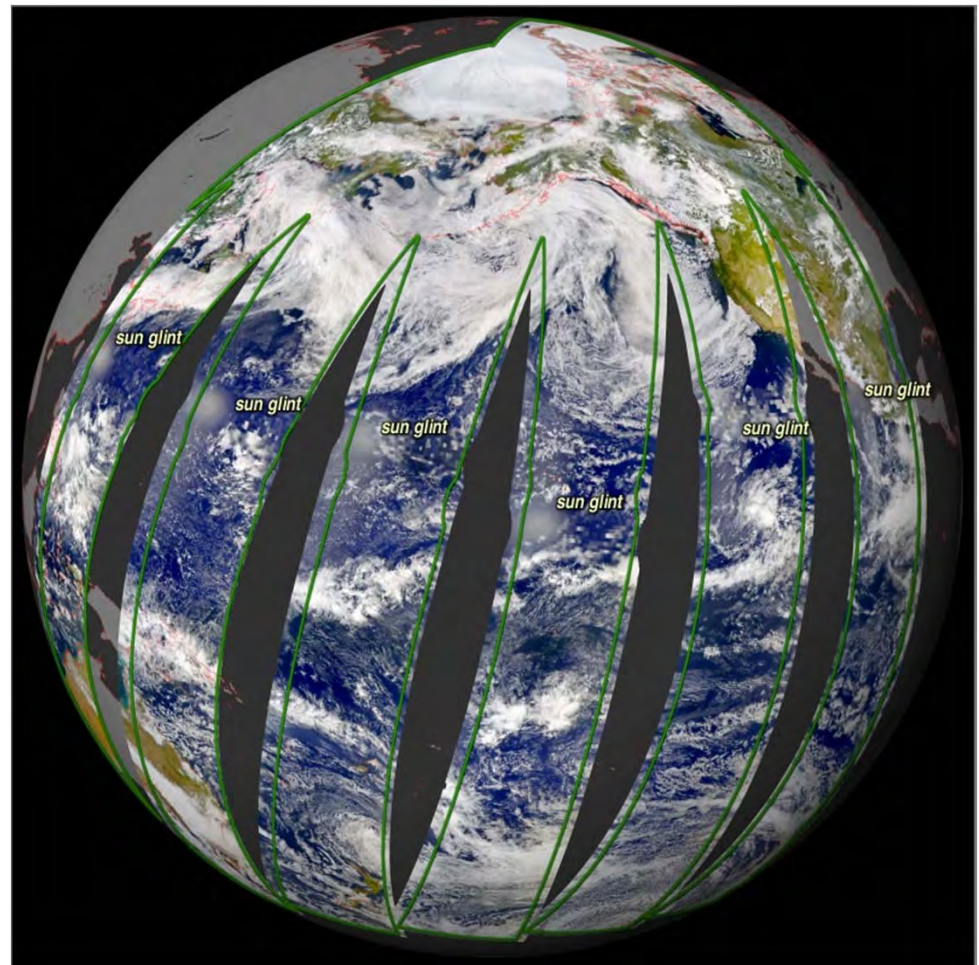
Standard ocean color chl products are not reliable in the Arctic Ocean



1. Ocean colour remote sensing in polar seas

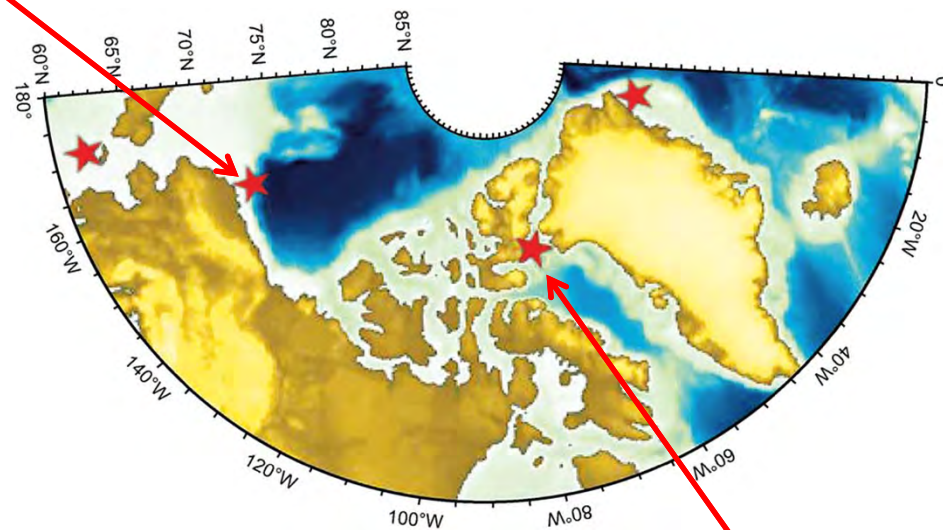
1.4. Availability of data as favoured by polar orbits
and limited by elevated cloudiness

Polar orbiting ocean color sensors provide data several several times a day



Cloud cover

BFL (August-September)

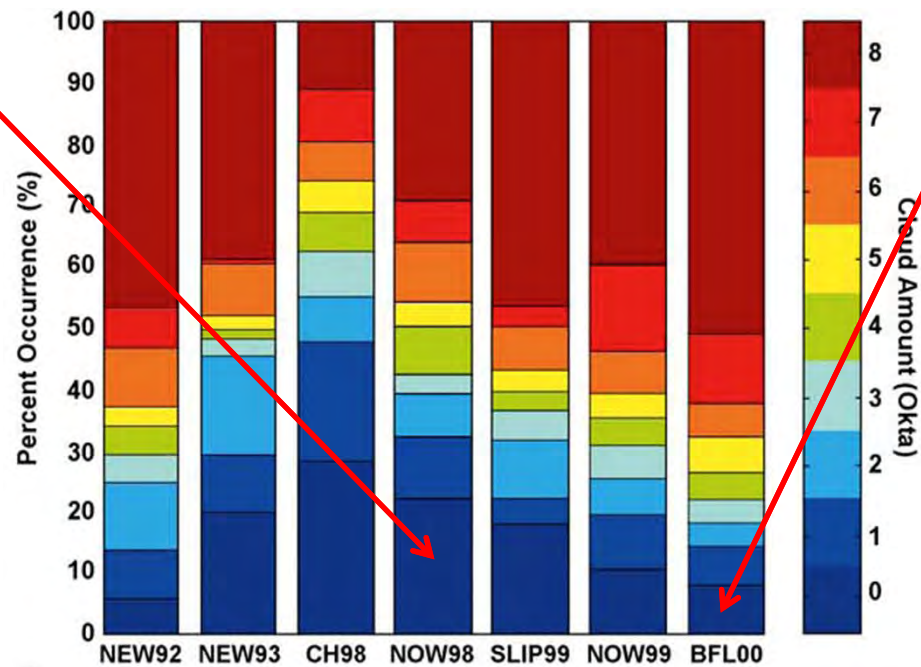


Key et al. (2004)

NOW98 (April-July)

NOW98 (April-July)

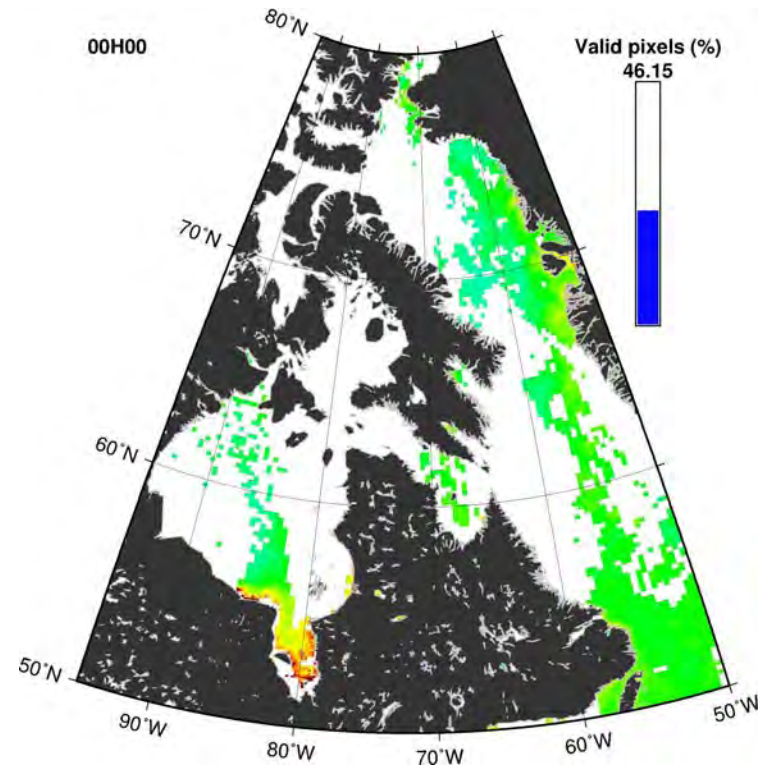
BFL (August-September)



Key et al. (2004)

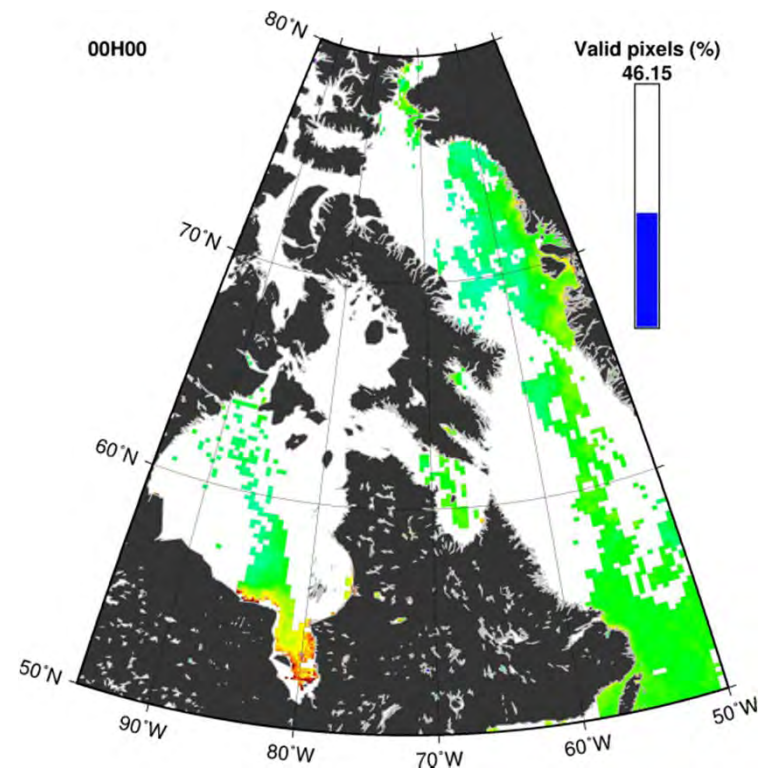
High-frequency observations may help

MERIS L3 Chlorophyll-a Concentration (mg m^{-3})

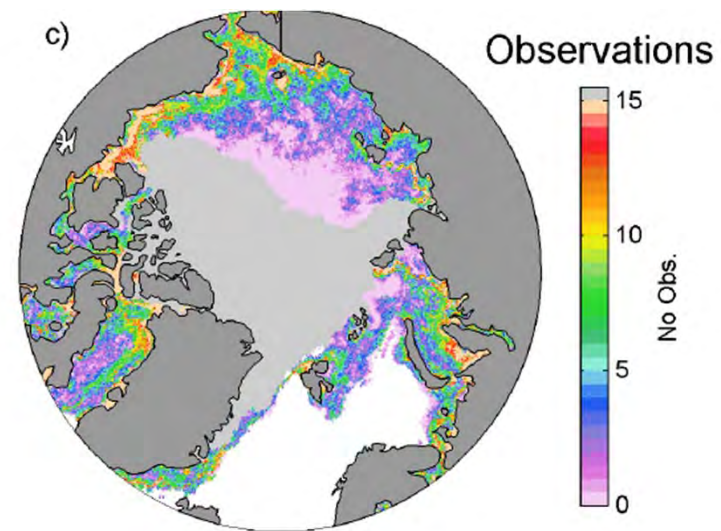
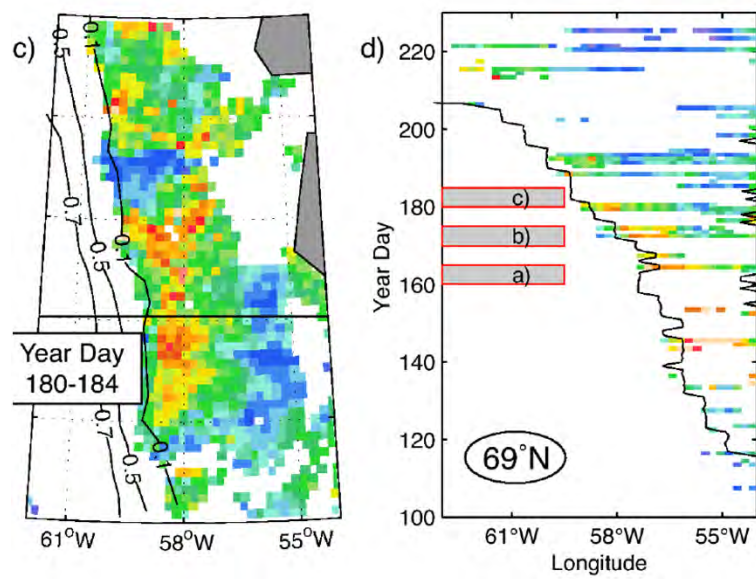


Cloud cover derived using METEOSAT GEMS02 Satellite (15 min resolution)

MERIS L3 Chlorophyll-a Concentration (mg m^{-3})



Cloud cover derived using METEOSAT GEMS02 Satellite (15 min resolution)

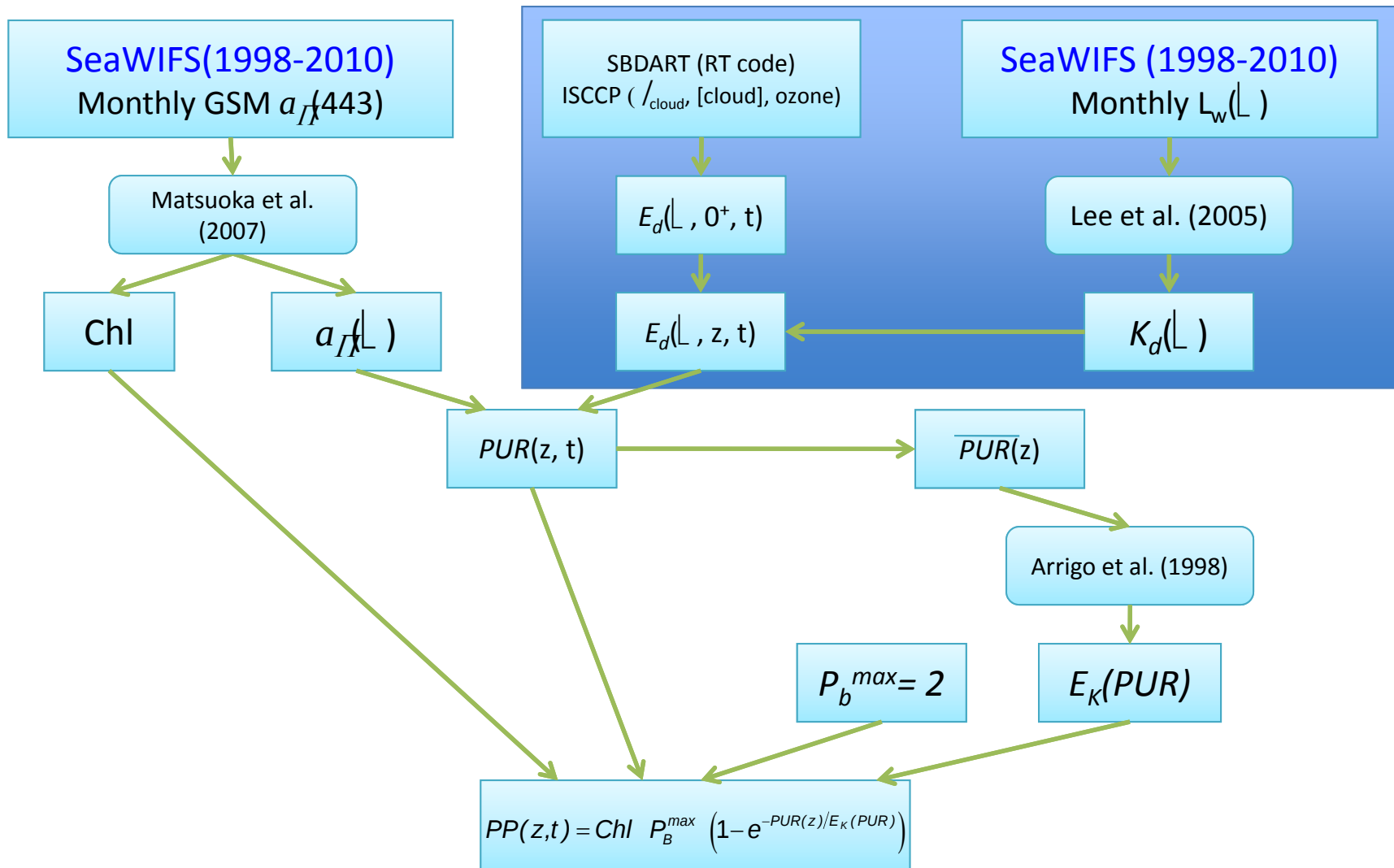


Perrette et al. (2011)

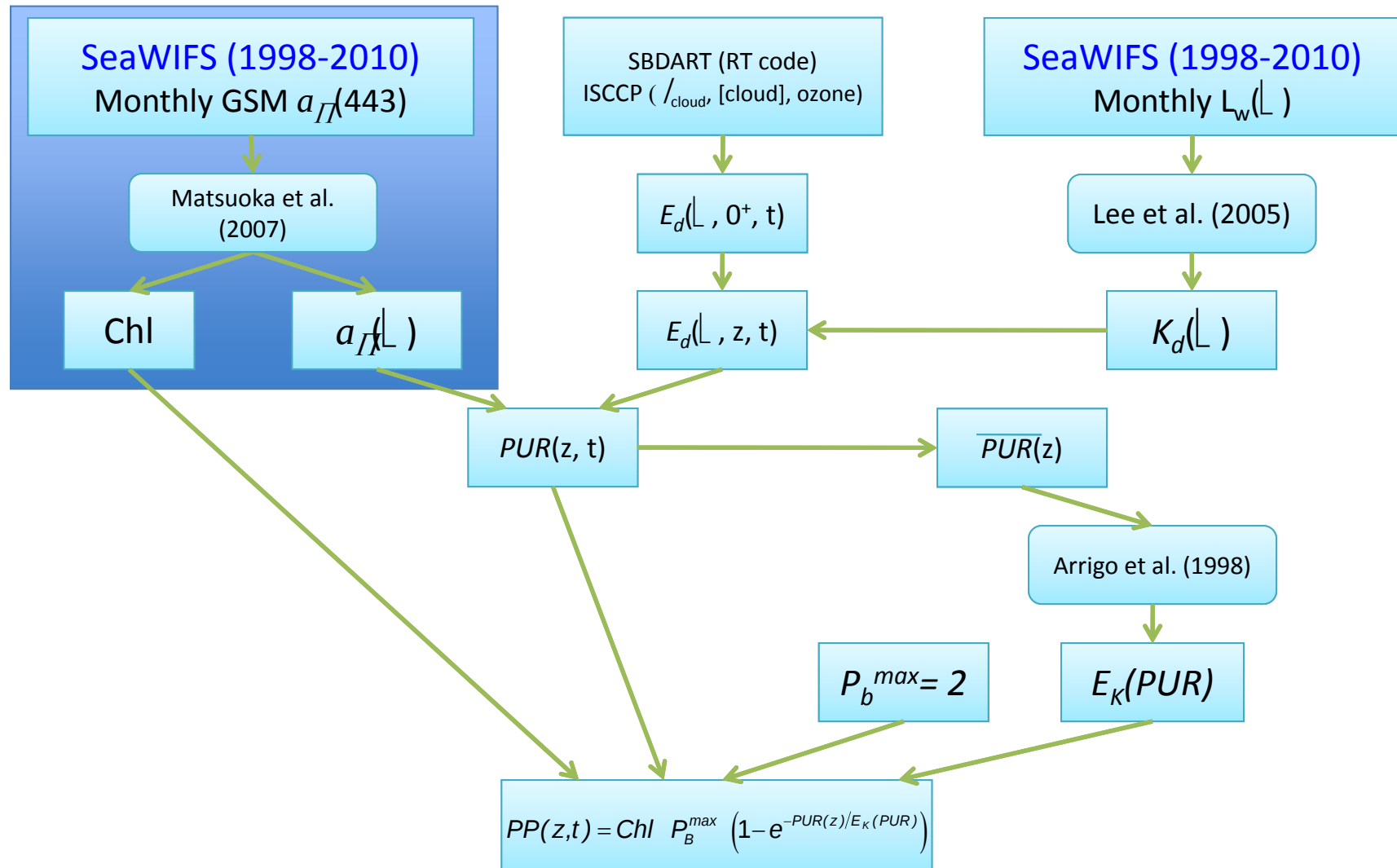
2. Primary production estimates from OC in polar seas

2.1. PP model and validation: an example

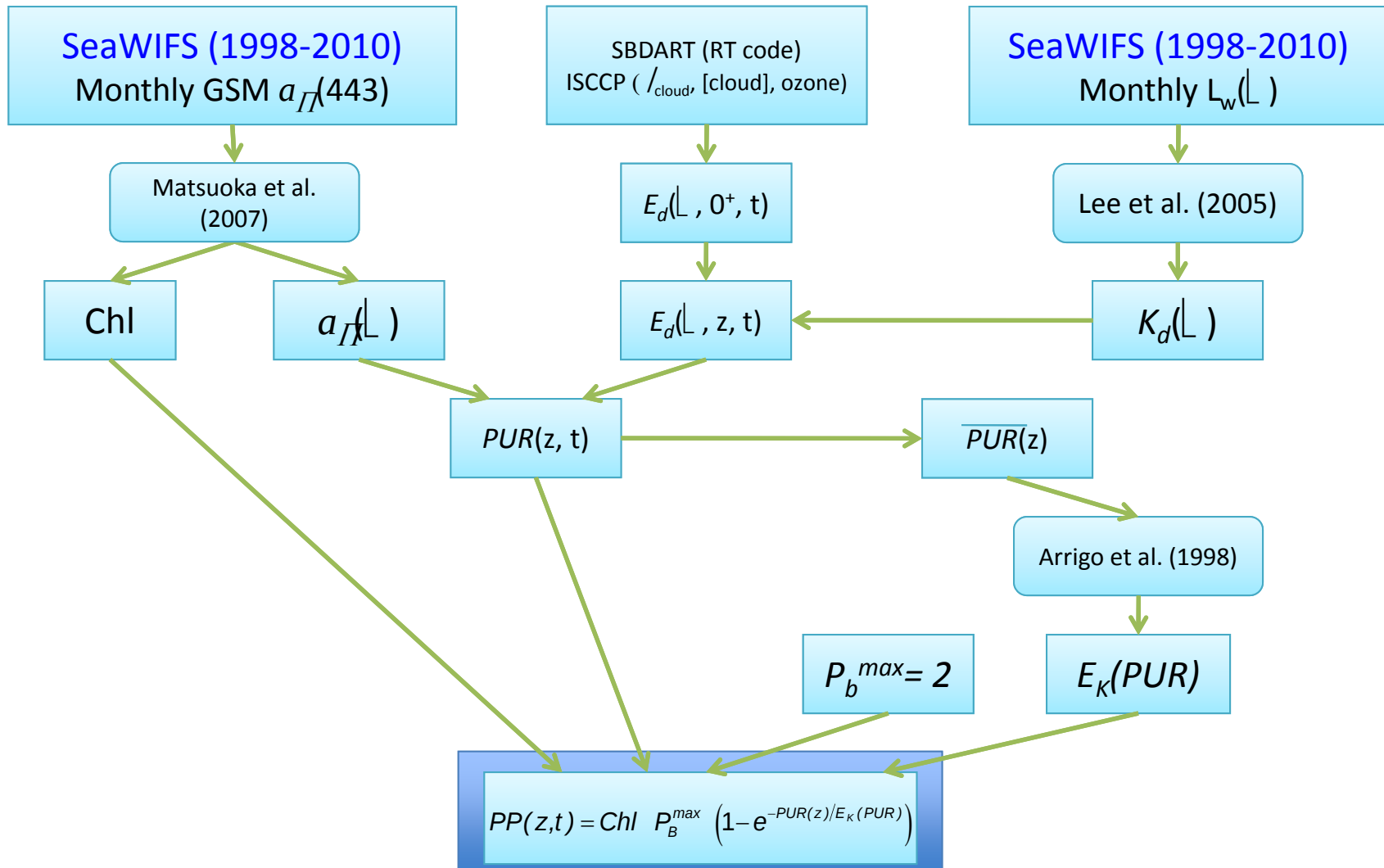
Methods – Primary production model



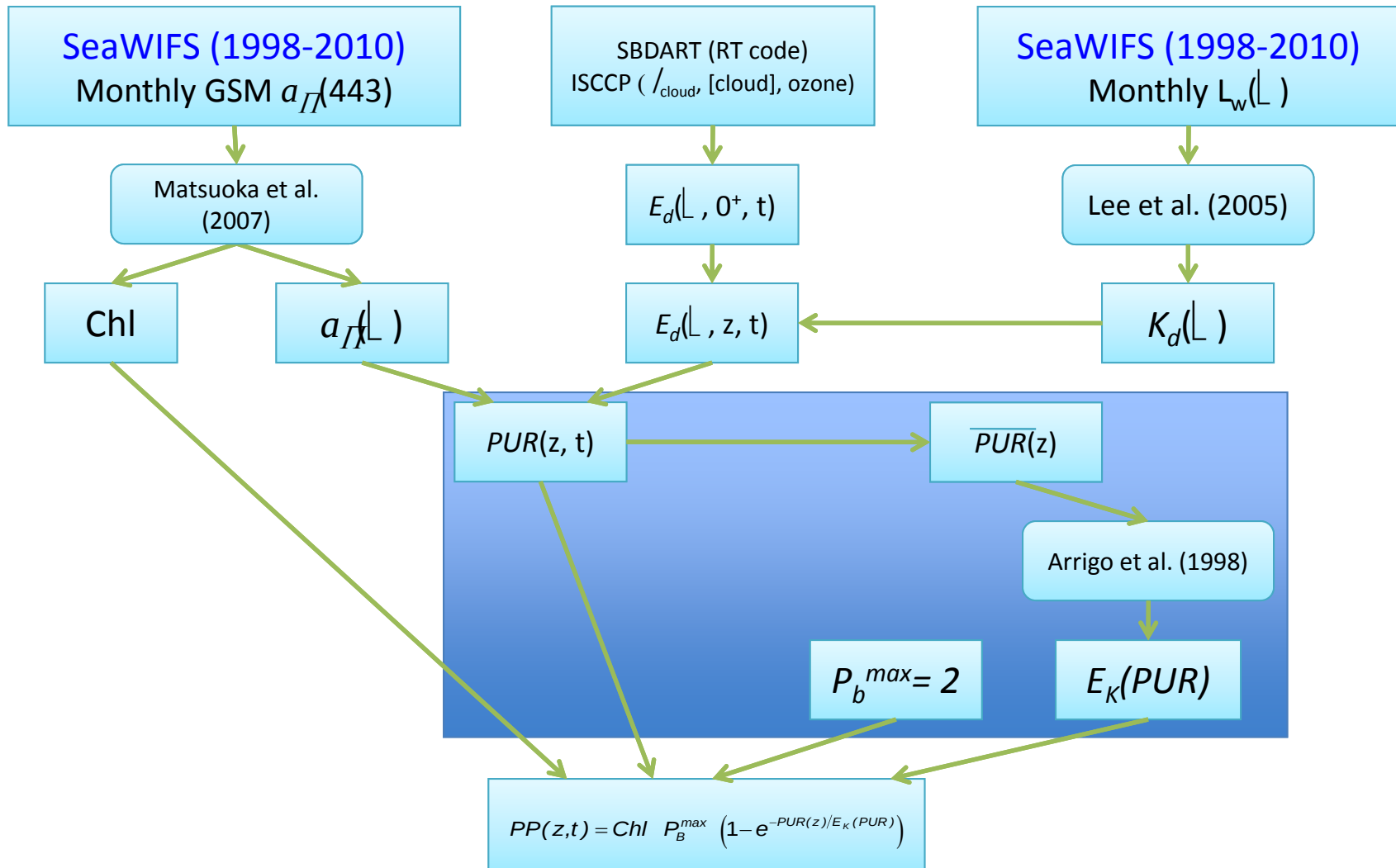
Methods – Primary production model



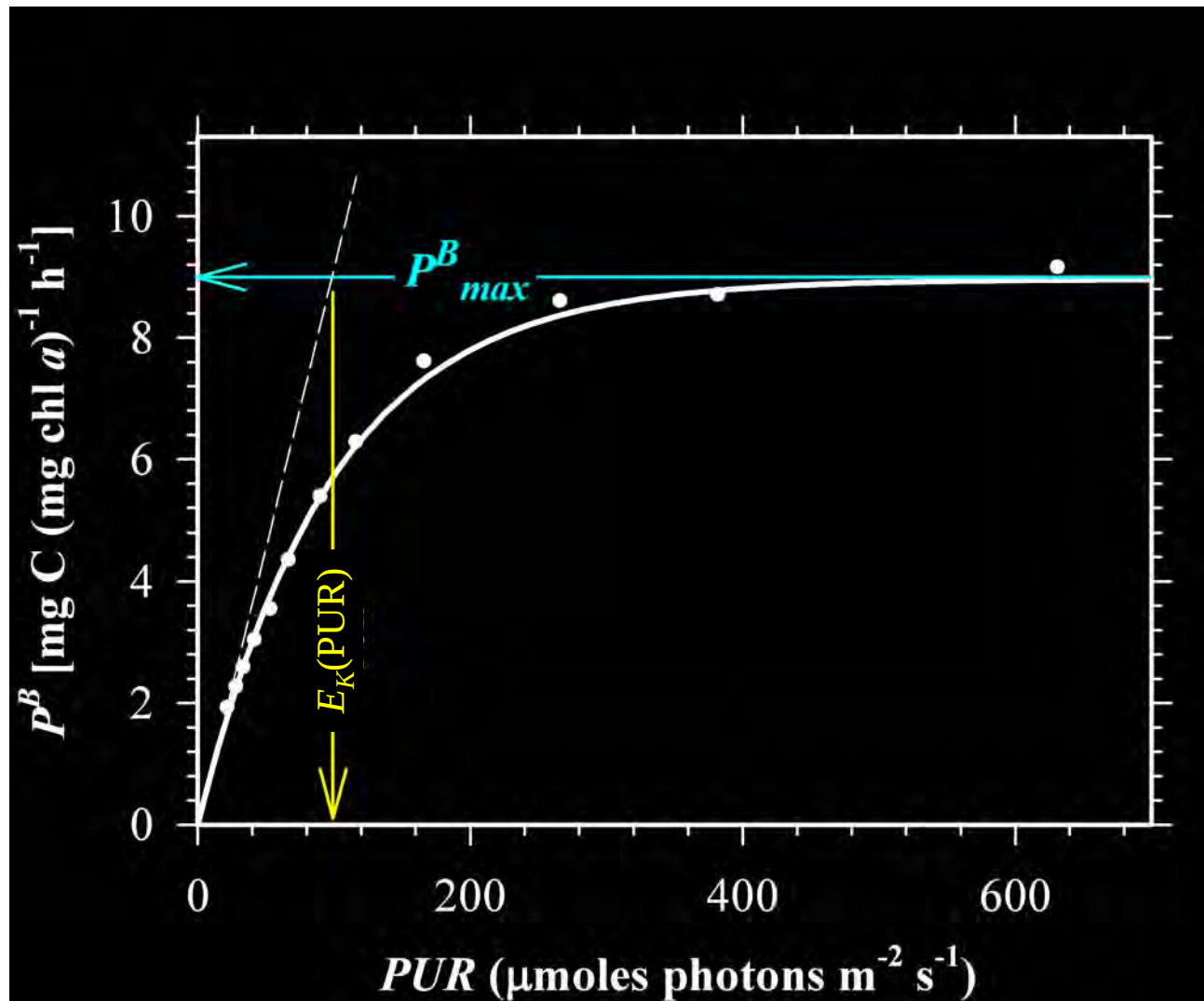
Methods – Primary production model



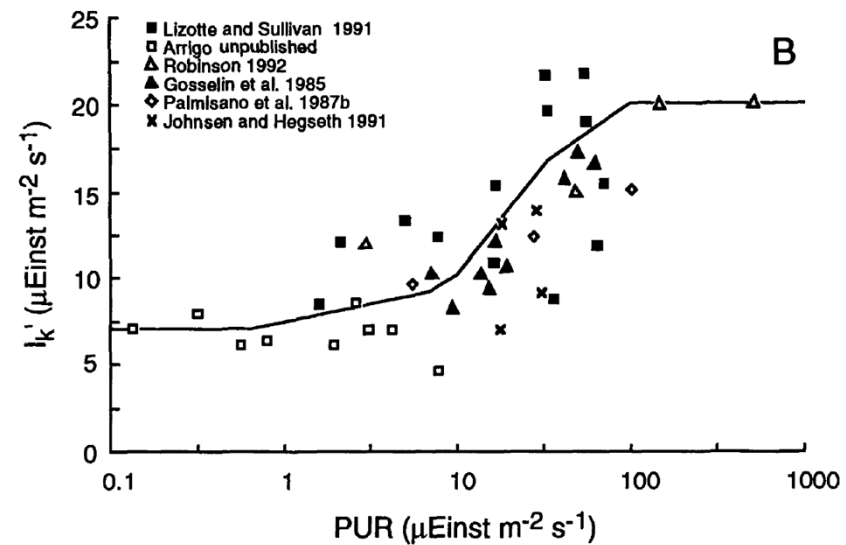
Methods – Primary production model



$$PP(z,t) = Chl \ P_B^{max} \left(1 - e^{-PUR(z)/E_K(PUR)}\right)$$



Photosynthetic parameters



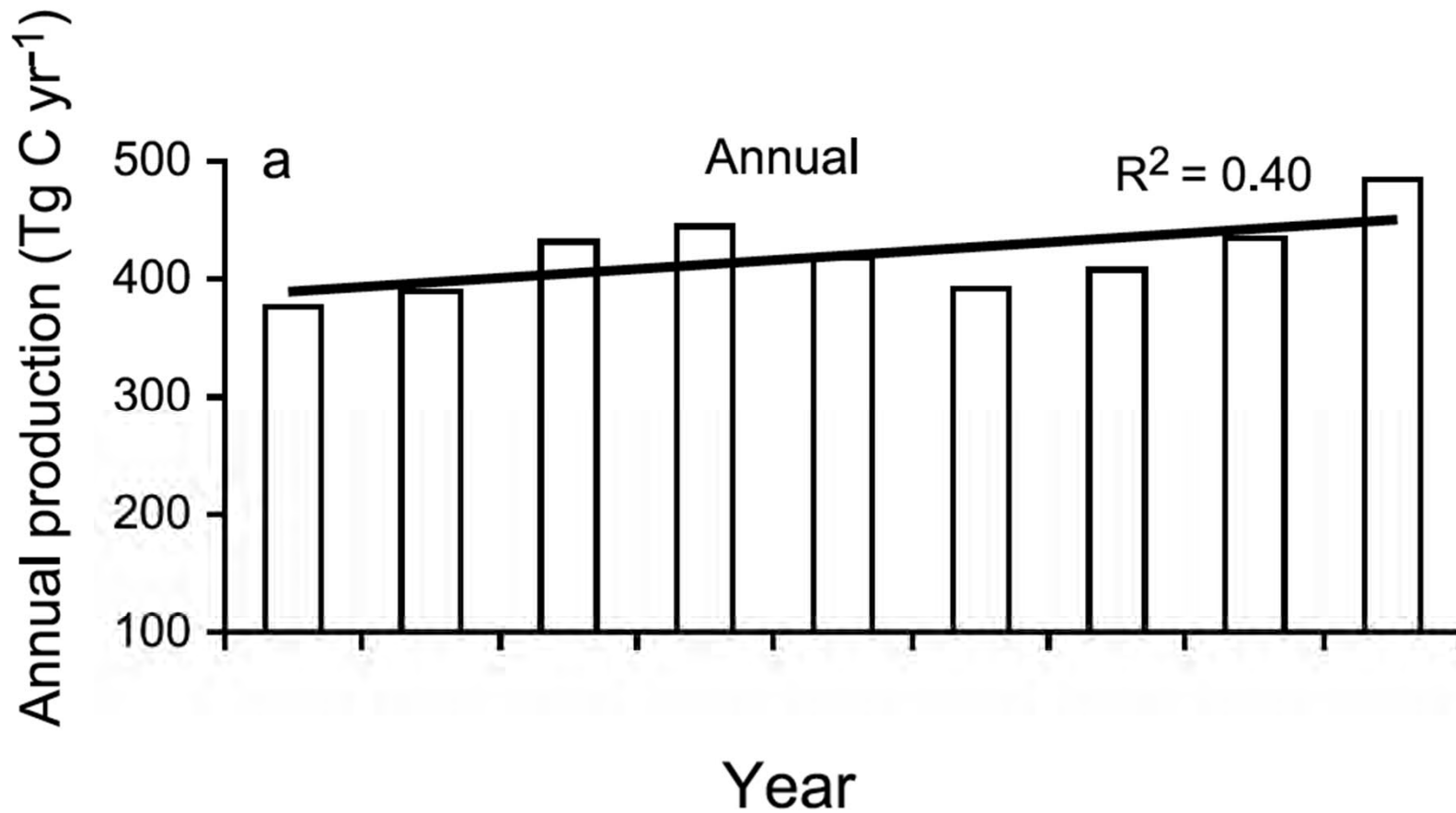
Arrigo and Sullivan (1994)

2. Primary production estimates from OC in polar seas

2.2. Results from this and other PP models

Arctic Primary Production

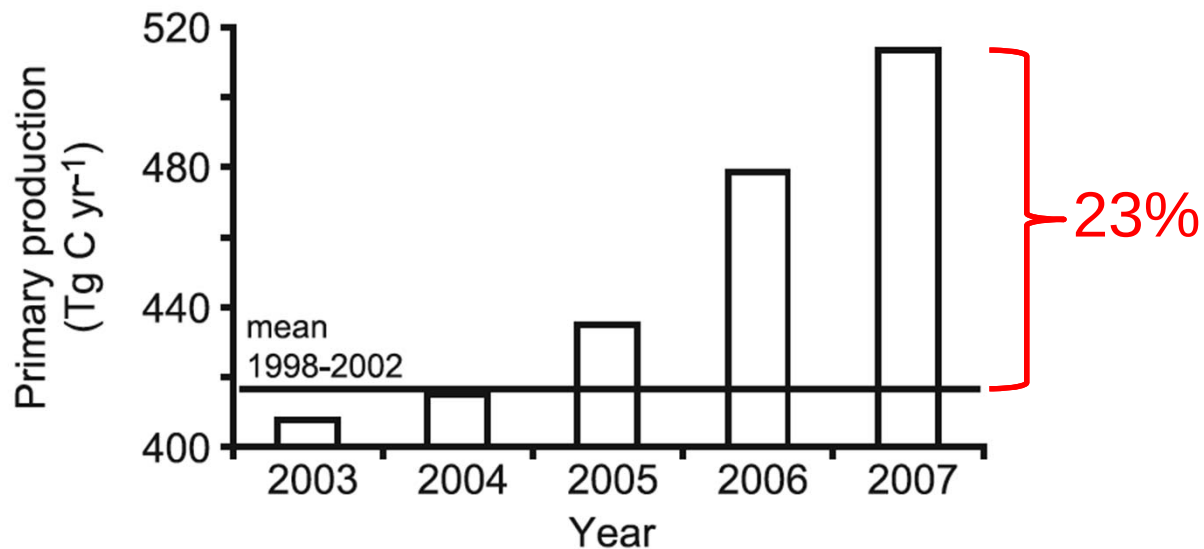
Pabi et al. 2008



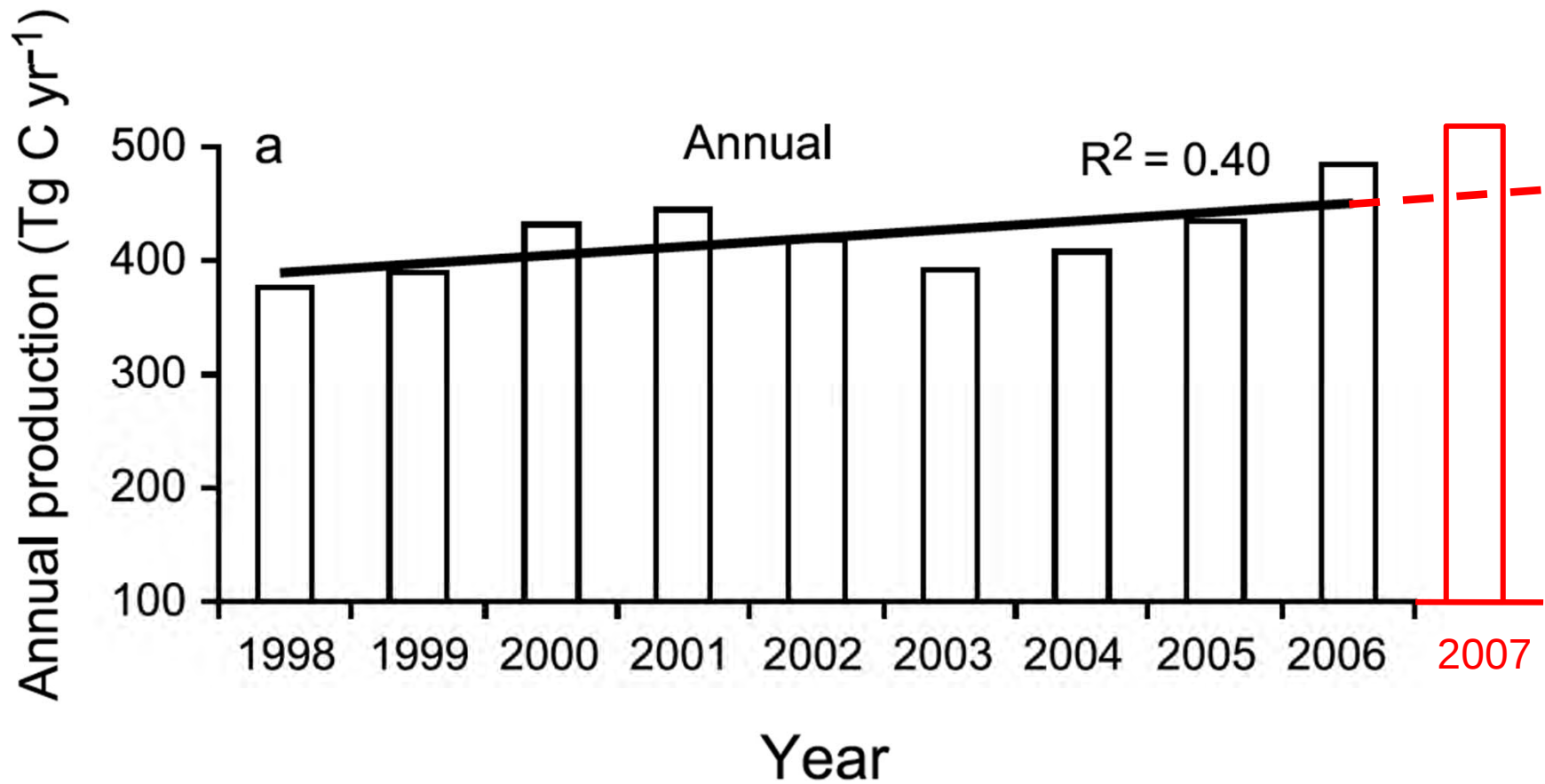
Arctic Primary Production

Arrigo et al. 2008

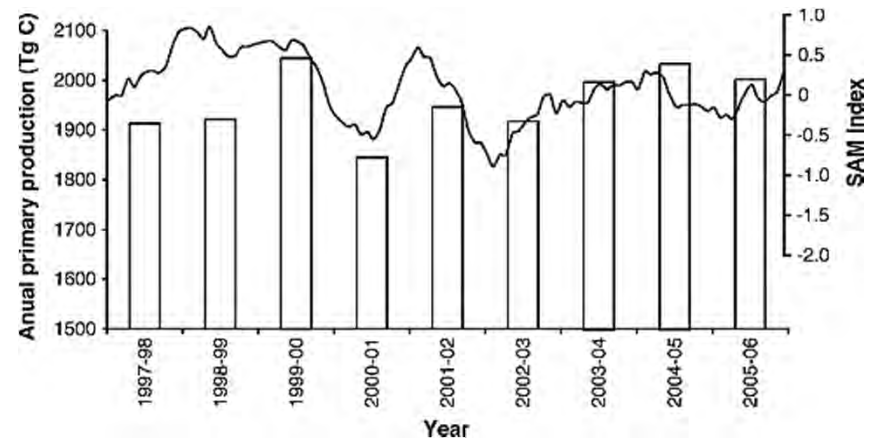
In 2007, an abrupt change!



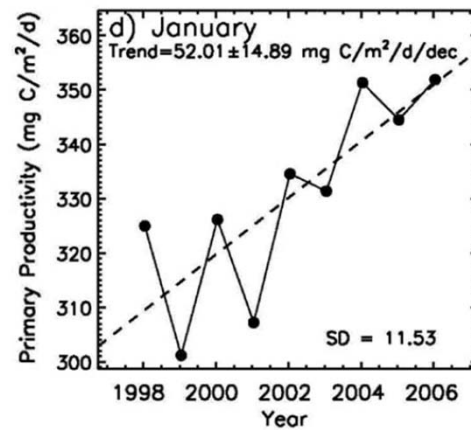
Arctic Primary Production



Antarctic Ocean



Arrigo et al. (2008) JGR



Smith & Comiso (2008) JGR

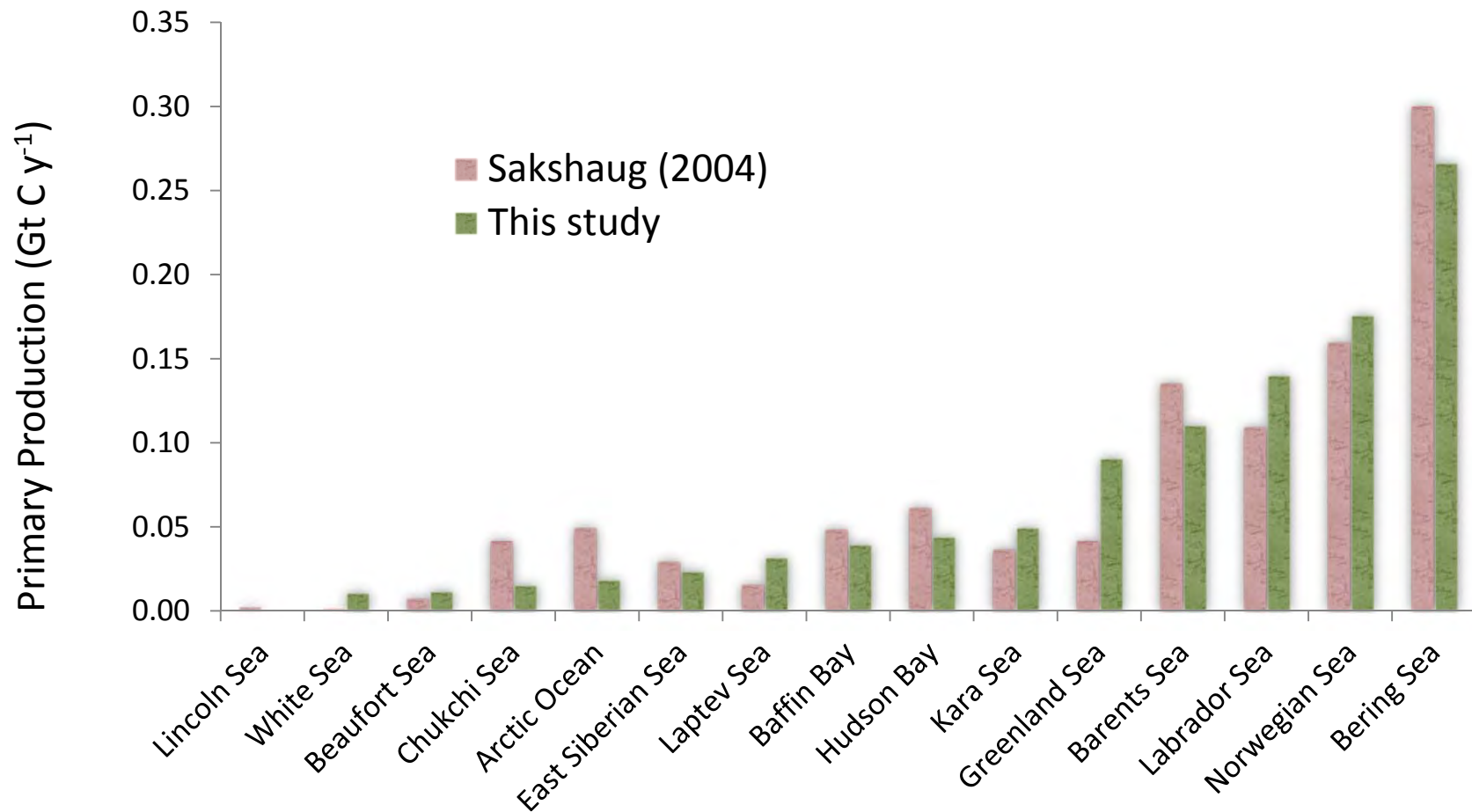
Results

Primary Production

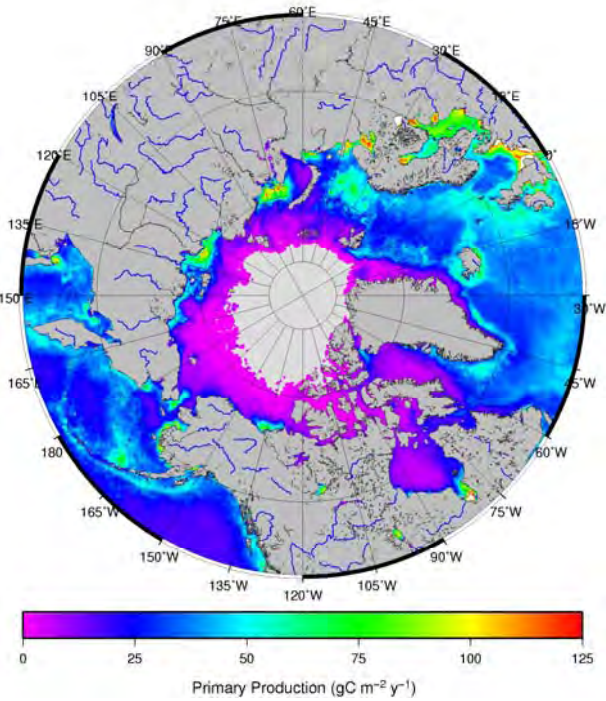
Latitude criterion	This study	Longhurst et al. (1995)	Behrenfeld & Falkowski (1997)	Antoine et al. (1996)	Pabi et al. (2008)
All Arctic (sensu IHO)	0.62*	-	-	-	-
>60°	0.46	-	-	-	-
>66.5°	0.26	-	-	-	0.42
>70°	0.19	1.4	0.4	0.6	-

* Gt C y⁻¹

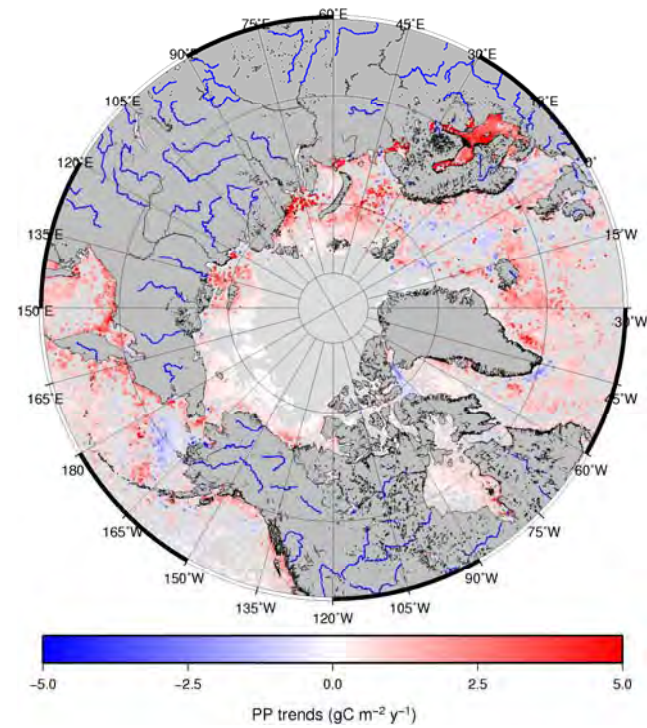
Results



Climatology (1998–2010)

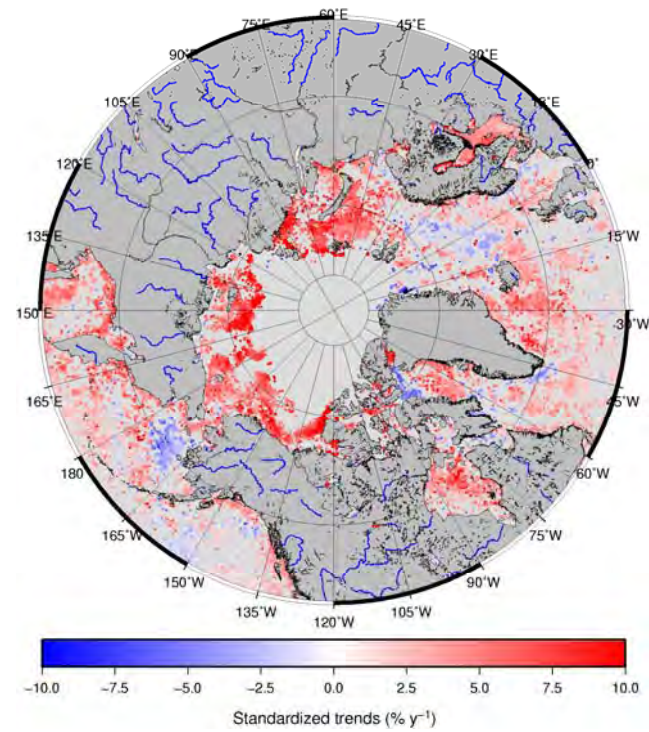


PP trends 1998–2010 (Sen Slope; $p < 0.05$)

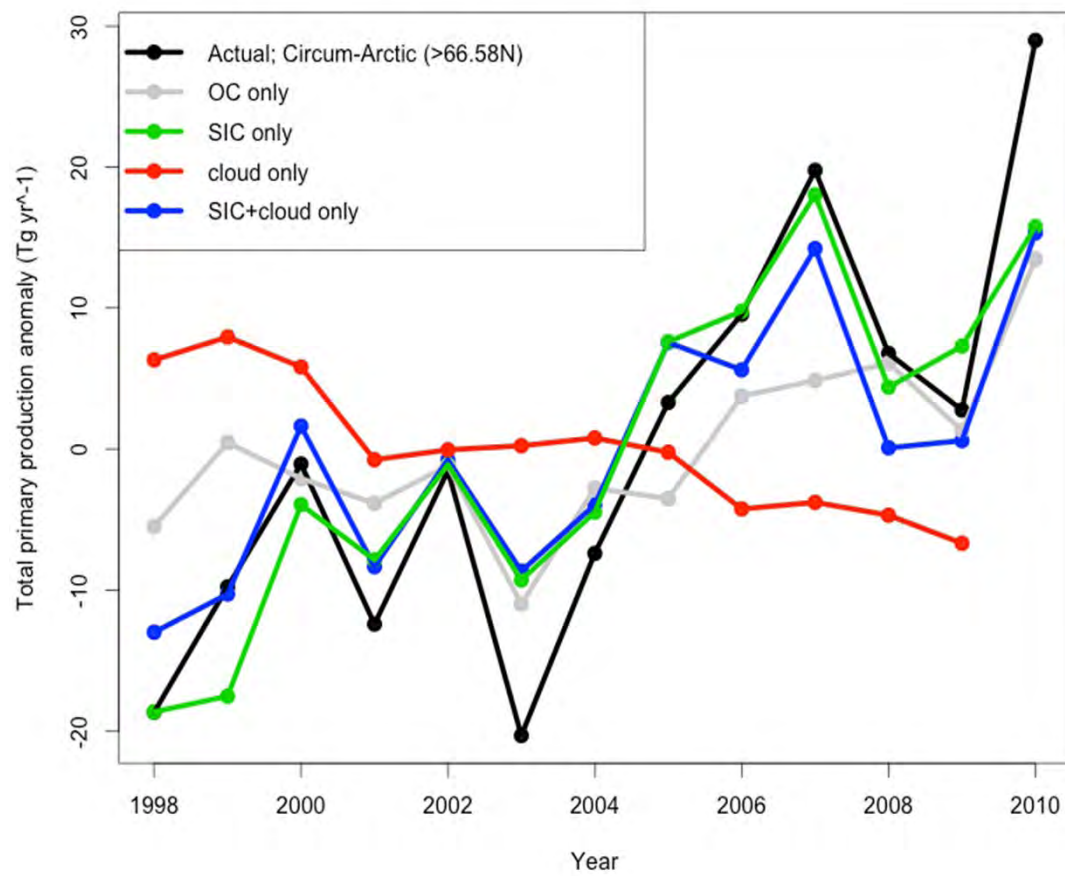


Bélangier et al.

PP trends 1998–2010 (Sen Slope; $p < 0.05$)

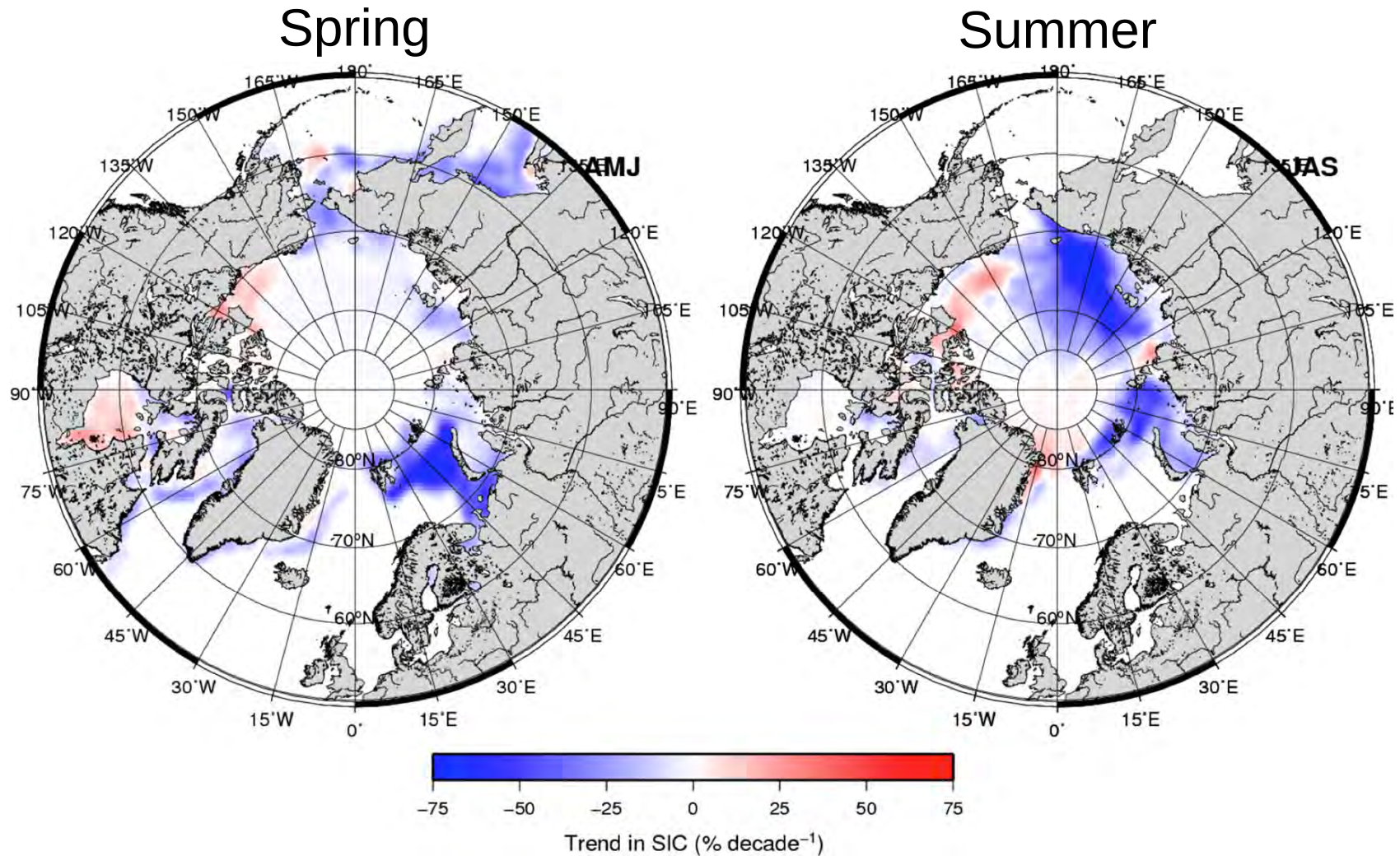


Bélanger et al.

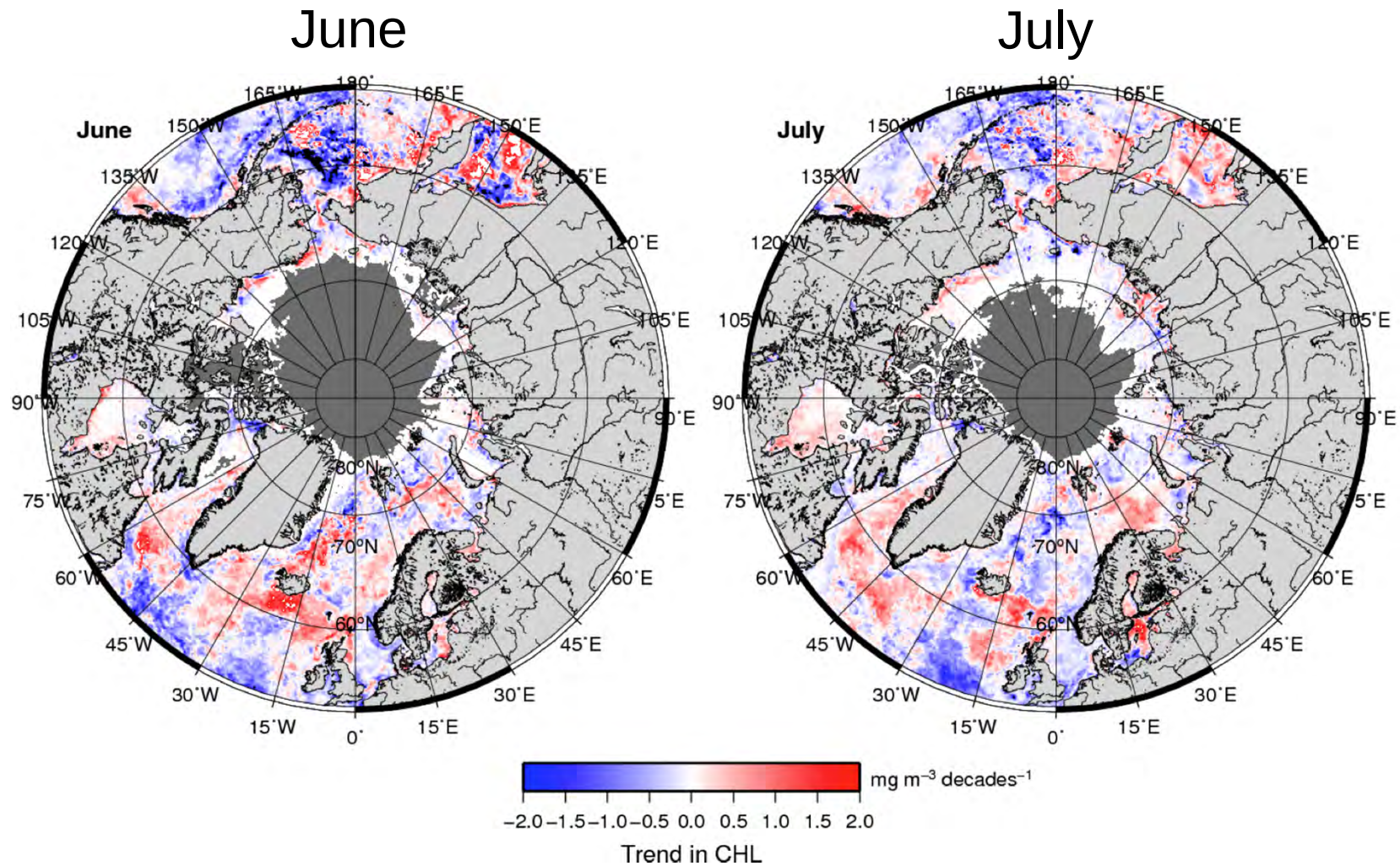


Bélangier et al.

Ice trend during the last decade

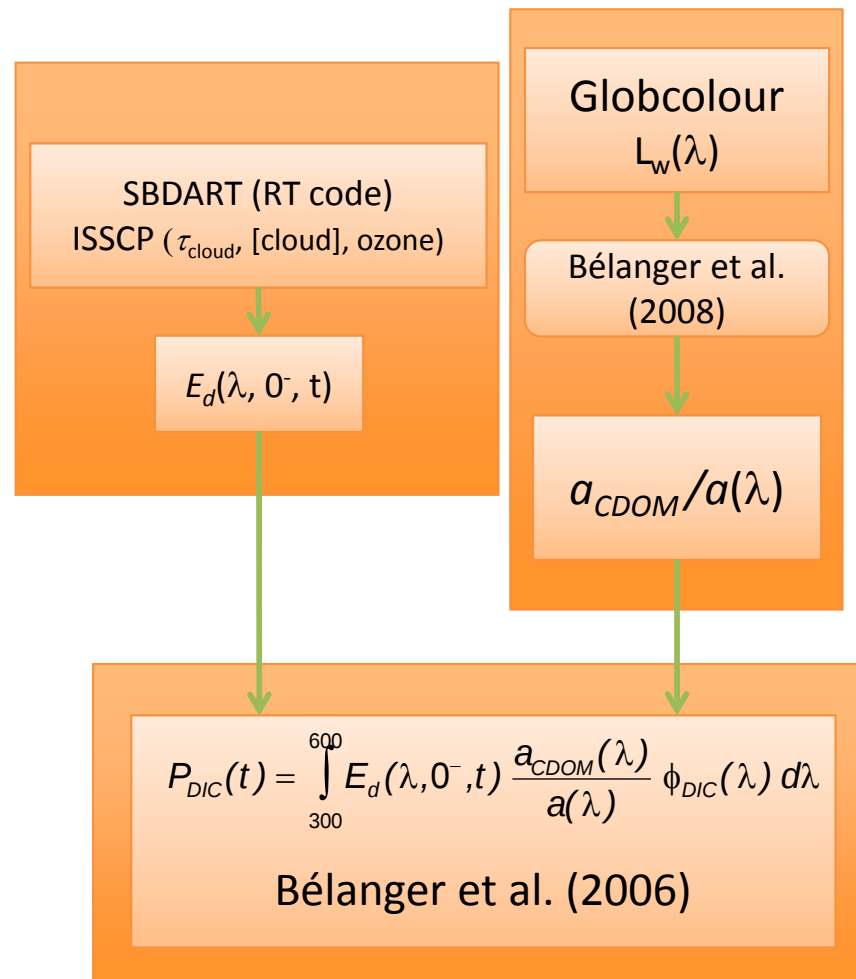


Trend in Chlorophyll concentration



What about CDOM photooxidation?

CDOM Photooxidation model



Primary Production vs. Photooxidation

Gross PP → $0.62 \times 10^{15} \text{ gC y}^{-1}$

Photooxidation (CO_2) → $2.7 \times 10^{13} \text{ gC y}^{-1}$

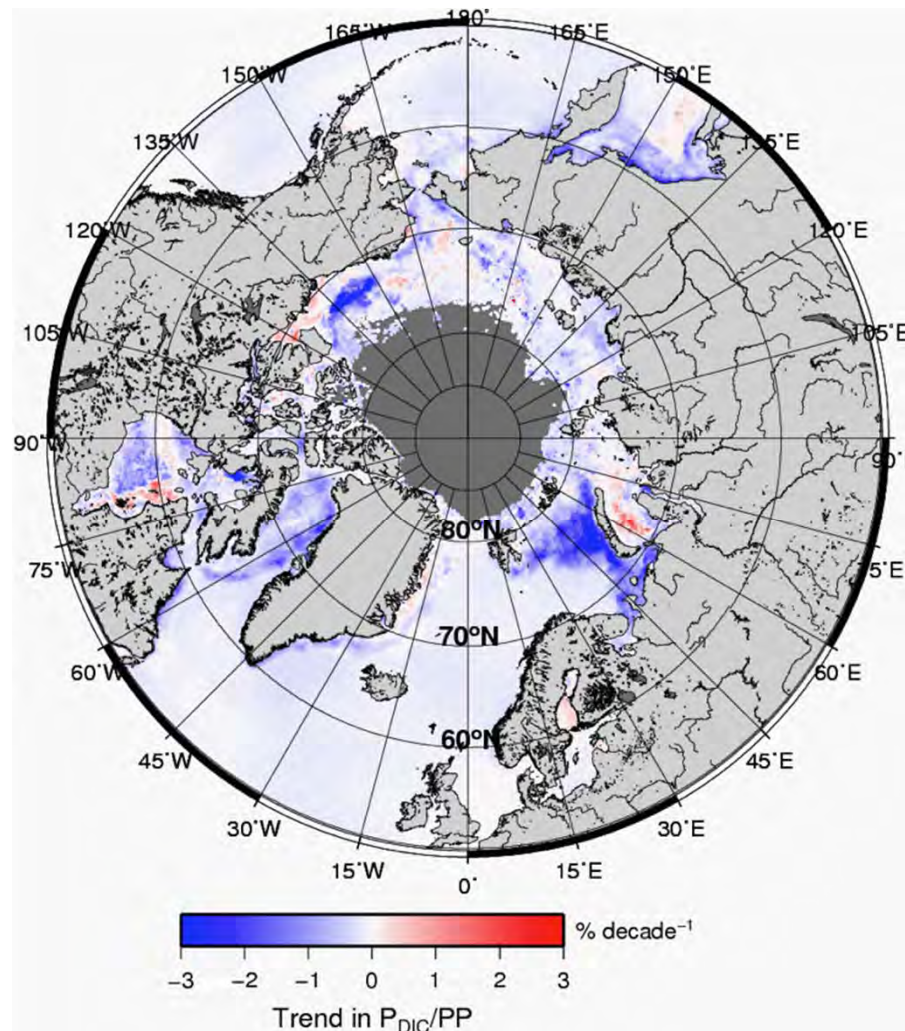
Photooxidation (CO_2) → 4.4% of PP

Sequestered C ~ 1% of PP (Stein & McDonald 2004)

So, potentially

PO of DOM > Sequestered C from PP

Trend in Photooxidation / Primary Production



Problems & Limitations

- Problems to solve:
 - Optical properties of Arctic seawater

PP : GSM VS OC4v4 for year 2007

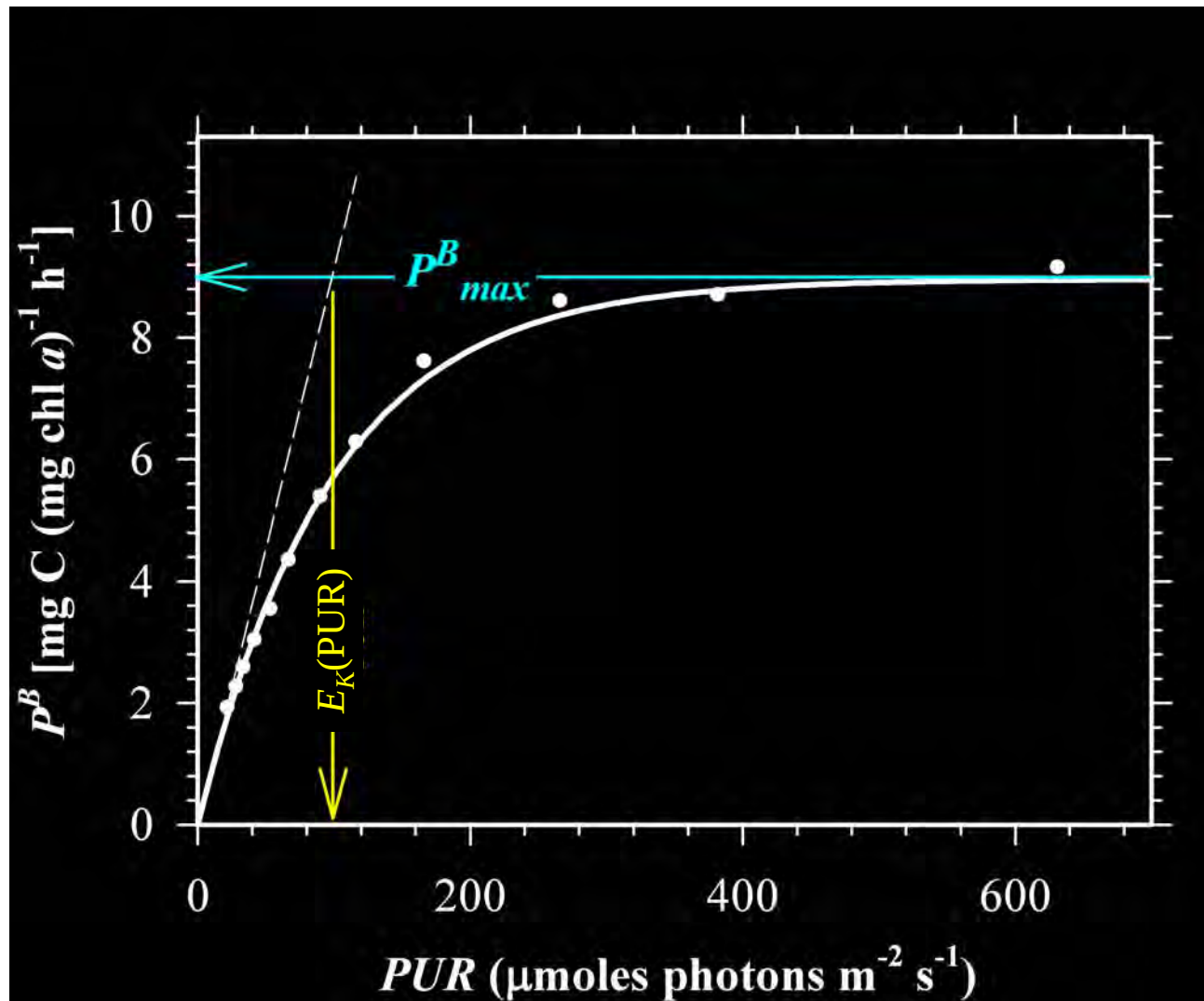
Region	GSM	OC4v4	%diff
Greenland	51.9	65	25.2%
Norwegian	97.7	98.3	0.6%
Barents	70.9	80	13%
Kara	24.5	39.5	61%
Laptev	15.3	22	44%
East Sib Sea	16.5	22.7	38%
Chukchi	9.6	14	46%
Beaufort	6.7	11.3	69%
Baffin Bay	11.9	18.1	52%
Hudson Bay	18.3	27.3	49%
Arctic Ocean	14.2	22.4	57%

Circumpolar Arctic: 293 378 29%

Problems & Limitations

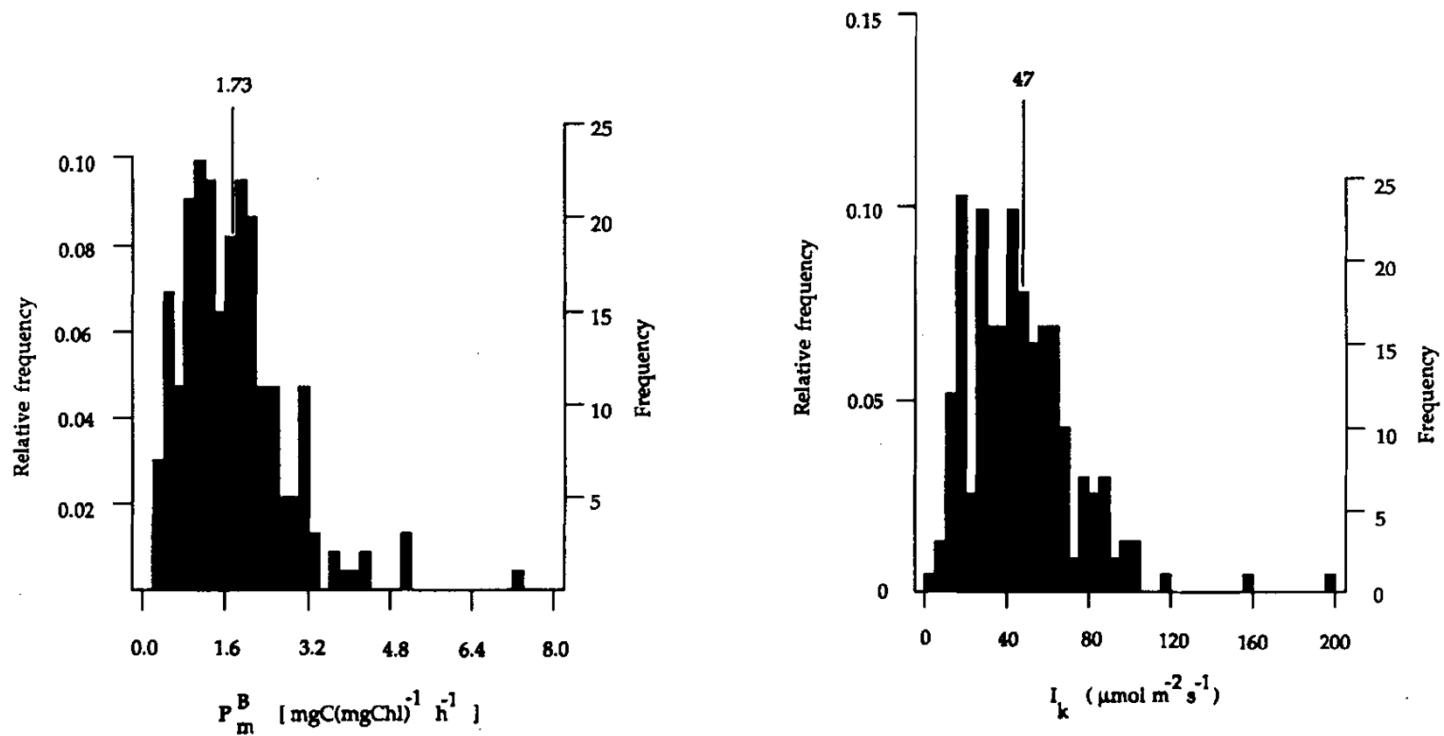
- Problems to solve:
 - Optical properties of Arctic seawater
 - **Photosynthetic properties of phytoplankton**

$$PP(z,t) = Chl \ P_B^{max} \left(1 - e^{-PUR(z)/E_K(PUR)}\right)$$



Photosynthetic parameters

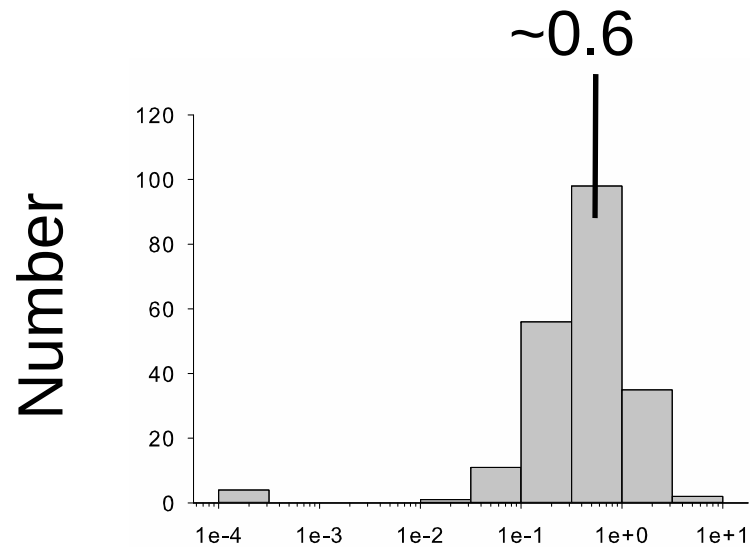
Barents Sea



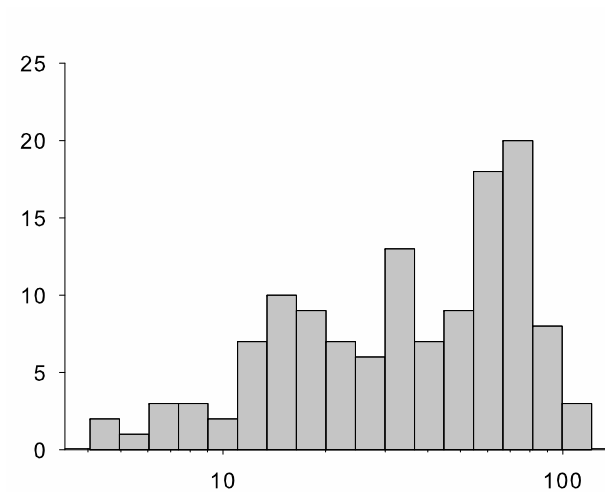
Rey (1991)

Photosynthetic parameters

Beaufort Sea in August 2009

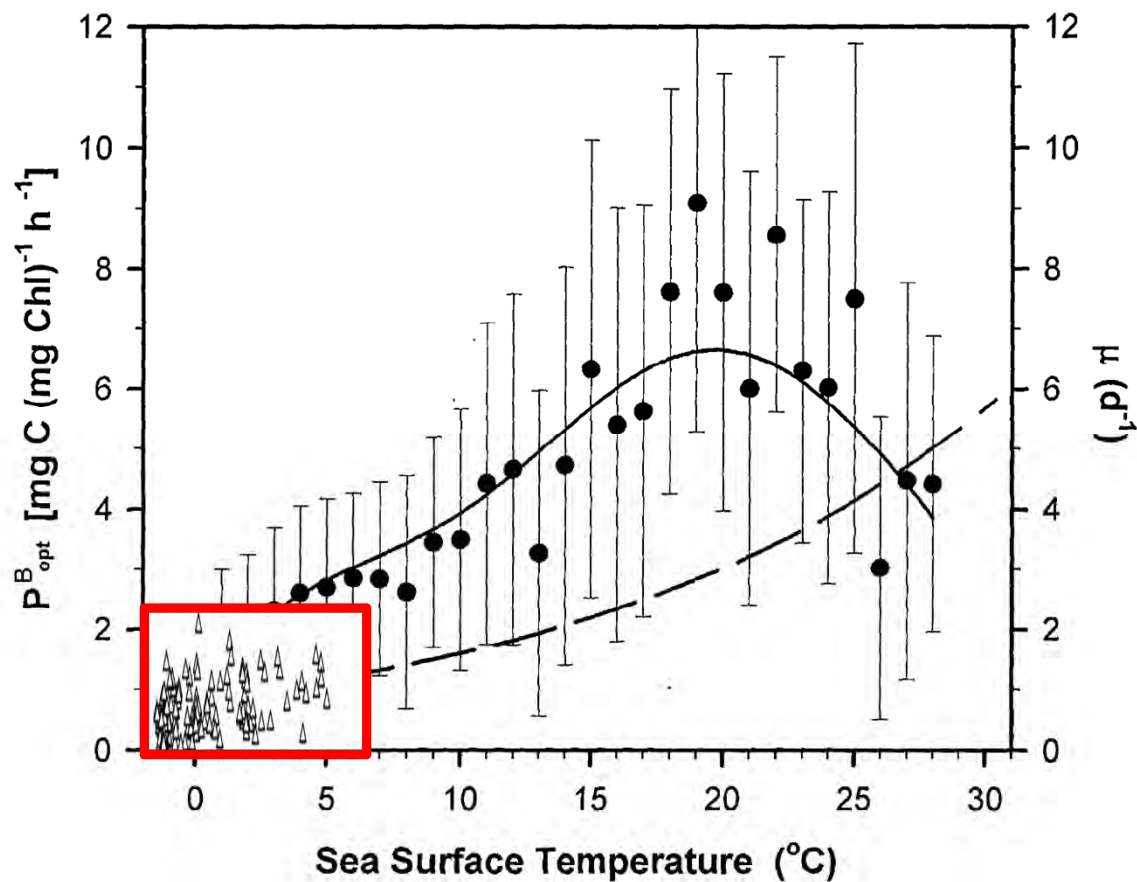


$P^B_{\max}$ [$\text{mg C (mg chl } a)^{-1} \text{ h}^{-1}$]



E_k ($\mu\text{mole photons m}^{-2} \text{ s}^{-1}$)

Photosynthetic parameters

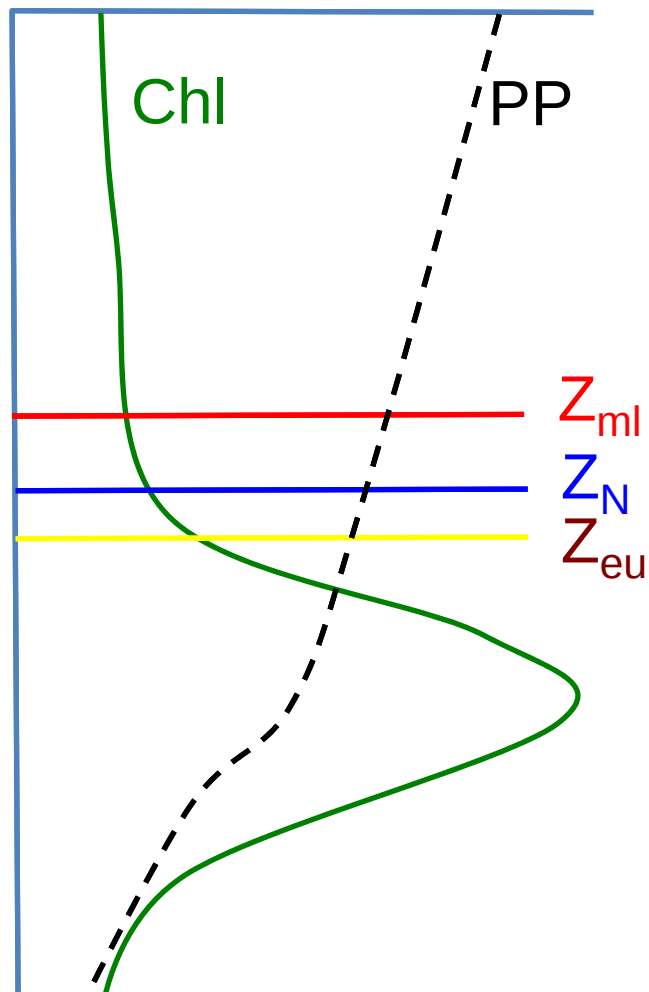


Problems & Limitations

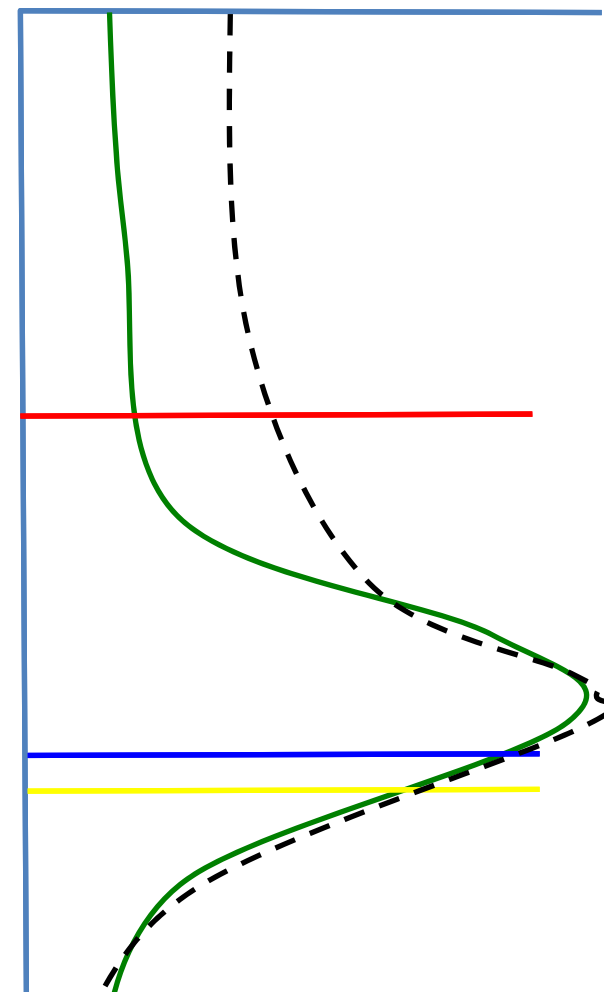
- Problems to solve:
 - Optical properties of Arctic seawater
 - Photosynthetic properties of phytoplankton
 - **Deep chlorophyll maximum**

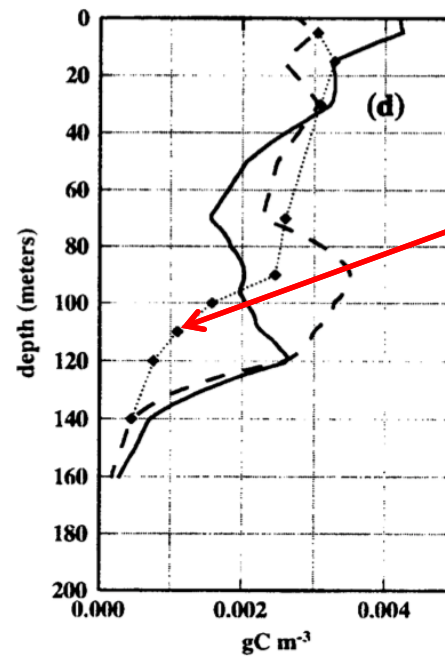
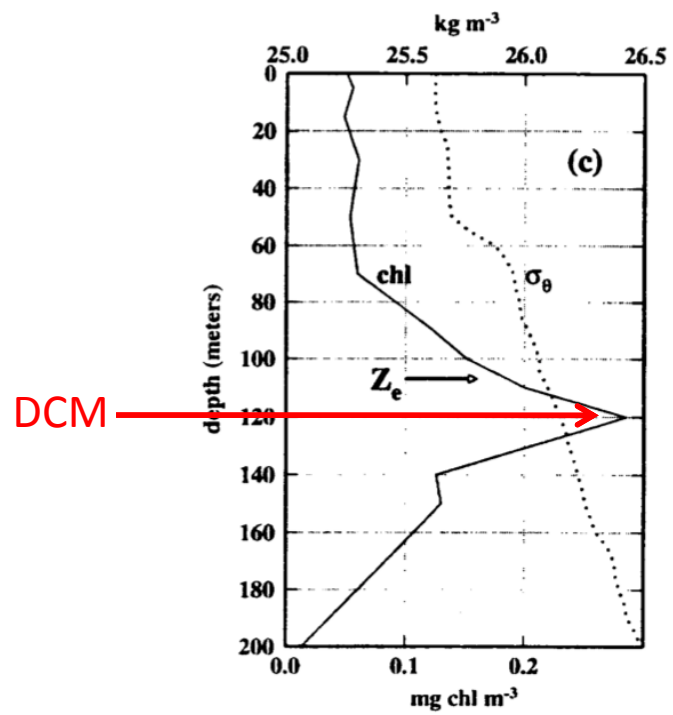
The DCM in the Arctic Ocean

At lower latitudes

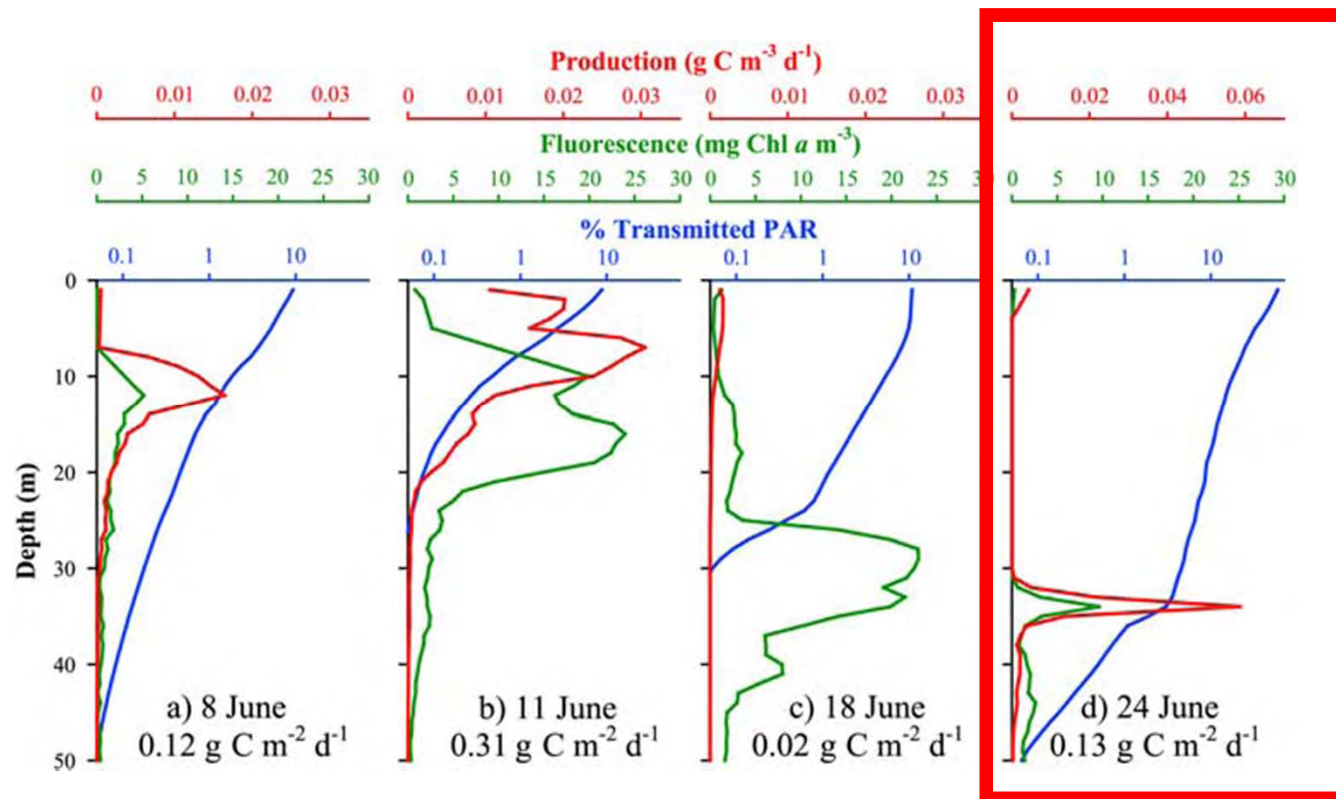


In the Arctic Ocean

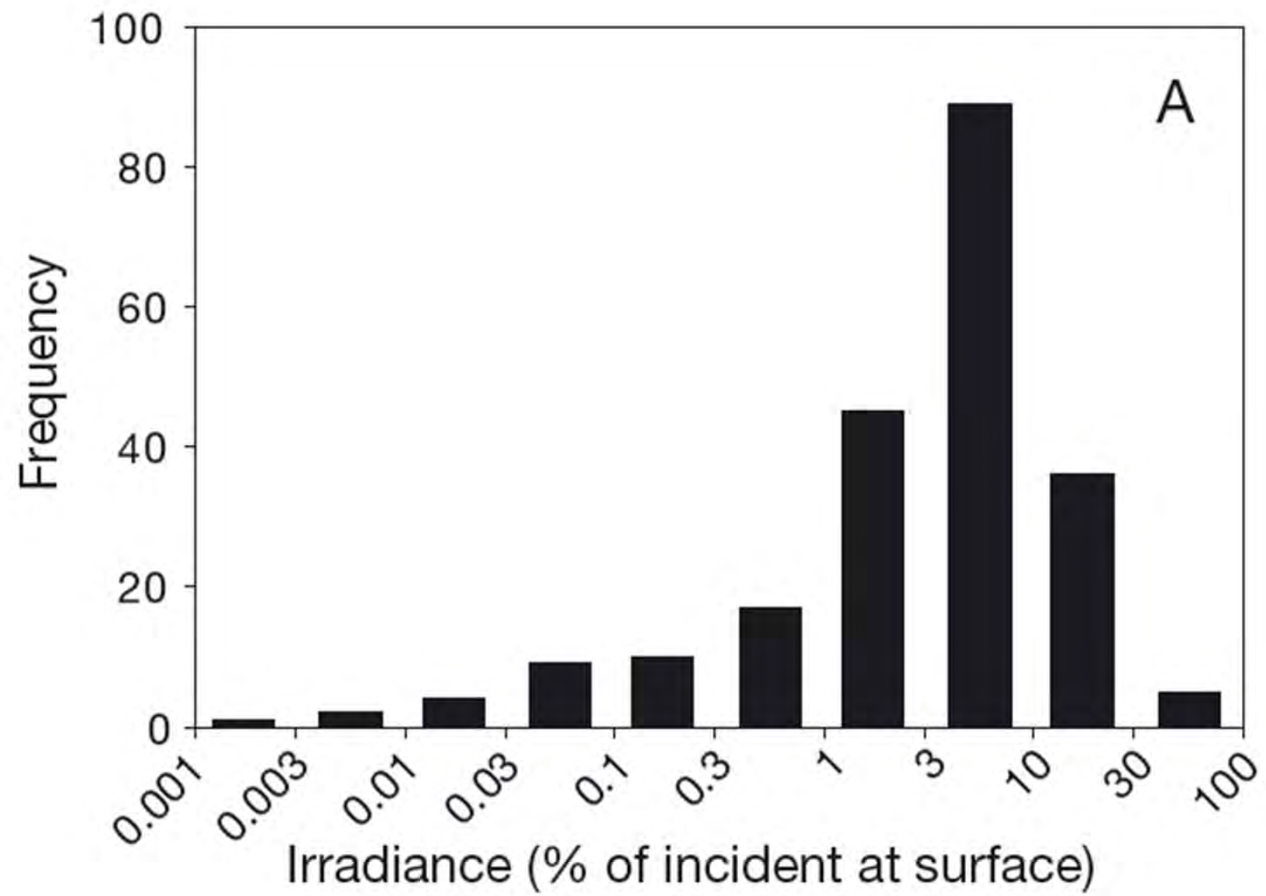




Morel et al. (1996)



Mundy et al. (2009)



Martin et al. 2010

PABI ET AL.: PRIMARY PRODUCTION IN THE ARCTIC OCEAN

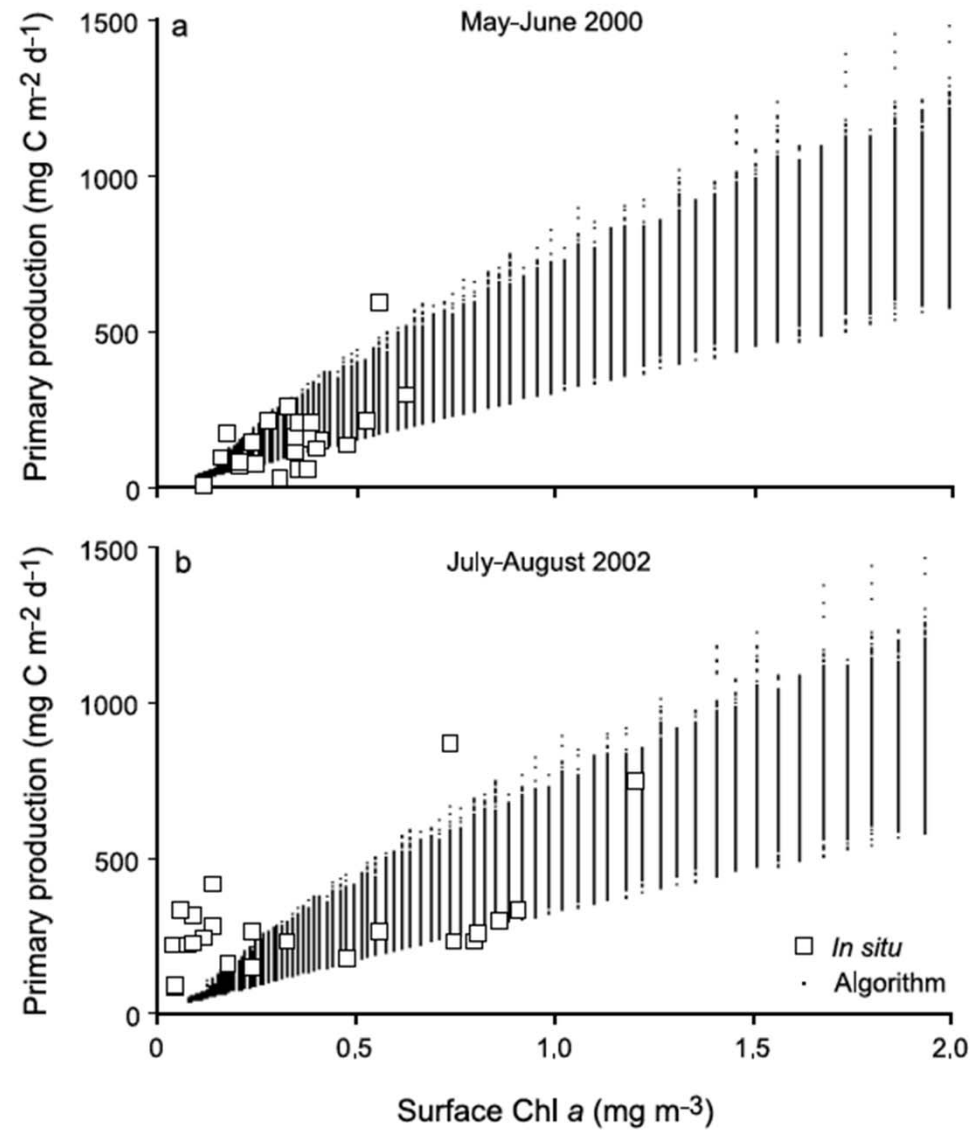


Table 3. Percent Change in Depth-Integrated Daily Net Primary Production Due to Removal of the Subsurface Chl *a* Maximum for Different Geographic Sectors and Different Time Periods^a

	Jan–Mar	Apr–Jun	Jul–Sep	Oct–Dec	Annual
<i>Chlorophyll a Constant in the Upper 100 m (Method 2)</i>					
Chukchi	–18.9	–12.0	–6.1	3.0	–7.6
Beaufort	–76.9	8.4	–20	18.3	–10.8
Baffin	3.1	2.2	–3.8	–2.9	–3.7
Greenland	–16.8	0	–5.1	13.7	–3.0
Barents	18.8	0.9	–0.1	7.5	1.9
Kara	ND ^b	8.6	9.1	2.4	8.9
Laptev	ND	ND	ND	ND	ND
Siberian	ND	ND	0.3	ND	0.3
All	10.3	–9.0	–6.6	–5.9	–7.4
<i>Chlorophyll a Constant in Upper 20 m and Declines Exponentially With Depth (Method 3)</i>					
Chukchi	–25.9	–15.5	–10.7	–9.1	–12.3
Beaufort	–85.4	2.1	–30.7	–12.8	–20.8
Baffin	–10.8	–5.2	–5.4	–16.4	–6.1
Greenland	–26.6	–8.4	–16.6	–10.9	–13.0
Barents	–10.9	–8.2	–14.6	–23.4	–9.9
Kara	ND	6.3	1.9	–5.9	4.7
Laptev	ND	ND	ND	ND	ND
Siberian	ND	ND	–7.9	ND	–7.9
All	–7.3	–14.7	–12.0	–19.7	–13.2
<i>Chlorophyll a Constant in Upper 40 m and Declines Exponentially With Depth (Method 4)</i>					
Chukchi	–19.9	–12.3	–6.5	0.7	–8.0
Beaufort	–81.0	7.8	–22.3	6.6	–12.6
Baffin	0.1	1.3	–3.9	–5.9	–3.8
Greenland	–18.7	–1.2	–7.4	5.9	–4.7
Barents	8.1	–0.5	–3.4	–4.9	–0.3
Kara	ND	8.5	8.3	1.2	8.6
Laptev	ND	ND	ND	ND	ND
Siberian	ND	ND	–0.9	ND	–0.9
All	6.0	–9.6	–7.1	–9.1	–8.1

^aChl *a* was distributed vertically using methods 2–4.

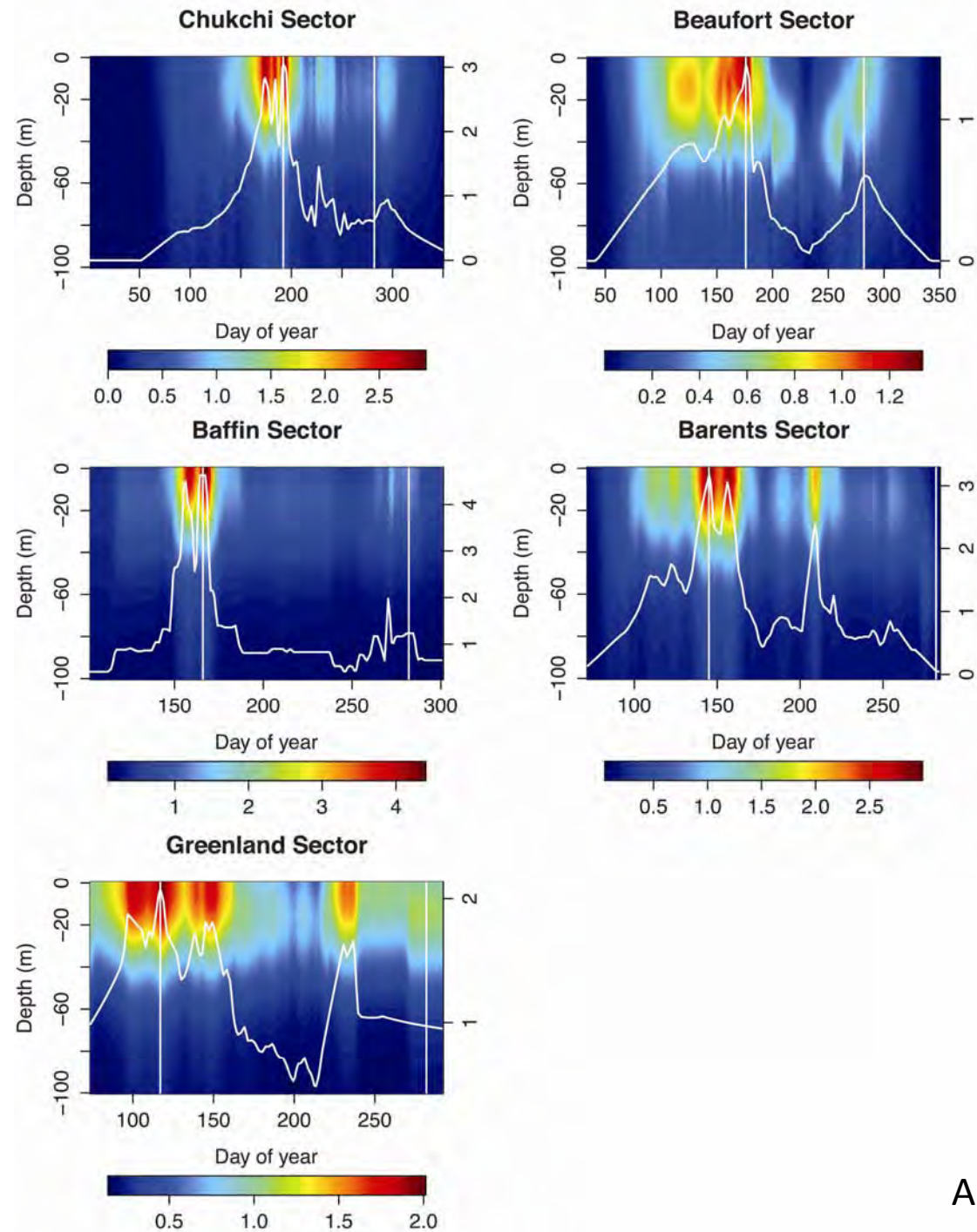
^bND indicates no in situ data were available.

Table 4. Percent Change in Depth-Integrated Net Primary Production Caused by Increasing in Situ Chl *a* by the RMSE of Satellite-Derived Chl *a* for Different Geographic Sectors and Different Time Periods^a

	January–March	April–June	July–September	October–December	Annual
<i>Chlorophyll a Constant in Upper 100 m (Method 2)</i>					
Chukchi	−10.7	−7.5	−0.4	18.1	−1.9
Beaufort	274.0	15.8	−6.0	121.8	1.6
Baffin	21.5	10.7	−0.9	15.2	0.0
Greenland	−4.5	9.7	9.2	66.1	9.1
Barents	104.0	11.4	20.1	145	16.5
Kara	ND ^b	12.4	17.5	11.9	14.5
Laptev	ND	ND	ND	ND	ND
Siberian	ND	ND	9.7	ND	9.7
All	36.4	−2.4	−0.3	13.4	−0.7
<i>Chlorophyll a Constant in Upper 20 m and Declines Exponentially With Depth (Method 3)</i>					
Chukchi	−17.4	−10.8	−4.8	6.4	−6.3
Beaufort	156.2	9.9	−16.4	77.6	−8.1
Baffin	7.9	3.8	−2.4	2.0	−2.3
Greenland	−14.0	1.8	−1.8	37.9	−0.5
Barents	65.0	2.8	5.7	93.0	5.2
Kara	ND	10.2	10.7	4.1	10.4
Laptev	ND	ND	ND	ND	ND
Siberian	ND	ND	2.0	ND	2.0
All	18.6	−7.7	−5.3	−0.3	−6.1
<i>Chlorophyll a Constant in Upper 40 m and Declines Exponentially With Depth (Method 4)</i>					
Chukchi	−11.6	−7.7	−0.8	16.1	−2.3
Beaufort	209.9	15.3	−7.9	108.6	−0.0
Baffin	19.0	9.9	−0.9	12.7	−0.1
Greenland	−6.0	8.7	7.3	58.8	7.6
Barents	92.8	10.2	17.3	128.6	14.7
Kara	ND	12.3	16.8	11.0	14.2
Laptev	ND	ND	ND	ND	ND
Siberian	ND	ND	8.8	ND	8.8
All	32.7	−2.9	−0.8	10.7	−1.2

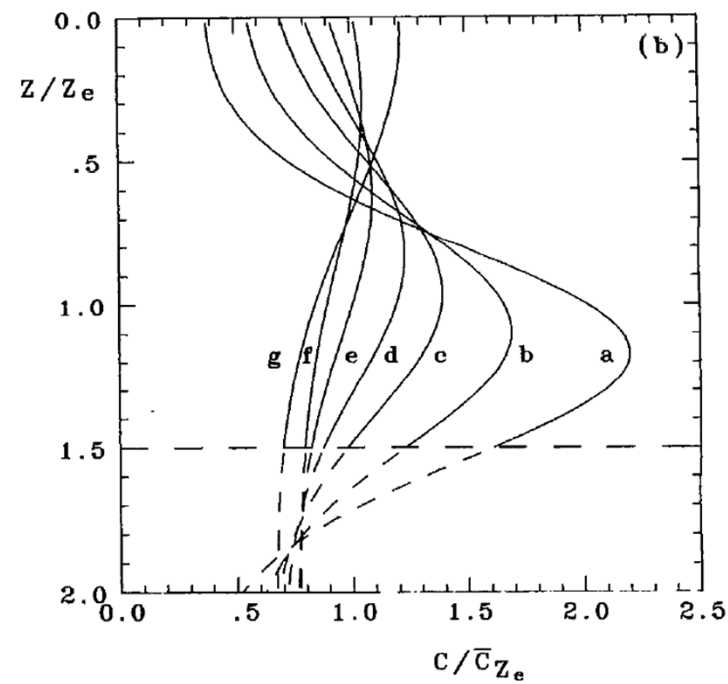
^aChl *a* was distributed vertically using methods 2–4.

^bND indicates no in situ data were available.



Ardyna et al.

The DCM in the Arctic Ocean



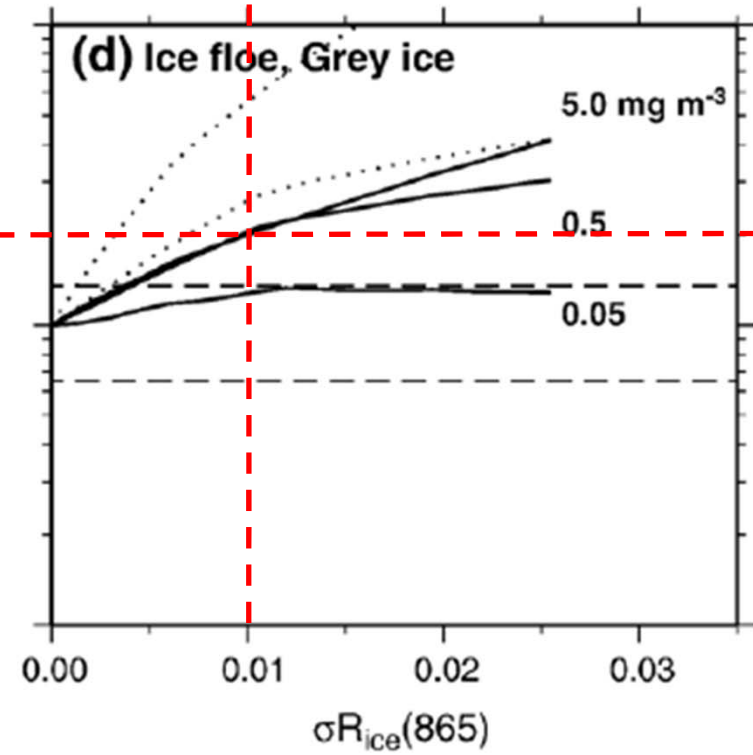
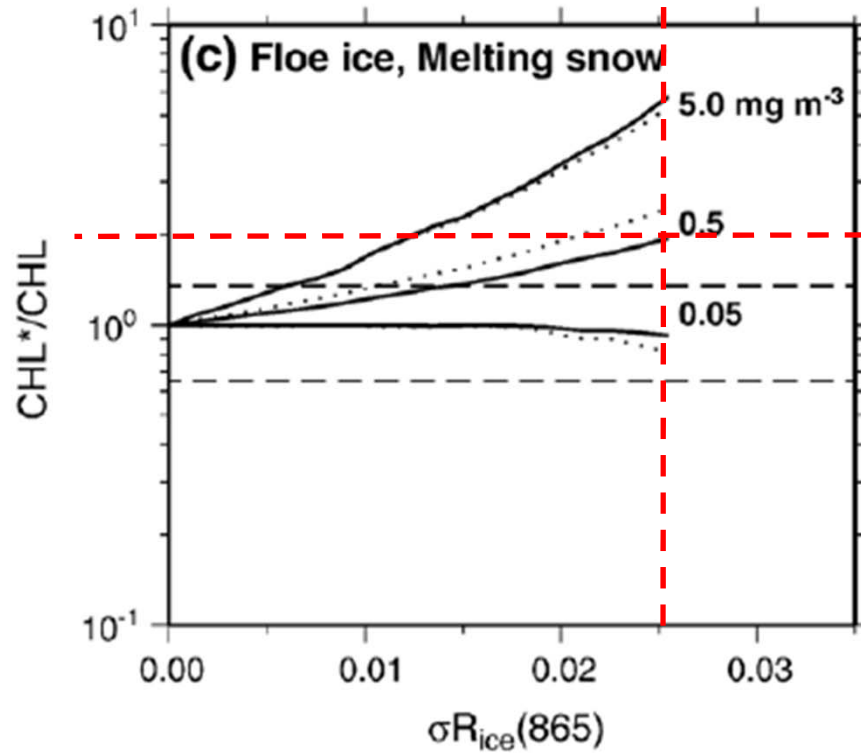
Morel & Berthon (1989)

Problems & Limitations

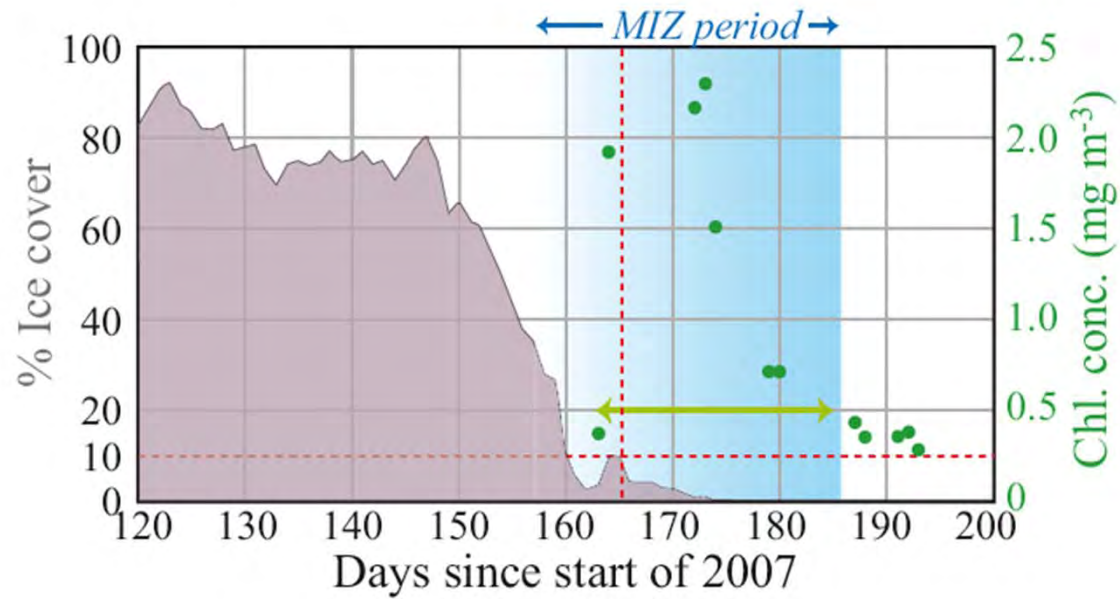
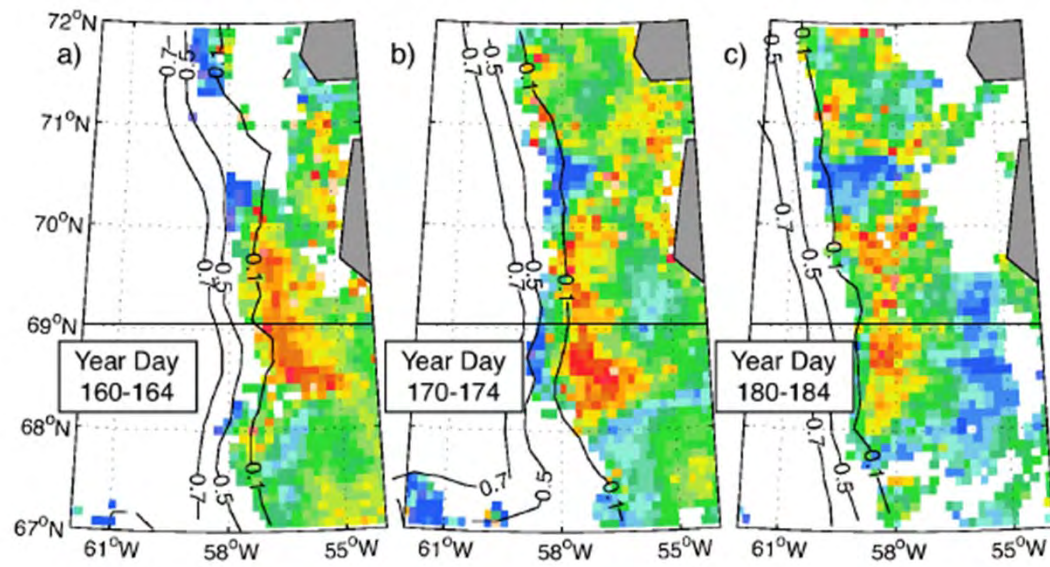
- Problems to solve:
 - Optical properties of Arctic seawater
 - Photosynthetic properties of phytoplankton
 - Deep chlorophyll maximum
 - **Ice**

% ice cover = 7.5%

% ice cover = 10%



Bélanger et al. (2007)



Perrette et al. 2011

Problems & Limitations

- Problems to solve:
 - Optical properties of Arctic seawater
 - Photosynthetic properties of phytoplankton
 - Deep chlorophyll maximum
 - Ice
- Limitations:
 - **Low Sun elevation**

Impact of pixels with no OC data

Region Considered	Relative difference between PP with OC only, and PP for all ice-free areas
All Arctic waters	-6.7%
> 60°	-30%
Arctic Basin	-47%

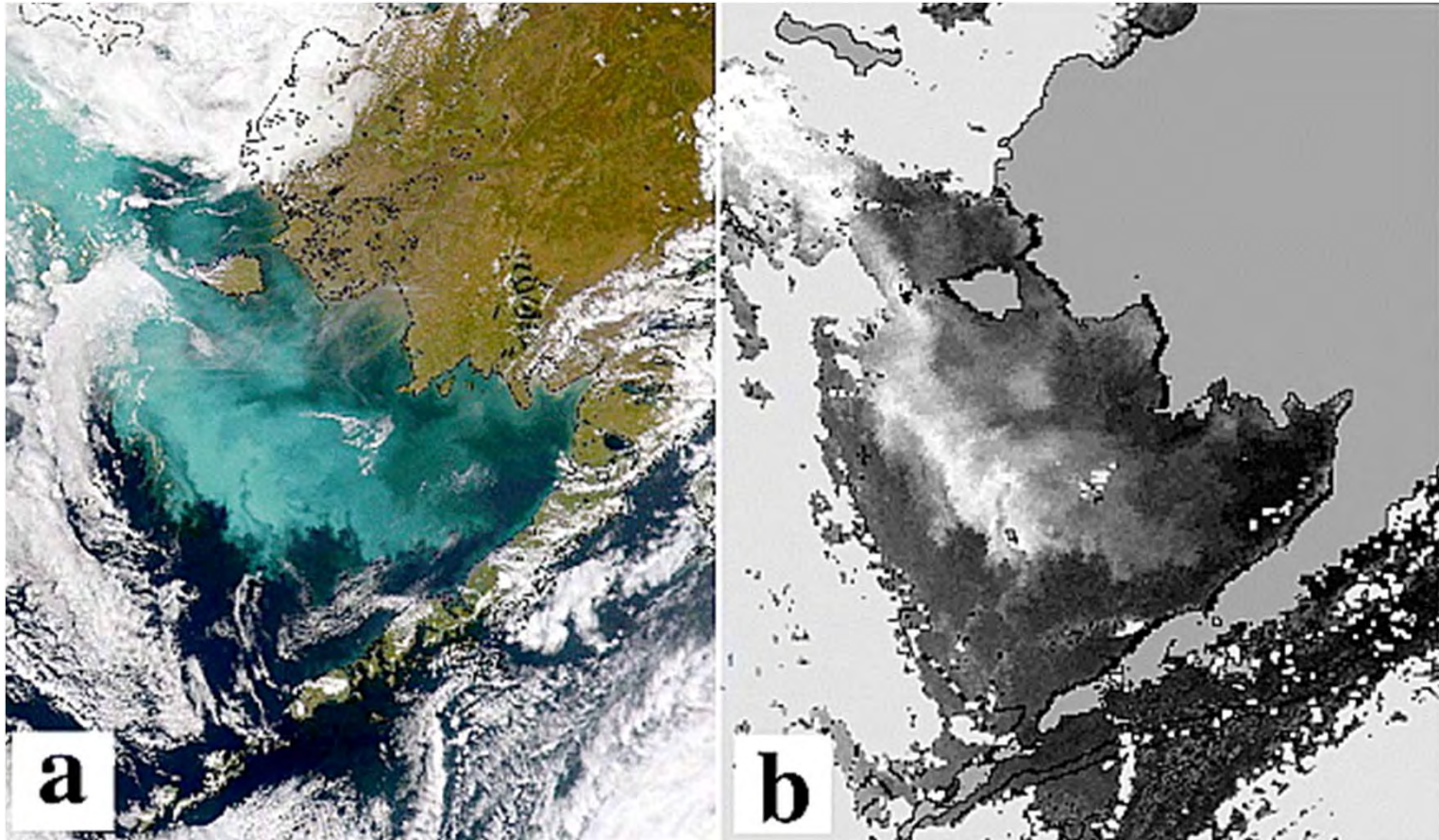
Problems & Limitations

- Problems to solve:
 - Optical properties of Arctic seawater
 - Photosynthetic properties of phytoplankton
 - Deep chlorophyll maximum
 - Ice
- Limitations:
 - Low Sun elevation
 - **Clouds & fog**

Tracking regime shifts with OC?

Other examples:

- Northward migration of phytoplankton groups/species



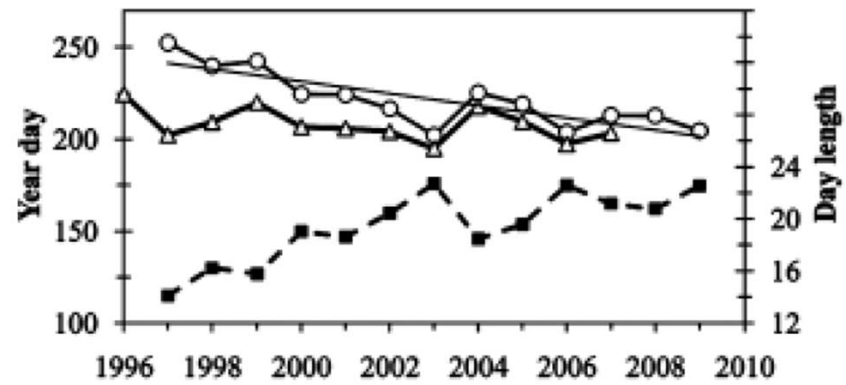
Merico et al. (2006)

Tracking regime shifts with OC?

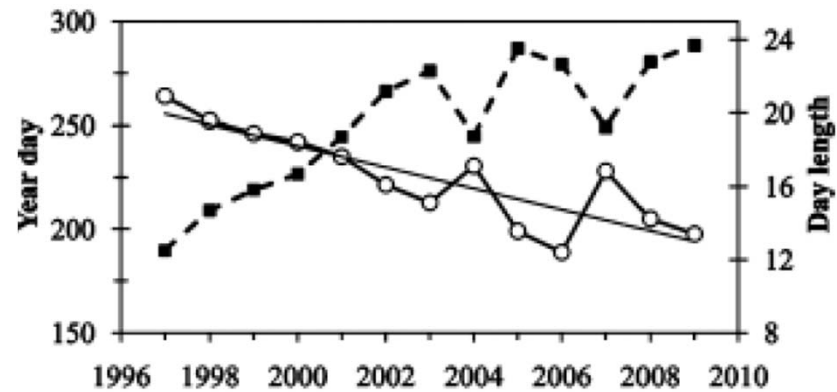
Other examples:

- Northward migration of phytoplankton groups/species
- **Timing of the phytoplankton bloom**

Baffin Bay



Kara Sea



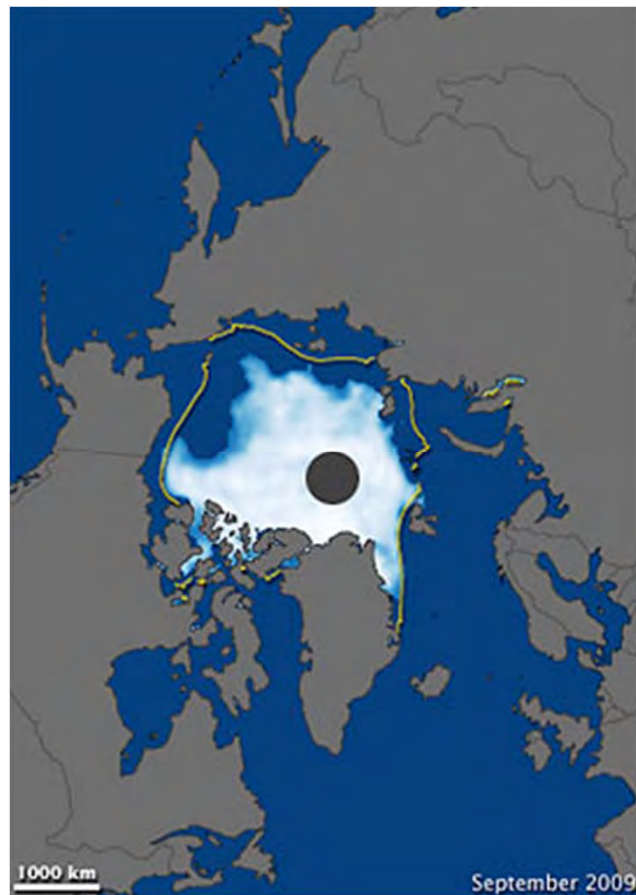
Kahru et al. 2010

What's next?

- To better document Arctic Ocean optical properties
- To improve and validate OC algorithms
- To address the DCM problem
- To further document phytoplankton photosynthesis
- ...

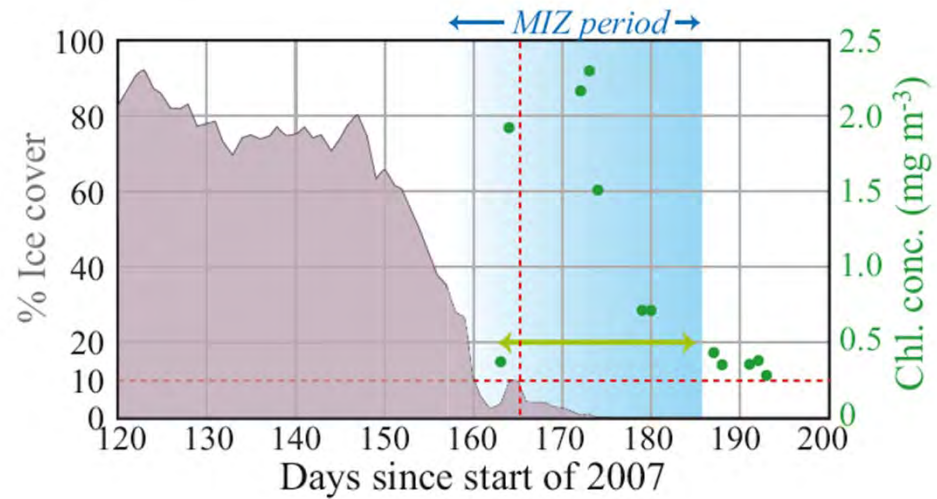
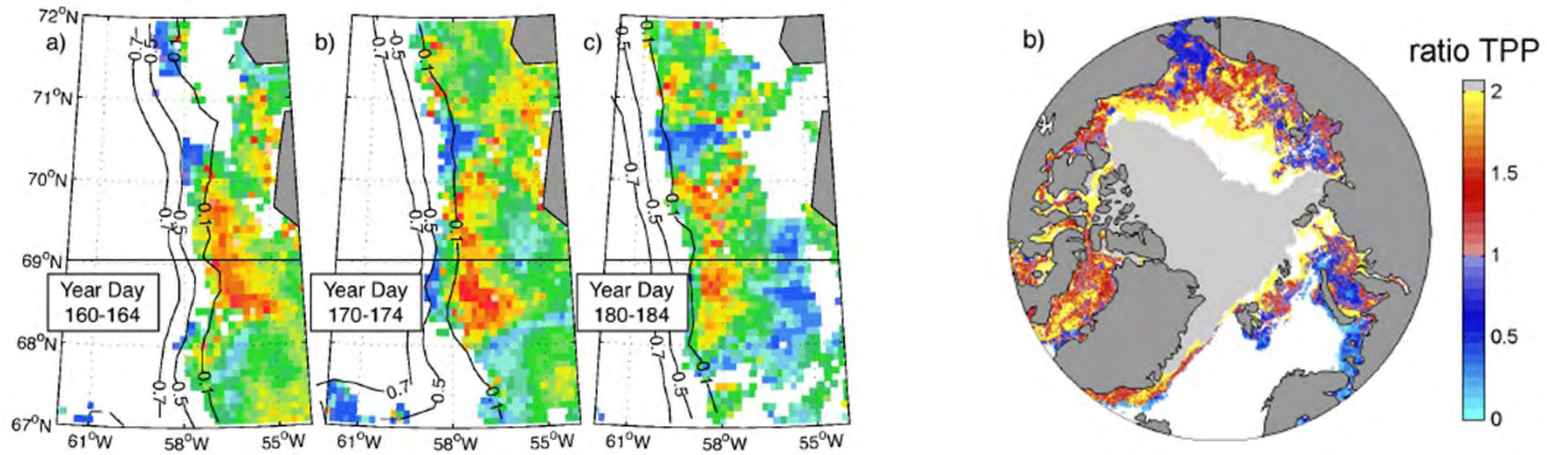
The importance of ice-edge blooms

Sea ice

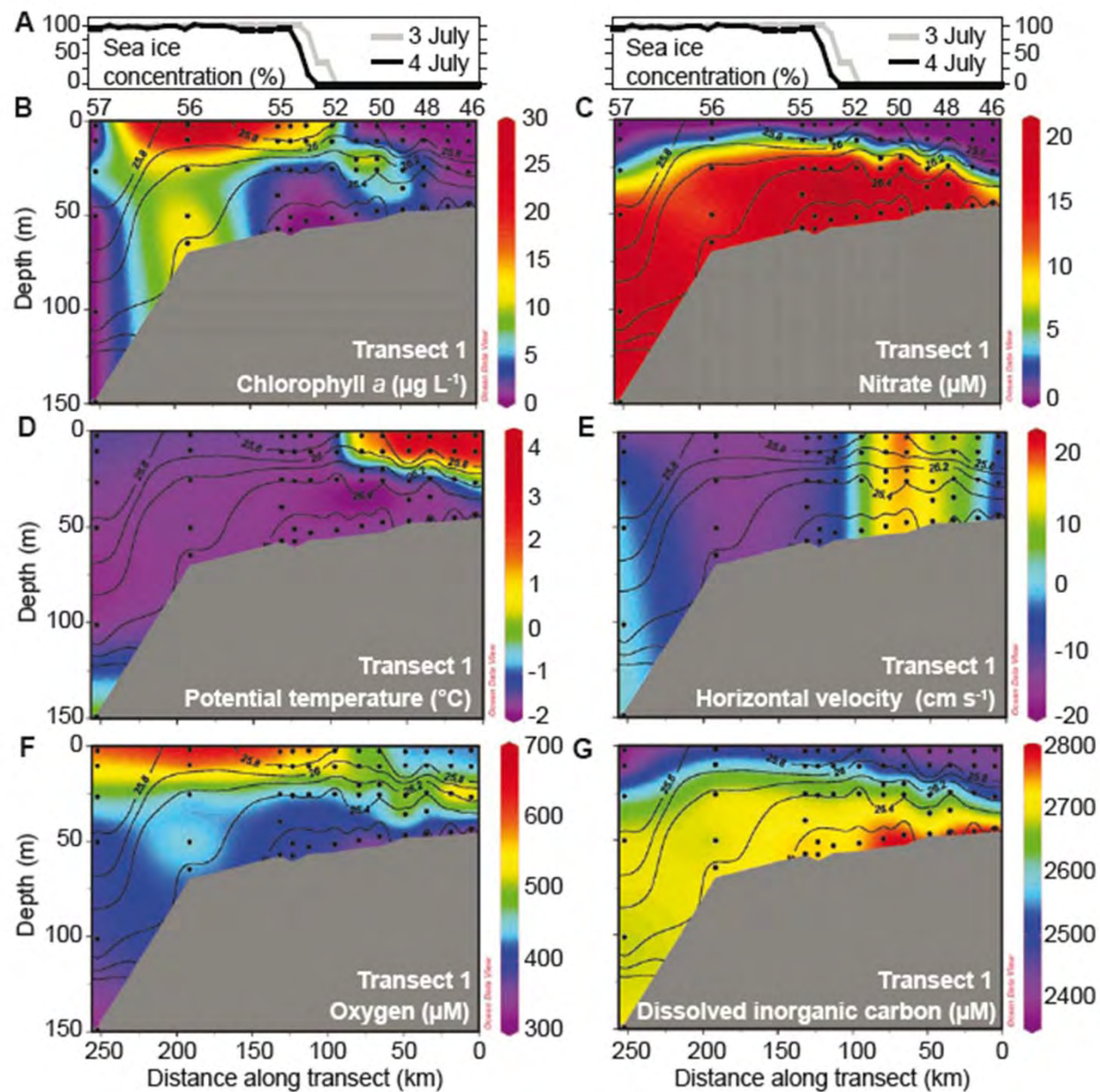




Efflorescences de marge de banquise



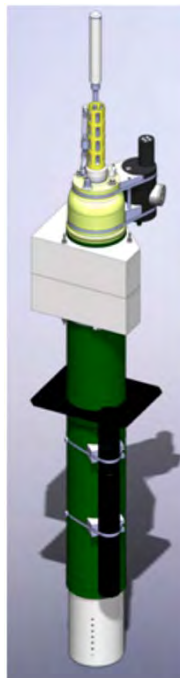
Perrette et al. 2011

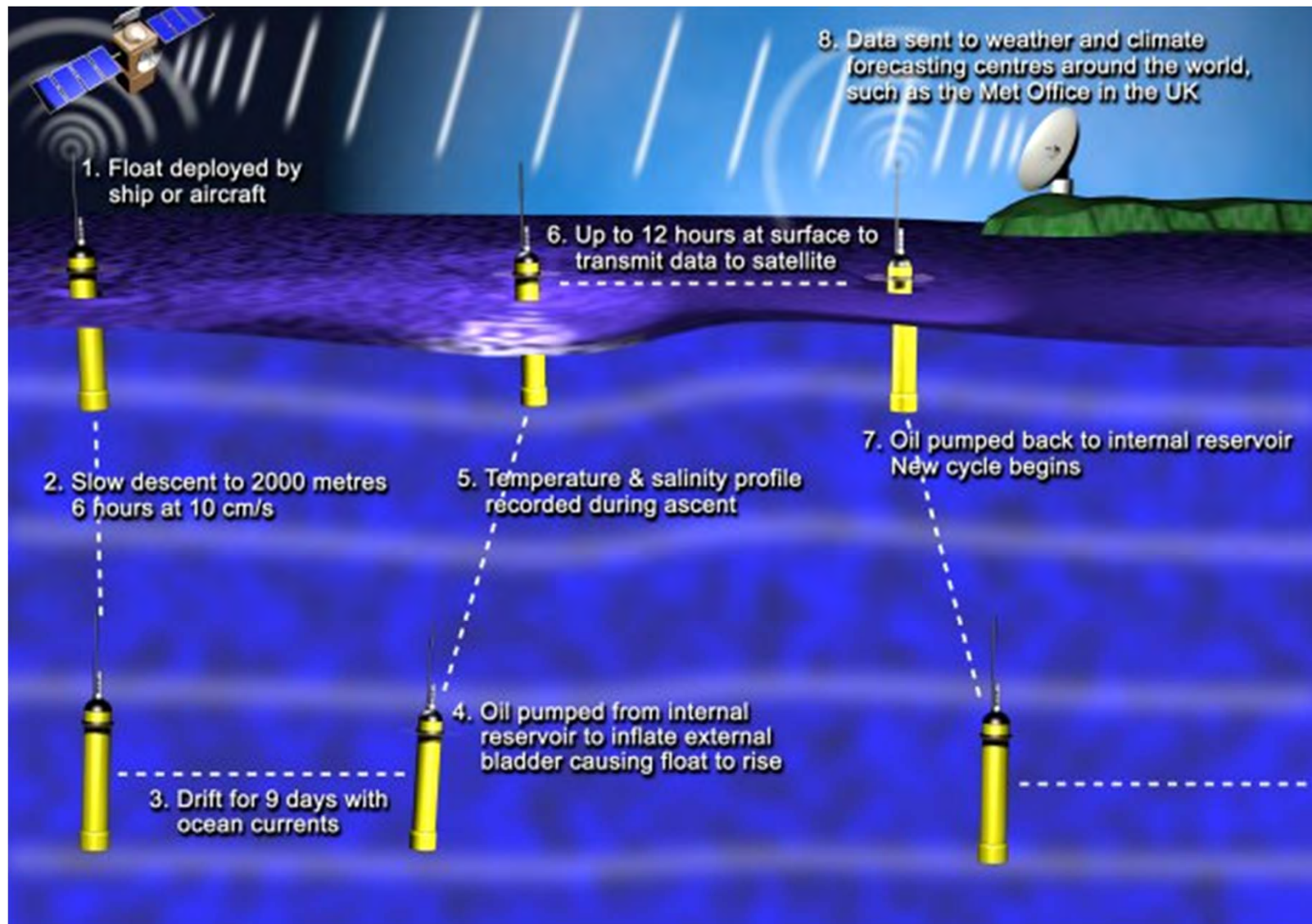


Arrigo et al. (2012), Science

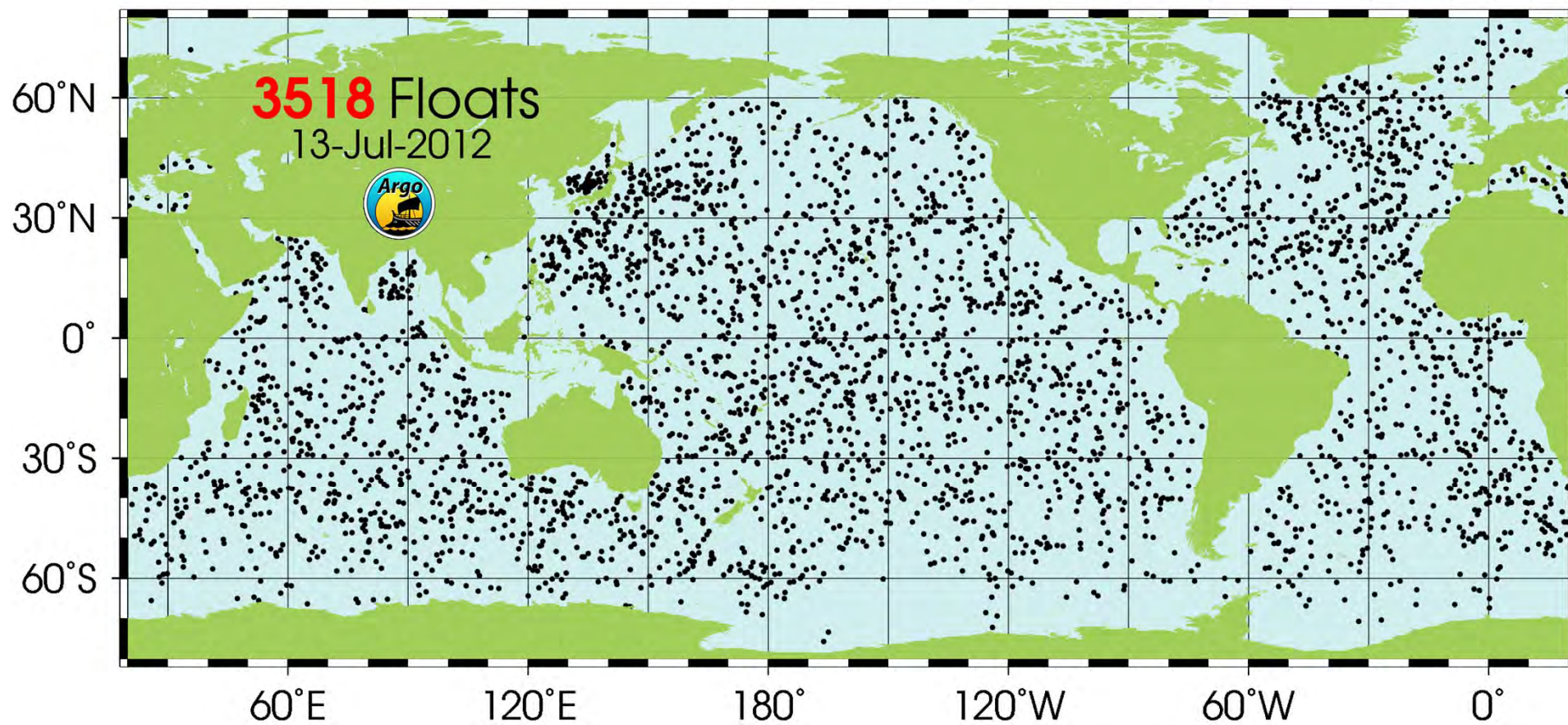


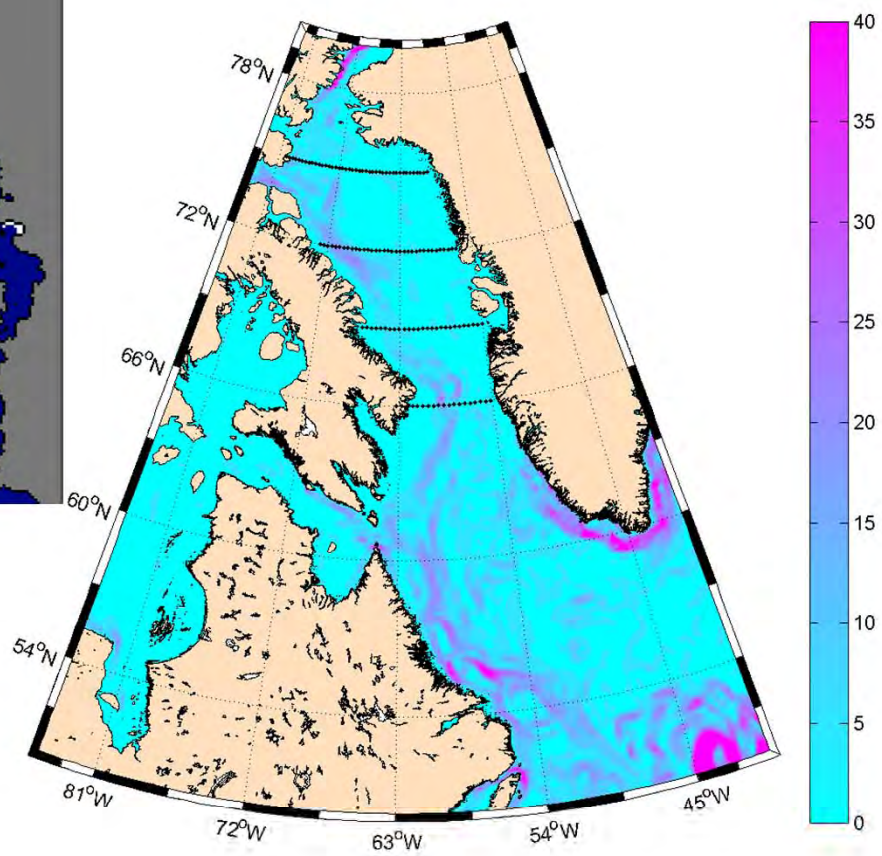
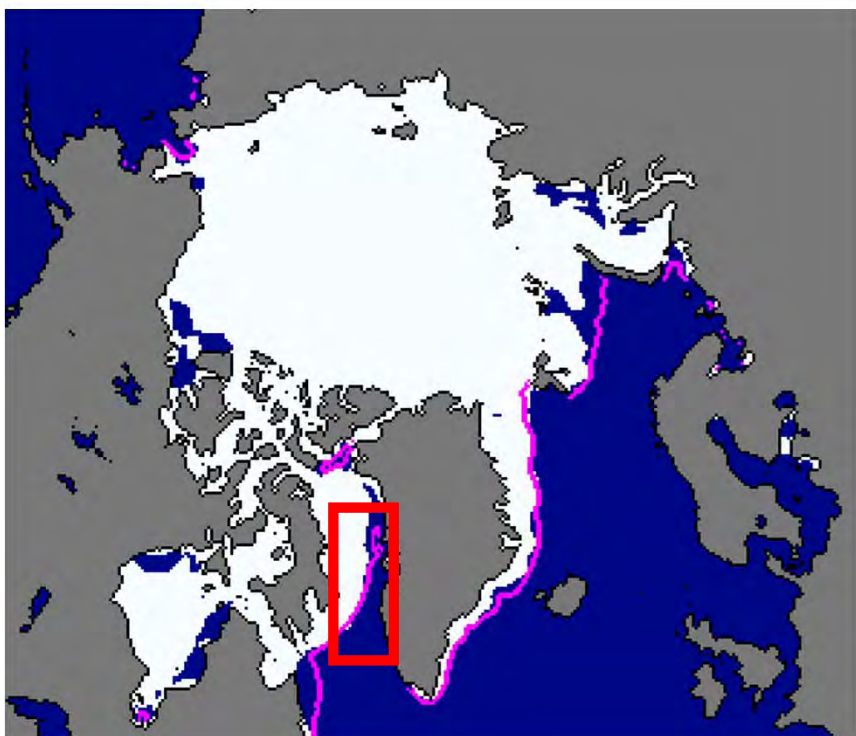
How to monitor phytoplankton
blooms under the ice pack?

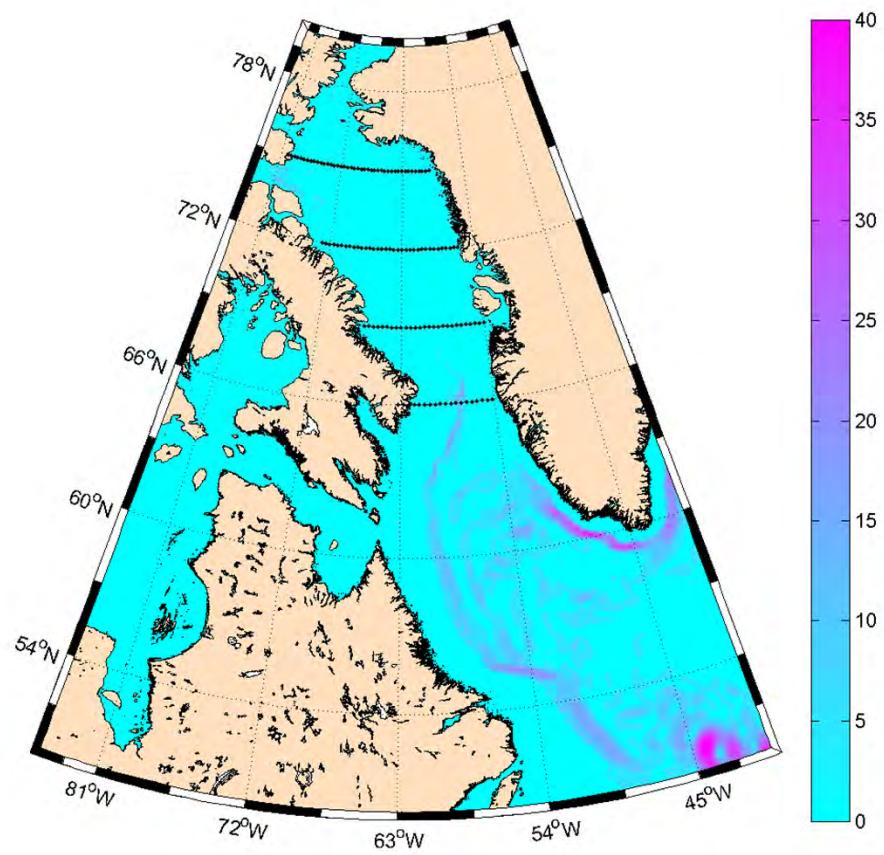


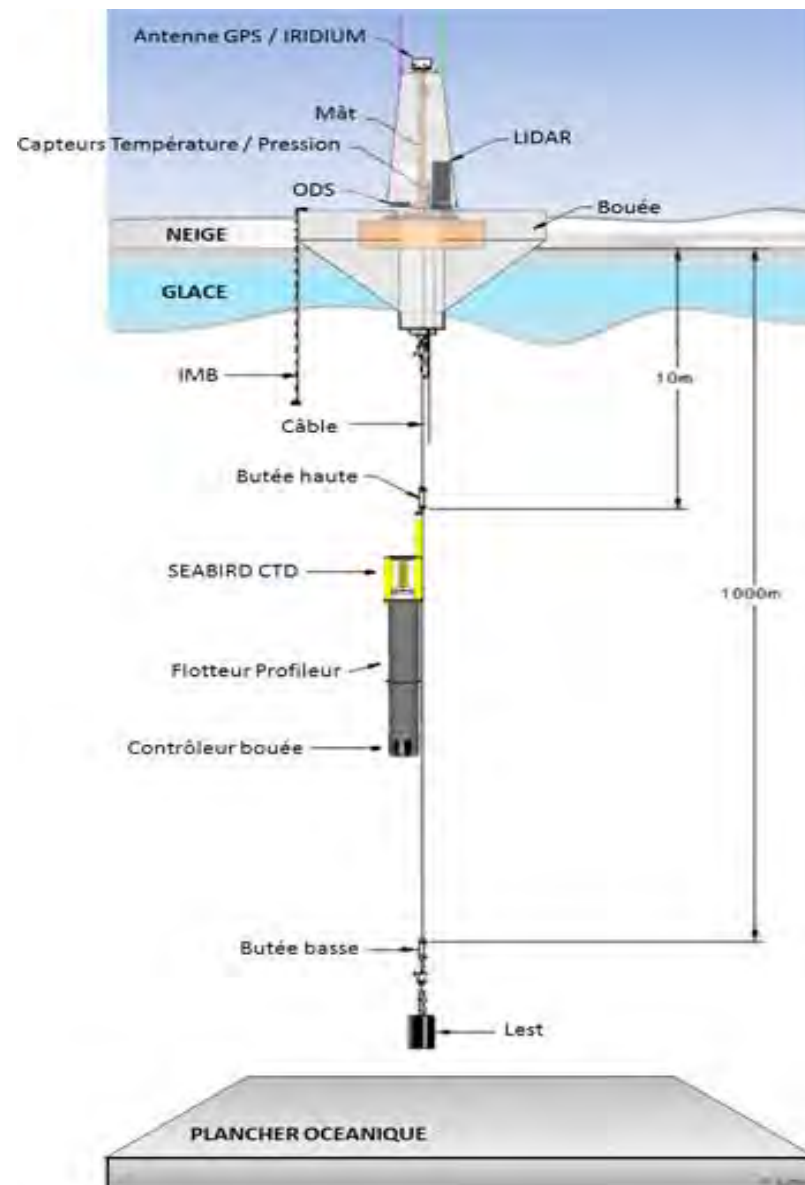


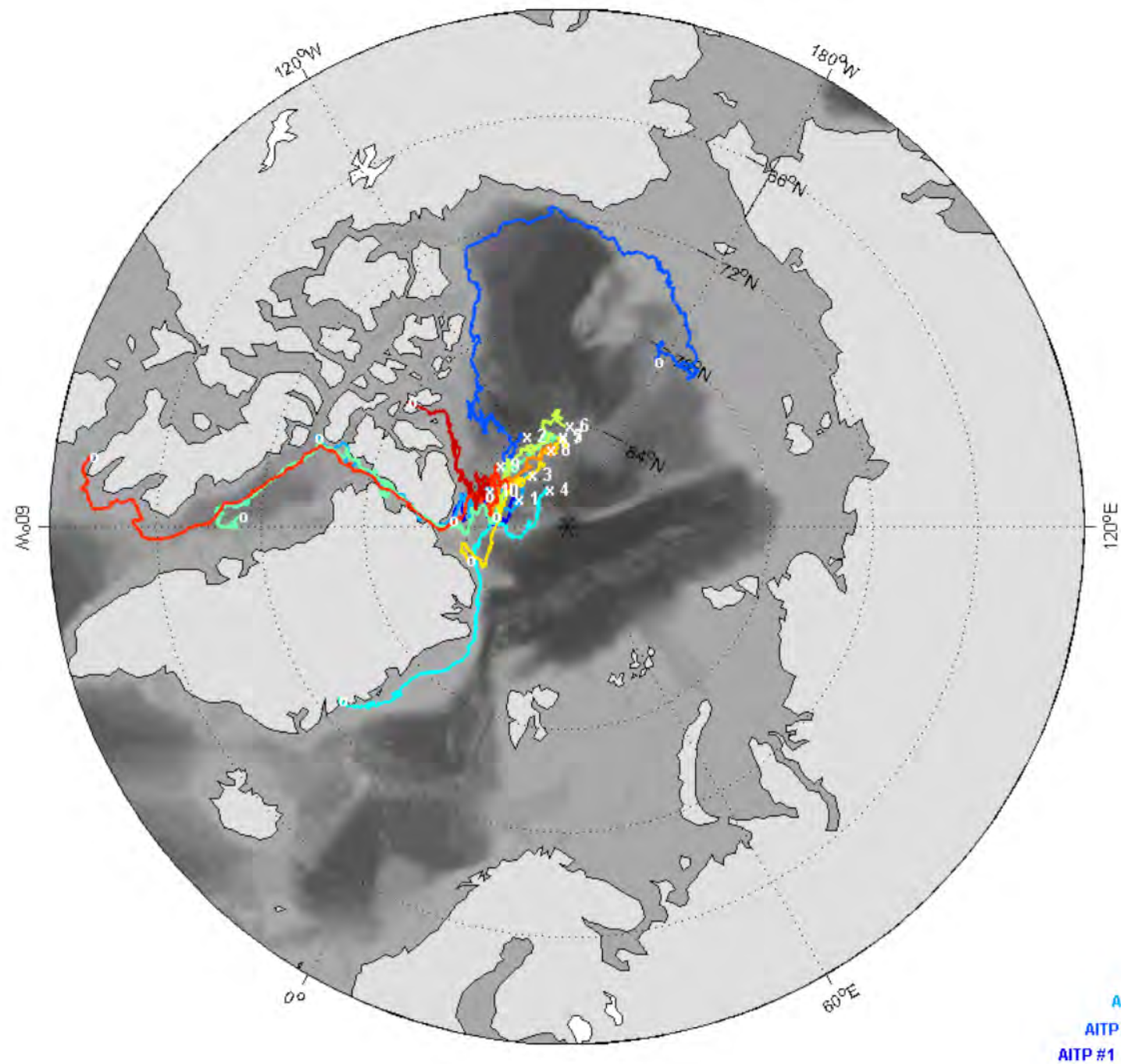












- AITP #10**
- AITP #9**
- AITP #8**
- AITP #7**
- AITP #6**
- AITP #5**
- AITP #4**
- AITP #3**
- AITP #2**
- AITP #1**

END